

WORLD METEOROLOGICAL ORGANIZATION

GLOBAL ATMOSPHERE WATCH

WORLD DATA CENTRE FOR GREENHOUSE GASES



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Acknowledgements

This issue of *Data Summary* reports the latest status of greenhouse and some reactive gases in the global atmosphere. This *Data Summary* has been prepared by the World Data Centre for Greenhouse Gases (WDCGG), established under the Global Atmosphere Watch (GAW) Programme of the World Meteorological Organization (WMO) and operated by the Japan Meteorological Agency (JMA). This *Data Summary* is based on the data submitted by many contributors worldwide (Appendix: LIST OF CONTRIBUTORS). These contributors include both organizations and individuals involved in observations and research of greenhouse and some reactive gases at stations and laboratories operating within the framework of GAW and some other monitoring and research programmes. The WDCGG thanks all of these organizations and individuals, including those from the global air sampling network of the National Oceanic and Atmospheric Administration (NOAA), for their efforts in maintaining the observation programme and continuous provision of observational data. Not all of the contributors may be explicitly acknowledged in this publication, owing to lack of space, but all the organizations and individuals that have submitted data to the WDCGG are nevertheless here acknowledged as invaluable contributors to this latest issue of *Data Summary*.

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SUMMARY

This *Data Summary* reports the results of basic analyses of greenhouse and some reactive gas data submitted to the WMO World Data Centre for Greenhouse Gases (WDCGG) by contributing organizations and individuals. This issue covers observations from 1968 through 2012, based on data reported to the WDCGG by November 2013. The *Data Summary* includes analyses of global, hemispheric and latitudinal monthly mean mole fractions of greenhouse and some reactive gases calculated on the basis of observations at marine and continental surface-based stations, and provides current information on the state of mole fractions of these gases.

Although only monthly mean mole fractions were used for the analyses, the WDCGG greatly appreciates those stations that submit daily and hourly mean mole fractions, which are important for analysis of variations on shorter time scales. All data submitted to the WDCGG are available on its website, <http://ds.data.jma.go.jp/gmd/wdogg/>. In this *Data Summary*, data are reported as dry air mole fractions defined as the number of molecules of a target gas species divided by the number of all molecules in the air including the target itself, but excluding water vapor. Mole fractions are expressed as parts per million (ppm), parts per billion (ppb), and parts per trillion (ppt), which correspond to the SI units of $\mu\text{mol/mol}$, nmol/mol and pmol/mol , respectively.

Variations in the mole fractions of some gases are presented as combinations of seasonal cycles and deseasonalized long-term trends. Growth rates are presented as time derivatives of the long-term trends. Global average mole fractions are presented with accompanying uncertainty. The analytical results are summarized below for each greenhouse and reactive gas.

Carbon Dioxide (CO₂)

The level of carbon dioxide (CO₂), which contributes the most to increases in anthropogenic induced radiative forcing, has been increasing since the beginning of the industrial era. The global average mole fraction of CO₂ reached a new high of 393.1 ± 0.1 ppm in 2012, which constitutes 141% of the pre-industrial level (in 1750). The annual average increase of 2.2 ppm from 2011 to 2012 was greater than the average growth rate for the 1990s (about 1.5 ppm/year) and the average growth rate for the past decade (about 2.0 ppm/year).

The global growth rate of CO₂ has a significant interannual variability driven by natural processes. Interannual changes higher than 2 ppm/year in 1987/1988, 1997/1998, 2002/2003 and 2009/2010

resulted from warmer conditions caused by El Niño-Southern Oscillation (ENSO) events. The anomalously strong El Niño event in 1997/1998 resulted in greater annual increases in CO₂ worldwide in 1998 than during any other one-year period. The high growth rate in 2006 may have been related to the global high temperature during the same year. The exceptionally low growth rate in 1992, including negative values in northern high latitudes, may have been due to low global temperatures following the eruption of Mount Pinatubo in 1991. Variations in CO₂ mole fraction can be seen both on seasonal and long-term scales. The seasonal amplitudes are large in northern high and mid-latitudes and small in the Southern Hemisphere. In southern low latitudes, there is no clear annual cycle, but a semiannual cycle can be determined.

Methane (CH₄)

Methane (CH₄) is the second most significant greenhouse gas which is largely influenced by anthropogenic activity and whose level has been increasing since the beginning of the industrial era. The annual average mole fraction was 1819 ± 1 ppb in 2012, an increase of 6 ppb since 2011. The mole fraction is now 260% of that in the pre-industrial period. This is the sixth year of marked methane increases since levelling-off at the beginning of this century.

The latitudinal gradient of CH₄ mole fraction is large from the northern mid-latitudes to the tropics, suggesting that the major sources of CH₄ are located in the Northern Hemisphere.

CH₄ growth rates decreased significantly in some years, including 1992, when negative values were recorded in northern high and mid-latitudes. However, both hemispheres experienced high growth rates in 1998, caused by higher than average global mean temperature. The global growth rates were generally low from 1999 to 2006, except during the El Niño event of 2002/2003. The global growth rate averaged over the period 1984-1990 was 11.5 ppb/year, but decreased markedly in the 1990s. The mean annual absolute increase during last 10 years was 3.7 ppb/year, but in the last six years through 2012, the global mole fraction increased by a total of 35 ppb.

CH₄ mole fractions vary seasonally, being relatively high in winter and low in summer. Unlike CO₂, the seasonal amplitudes of CH₄ are large, not only in the Northern Hemisphere but also in southern high and mid-latitudes which are associated with methane sinks. In southern low latitudes, a distinct secondary maximum in boreal winter overlies the annual cycle.

Nitrous Oxide (N₂O)

Nitrous oxide (N₂O) is an important greenhouse gas whose level is increasing globally. N₂O data submitted to the WDCGG show that mole fractions are increasing in both hemispheres. The global mean mole fraction reached a new high of 325.1±0.1 ppb in 2012, which is 0.9 ppb higher than that in the previous year. This mole fraction corresponds to 120% of that in the pre-industrial period. The mean annual absolute increase during last 10 years was 0.80 ppb/year and the inter-hemispheric gradient in N₂O is 1.1 ppb (averaged over the years 1980 to 2012), indicating that the majority of N₂O sources are situated in the Northern Hemisphere.

Halocarbons and Other Halogenated Species

Halocarbons, most of which are anthropogenic and generated since the 20th century, are potent greenhouse gases, with some also acting as ozone-depleting compounds. Levels of some halocarbons (e.g. CFCs) increased in the 1970s and 1980s, but this increase has almost ceased by now, due to the production and consumption control of halocarbons under the Montreal Protocol on Substances that Deplete the Ozone Layer and its subsequent Adjustments and Amendments. However, some substances targeted by the Kyoto Protocol but not regulated by the Montreal Protocol, such as HFCs and SF₆, are increasing.

The mole fraction of CFC-11 peaked around 1992 and then started decreasing. The mole fraction of CFC-12 increased until around 2005 and then started decreasing gradually. The mole fraction of CFC-113 stopped increasing in the 1990s, followed by a slight decrease over about twenty years. The mole fractions of HCFCs, used mainly as substitutes for CFCs, have increased significantly during the last decade, but the growth of HCFC-141b decelerated rapidly in the last decade. The mole fraction of Halon-1211 has not increased since 2005, whereas the mole fraction of Halon-1301 is increasing. The mole fraction of CCl₄ was maximal around 1991 and has since decreased slowly. The mole fraction of CH₃CCl₃ peaked around 1992 and decreased thereafter. The mole fractions of HFC-134a, HFC-152a and SF₆ are increasing, but the growth of HFC-152a decelerated in the second half of the decade.

Surface Ozone (O₃)

Ozone (O₃) plays important roles in the atmospheric environment through radiative and chemical processes. It absorbs solar UV radiation in the stratosphere, influencing the vertical temperature profile as well as terrestrial IR radiation, and contributing to the greenhouse effect as a greenhouse gas. Ozone is also involved in the chemical transformations of the primary air pollutants, as its mole fraction in the boundary layer serves as an indicator of air quality.

The mole fraction of O₃ near the surface, so-called surface ozone, reflects various processes. While some of the O₃ in the troposphere comes from the stratosphere, the rest is chemically produced in the troposphere through oxidation of CO or hydrocarbons in the presence of NO_x.

The mole fraction of surface ozone is measured at many locations in various environments. Continuous ozone observations are reported mostly as wet mole fraction. Due to uneven geographic distribution of surface ozone, it is difficult to identify its global long-term trend (WMO, 2011b).

Carbon Monoxide (CO)

Carbon monoxide (CO) is not a greenhouse gas itself but influences the mole fractions of greenhouse gases by affecting hydroxyl radicals (OH). Beginning in 1950, the CO mole fraction increased at a rate of 1% per year but started to decrease in the late 1980s (WMO, 1999). In 2012, the global mean mole fraction of CO was about 92±2 ppb. The mole fraction is high in the Northern Hemisphere and low in the Southern Hemisphere, suggesting substantial anthropogenic emissions in the Northern Hemisphere.

There is a large interannual variability of CO growth rates. The growth rate increases are usually attributed to biomass burning emissions during El Niño conditions.

The monthly mean mole fractions show seasonal variations, with large amplitudes in the Northern Hemisphere and small amplitudes in the Southern Hemisphere with opposite phase.

Nitrogen Monoxide (NO) and Nitrogen Dioxide (NO₂)

Nitrogen oxides (NO_x, *i.e.*, NO and NO₂) are not greenhouse gases, but they are involved in the photochemical production of ozone in the troposphere. In the presence of NO_x, CO and hydrocarbons are oxidized to produce ozone (O₃), which affects the Earth's radiative balance as a greenhouse gas and the oxidization capacity of the atmosphere by reproducing OH.

Most of the stations that have so far reported NO_x data to the WDCGG are located in Europe. NO_x has a large temporal and spatial variability, and it is difficult to identify its long-term global trend based on a spatially limited dataset.

Sulphur Dioxide (SO₂)

Sulphur dioxide (SO₂) is not a greenhouse gas but a precursor of atmospheric sulphate aerosols. Sulphate aerosols are produced by SO₂ oxidation through photochemical gas-to-particle conversion. SO₂ has also been a major source of acid rain and deposition throughout the industrial era.

Most of the stations reporting SO₂ data to the

WDCGG are located in Europe, and it is difficult to identify its long-term global trend based on a spatially limited dataset.

Volatile Organic Compounds (VOCs)

Volatile organic compounds (VOCs) are organic chemicals that easily evaporate or sublime at ordinary atmospheric temperatures (vapor pressure \geq 0.01 kPa at 20 °C). Many are in the form of non-methane hydrocarbons (NMHCs) of different complexity including aliphatics and aromatics which dominate anthropogenic emissions, and unsaturated molecules including terpenes which dominate natural emissions. They also exist as oxygenated hydrocarbons such as acetone and methanol, and sulphur-containing molecules such as dimethyl sulphide.

Although they are not important greenhouse gases in themselves, they are of interest to Global Atmosphere Watch (GAW) because of their environmental impacts, including ozone production and precursors to aerosols. Their main interest to GAW is as tracers of processes which either produce or destroy other major species in the atmosphere measured by GAW. An example is given in chapter 11 of how ethane measurements increase understanding of the behavior of methane.

In its function of the World Data Centre for reactive gases and GHGs, as of November 2013, WDCGG records 53 individual species of VOCs from which a subset, widespread throughout the global atmosphere has been selected as a focus for the GAW VOC Programme. Temporal coverage of VOC measurements has grown extensively over the last decade, since a global flask network came into operation in 2005, supplementing the longer series of measurements made at a few continental sites.

In this 38th edition of the WDCGG *Data Summary*, global analyses are performed for ethane and propane with their relatively long lifetime and wide measurement network. Ethane mole fractions are relatively high in winter and low in summer. The seasonal amplitudes are large in northern high and mid-latitudes but very small in the southern latitudes. Observed seasonal differences are connected with photochemical processes, whereas hemispheric differences indicate the majority of ethane sources in northern latitudes. The propane distribution shows basically similar general features as for ethane with stronger seasonality and more pronounced latitudinal gradients in the Northern Hemisphere (NH) during winter. The stronger seasonality is mainly because of its higher reactivity. Also, due to its shorter atmospheric lifetime compared to ethane, the NH seasonal cycle of propane shows a maximum close to mid-winter.

1. INTRODUCTION

Human activities have had major impacts on the global environment. Since the beginning of the industrial era, mankind has increasingly made use of land, water, minerals and other natural resources, and continuous growth of the world human population and economies may further increase our impact on the environment. As the climate, biogeochemical processes and natural ecosystems are closely interlinked, changes in any one of these may affect the others and be detrimental to humans and other organisms. Emissions of man-made gaseous species and particulate matter alter the energy balance of the atmosphere, which in turn has implications for the multiple interactions within the complex Earth's system. These interactions are not fully understood, partly due to the lack of high quality observations.

The World Meteorological Organization (WMO) established the Global Atmosphere Watch (GAW) Programme in 1989 to promote systematic and reliable observations of the global environment, including but not limited to greenhouse gases (*e.g.*, CO₂, CH₄, CFCs, and N₂O) and some reactive gases (*e.g.*, O₃, CO, VOCs, NO_x, and SO₂) in the atmosphere. In October 1990, WMO designated the Japan Meteorological Agency (JMA) in Tokyo to serve as the World Data Centre for Greenhouse Gases (WDCGG). The WDCGG is responsible for collecting, archiving and providing data on greenhouse and reactive gases in the atmosphere and oceans from a number of observational sites throughout the world that participate in GAW and other scientific monitoring programmes (Appendix: LIST OF OBSERVATIONAL STATIONS). In August 2002, the WDCGG took over the role of the World Data Centre for Surface Ozone from the Norwegian Institute for Air Research (NILU).

With regard to the issue of climate change the Kyoto Protocol to the United Nations Framework Convention on Climate Change came into force in February 2005. In March 2006, WMO commenced annual publication of the WMO Greenhouse Gas Bulletin, which summarizes the state of greenhouse gases in the atmosphere. The ninth issue of the Bulletin was published in November 2013. The WDCGG contributes to the production of the Bulletin through timely and adequate collection and analysis of data in cooperation with the contributors of the data.

Since its establishment, the WDCGG has provided its users with data and other information through its regular publications, including the *Data Summary* and *DVD* (Appendix: LIST OF WMO WDCGG PUBLICATIONS). In accordance with the GAW Strategic Plan: (2008–2015) and its Addendum, all data and information have been made available on the WDCGG web site, improving the accessibility of data, information and products (WMO, 2007a; WMO, 2011a). The WDCGG published the Data Submission and Dissemination Guide in 2007 (WMO, 2007b), which,

with its revision in 2009 (WMO, 2009b), is designed to facilitate submission of observational data and provide access to archived data in the WDCGG. Clear guidelines for data submission are included in the measurement guidelines published by GAW for the variables, which are under responsibility of WDCGG.

The GAW Strategic Plan requests that World Data Centres assist data users by providing the data and analysis related to atmospheric observations. To this end, the WDCGG provides global and integrated diagnostics on the state of greenhouse and some reactive gases as analytical information in the *Data Summary*. The WDCGG global analysis method has been described in a GAW technical report (WMO, 2009a). The content of the *Data Summary* is revised and improved based on comments from data contributors and scientists. We hope the diagnostic information presented here will promote the use of data on greenhouse and reactive gases and will enhance appreciation of the value of the GAW Programme.

All users are required to accept the following statement endorsed by the Commission for Atmospheric Sciences (CAS) at its thirteenth session: "For scientific purposes, access to these data is unlimited and provided without charge. By their use you accept that an offer of co-authorship will be made through personal contact with the data providers or owners whenever substantial use is made of their data. In all cases, an acknowledgement must be made to the data providers or owners and to the data centre when these data are used within a publication." The WDCGG requests data users to make appropriate acknowledgements. The principal investigators and other contacts can be obtained from the WDCGG website, as well as from the GAW Station Information System (GAWSIS) website, <http://gaw.empa.ch/gawsis/>. Information on these websites is updated in cooperation with the data contributors and the WMO Secretariat.

Finally, the WDCGG would like to thank all data contributors worldwide, including those involved in on-site measurements, for their efforts in maintaining the observational programmes and for continuous data provision.

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2. ANALYSIS

The WDCGG reposit, archives and provides observational data on the mole fractions of greenhouse and some reactive gases, and publishes diagnostic information on these gases based on the reported data.

The long-term trends and seasonal variations in the mole fractions of CO₂, CH₄, N₂O and CO are calculated for the whole globe (global means) and for latitudinal belts (zonal means). Global long-term trends in the surface O₃, are not analyzed due to its substantial spatial gradients, and its uneven geographic distribution which is poorly covered by observational sites. Zonal mean time series for selected VOCs (ethane and propane) are newly introduced. For halocarbons, NO_x and SO₂, only monthly mean mole fractions over time are presented without global, hemispheric or zonal averaging, due to insufficient number of reporting sites for each compound.

Mole fractions are expressed as parts per million (ppm), parts per billion (ppb), and parts per trillion (ppt), which correspond to the SI units of $\mu\text{mol/mol}$, nmol/mol and pmol/mol , respectively.

The method of analysis for CO₂, CH₄, N₂O and CO is summarized below. The details of the global analysis method are provided in the *Technical Report of Global Analysis Method for Major Greenhouse Gases by the World Data Centre for Greenhouse Gases*, published as a GAW technical report (WMO, 2009a). Additional uncertainty can be expected in the result of CO global analysis due to diversity of scales. When assessing long-term trends for CO₂, CH₄ and N₂O, the growth rates at both ends of the period were assumed to be simple linear extensions of the adjacent year, thus avoiding end effects. For simplicity, the rates for the rest of the period were approximated by linear functions.

(1) Site selection

For CO₂, CH₄ and N₂O, the diagnostic analyses, including global, hemispheric and zonal means, were based on data from sites that have adopted a standard scale traceable to the Primary Standard designated by WMO. These analyses also utilize data on other standard scales that are convertible to the WMO scale through a proven equation. Letters informing data submitters of the most recent WMO scales are sent out regularly by the WDCGG as well as discussed at the regular expert meetings (WMO, 2012).

Selection of observational sites is also based on whether they provide data representing a reasonably large geographical area, considering the fact that some sites may be susceptible to local sources and sinks. Sites are selected objectively using data submitted to the WDCGG. For CO₂, CH₄ and CO, only those sites that provide annual mean mole fractions falling within

a range of $\pm 3\sigma$ from a curve fitted to the LOESS model curve (Cleveland and Devlin, 1988) have been selected, with outliers rejected in an iterative manner. This procedure does not affect the datasets residing in the WDCGG, and these data may be useful for purposes other than global analysis, such as identification of sources and sinks.

The sites selected according to the above criteria are marked with asterisks in Plate 3.1 for CO₂, Plate 4.1 for CH₄, Plate 5.1 for N₂O and Plate 8.1 for CO, which represent 136 (72%), 127 (78%), 56 (70%) and 118 (87%) of the submitted datasets respectively (detailed in 'LIST OF OBSERVATIONAL STATIONS' in this issue).

(2) Analysis of long-term trends

The mole fractions of greenhouse and reactive gases over time, measured under unpolluted conditions, exhibit variations on different time scales. The two major components are seasonal variations and long-term trends. Several attempts have been made to separate these various scales from the measured data, including objective curve fitting (Keeling *et al.*, 1989), digital filtering (Thoning *et al.*, 1989; Nakazawa *et al.*, 1991), or both (Conway *et al.*, 1994; Dlugokencky *et al.*, 1994).

In this report, seasonal variations derived from components of Fourier harmonics and long-term trends are extracted by low-pass filtering with a cut-off frequency of 0.48 year^{-1} for each selected site. Details are described in WDCGG *Data Summary* No. 22 (WMO, 2000).

(3) Estimation for missing periods and gaps

The number and distribution of sites used to assess trends during the analysis period should be kept as invariable as possible to avoid the effects of changes in the availability of data over time. However, only a small number of sites provided data throughout the entire analysis period; others may have covered shorter periods or had gaps in measurements due to different reasons. To use as many sites as possible, data for missing values are constructed using interpolation and extrapolation in the calculation of zonal means as described below.

Gaps in some data were filled by linear interpolation based on available data, by subtracting the seasonal variation calculated from the longest consecutive period of data with Lanczos filters (Duchon, 1979). The subtracted variation was added back to the data to obtain estimated mole fractions in a single sequence.

In the case of extrapolation, long-term trends from the existing or interpolated series of data were extrapolated based on zonal mean growth rates

calculated from other sites in the same latitudinal zone. The seasonal variation was added to the extrapolated long-term trend to obtain estimated mole fractions for the entire period of analysis.

Using these statistical procedures, the future addition of new stations should not affect the consistency in global estimates over time.

Nevertheless, while adding new sites in the analysis WDCGG performs calculations both with and without the new information to ensure the consistency of the global average calculations.

(4) Calculation of global, hemispheric and zonal means

Zonal means were calculated by determining the arithmetic average of the mole fractions in each latitudinal zone, based on consistent datasets derived as above.

Global and hemispheric means were calculated as the weighted averages of the zonal means taking account of the area of each latitudinal zone.

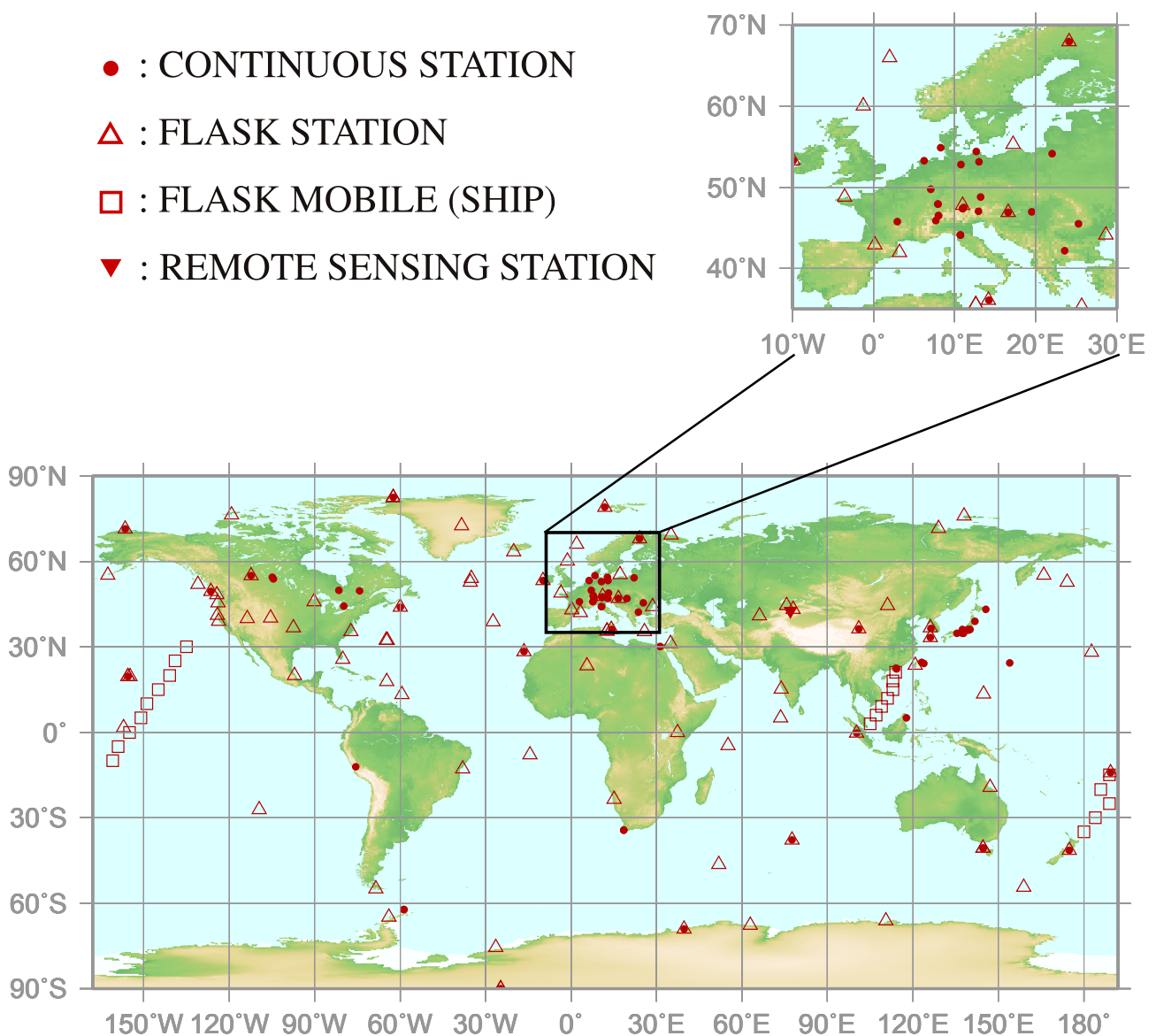
Deseasonalized long-term trends and growth rates for the globe, each hemisphere and each latitudinal zone were calculated from the global, hemispheric and zonal means, respectively, using the low-pass filter mentioned above and the time derivatives after filtering.

Error ranges estimated by a bootstrap method (Conway *et al.*, 1994) are included with the global means of major GHGs (CO₂, CH₄, N₂O and CO), where uncertainty is estimated as the standard deviation of many global means calculated by each bootstrap network.

3.

CARBON DIOXIDE (CO₂)

- : CONTINUOUS STATION
- △ : FLASK STATION
- : FLASK MOBILE (SHIP)
- ▼ : REMOTE SENSING STATION



This map shows locations of the stations that have submitted data for monthly mean mole fraction.

CO₂ Monthly Data

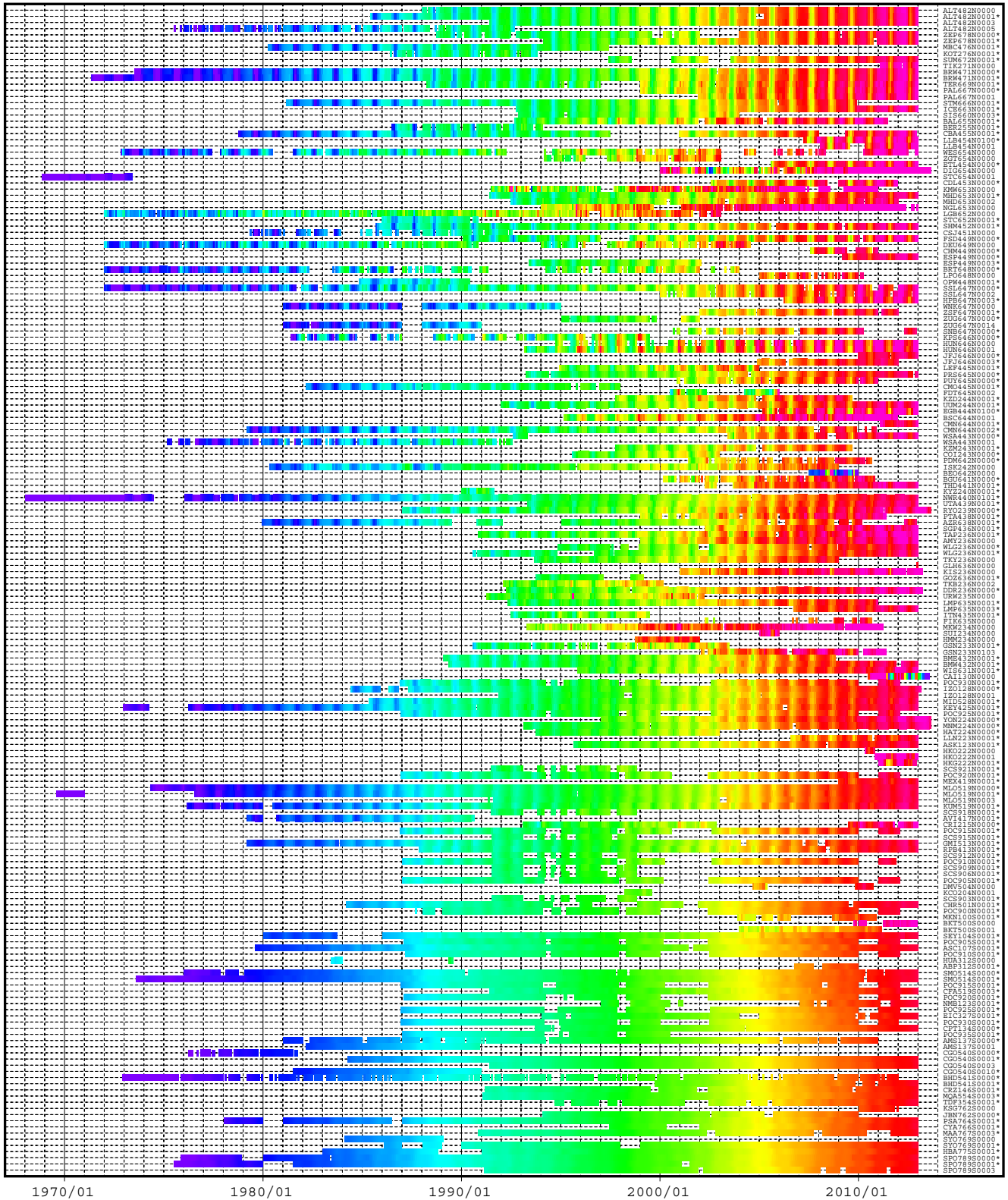
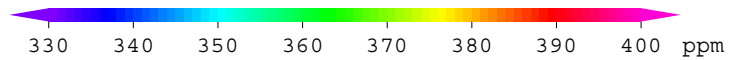
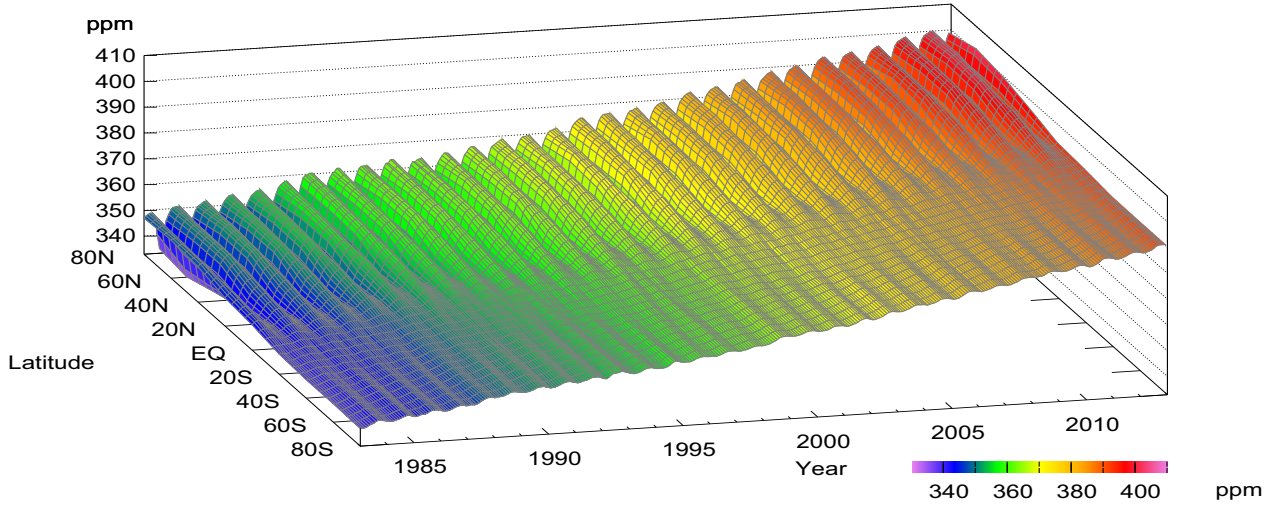
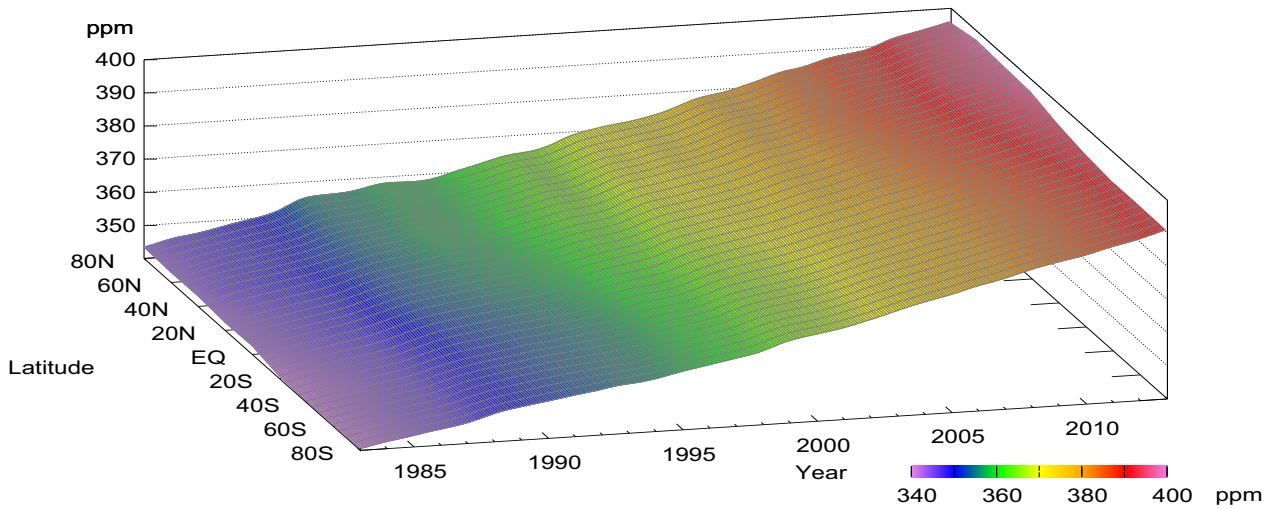


Plate 3.1 Monthly mean CO₂ mole fractions that have been reported to the WDCGG. The mole fractions are illustrated in different colours. The sites are listed in order from north to south. In the case where data are reported for two or three different altitudes, only the data at the highest altitudes are illustrated. In the case where monthly means are not reported, the WDCGG calculates them from hourly or other mole fractions reported to the WDCGG by simple arithmetic mean. The data from the sites with an asterisk at the end of the station index are used for the analysis shown in Plate 3.2. (see Chapter 2)

CO₂ mole fraction



CO₂ deseasonalized mole fraction



CO₂ growth rate

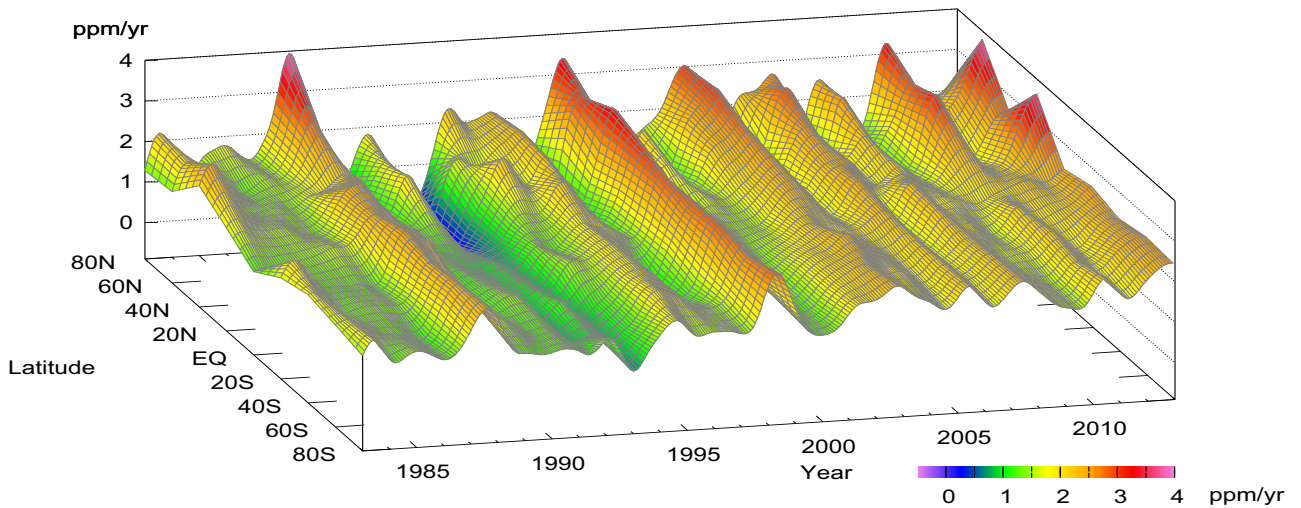


Plate 3.2 Variation of zonally averaged monthly mean CO₂ mole fractions (top), deseasonalized long-term trends (middle), and growth rates (bottom). The zonally averaged mole fractions are calculated for each 20° zone. The deseasonalized trends and growth rates are derived as described in Chapter 2.

3. CARBON DIOXIDE (CO₂)

Basic information on CO₂ with regard to environmental issues

Carbon dioxide (CO₂) has strong absorption bands in the infrared region and is the biggest anthropogenic contributor to the man-made greenhouse effect. CO₂ accounts for about 64% of total increase in the radiative forcing (since 1750) due to long-lived greenhouse gases in the atmosphere (WMO, 2013). It is responsible for 84% of the increase in radiative forcing over the past decade and 82% over the past five years.

The balance of the fluxes between the atmosphere, the oceans and the biosphere determines the mole fraction of CO₂ in the atmosphere. An amount of 515 [445 to 585] PgC was emitted between 1870 to 2011 (IPCC, 2013) and annual anthropogenic emissions mainly due to fossil fuel combustion reached 9.5 ± 0.5 gigatonnes in 2011 (<http://www.globalcarbonproject.org/>). Carbon in the atmosphere is exchanged with two other large reservoirs, the terrestrial biosphere and the oceans. CO₂ exchanges between the atmosphere and terrestrial biosphere occur mainly through absorption by photosynthesis and emission from the respiration of plants and the decomposition of organic soils. These biogenic activities vary seasonally, resulting in large seasonal variations in the level of CO₂. The direction of CO₂ exchange between the atmosphere and oceans is determined by the gradient of CO₂ mole fraction, and varies in time and space.

The current mole fractions of atmospheric CO₂ far exceed historic records, dating back at least the 2.1 million years (Tans, 2009). Based on the results of ice core studies, the mole fraction of atmospheric CO₂ in pre-industrial times was about 278 ppm (IPCC, 2013). The emission of CO₂ due to human activities has increased dramatically since the beginning of the industrial era, impacting CO₂ exchange rates between different reservoirs and CO₂ levels not only in the atmosphere but in the oceans and terrestrial biosphere. The global carbon cycle, which is comprised mainly of CO₂, is not fully understood. About half of anthropogenic CO₂ emissions has remained in the atmosphere, with the remainder removed by sinks, including the terrestrial biosphere and oceans. However, the amount of CO₂ removed from the atmosphere varies significantly over time (Figure 3.1) without confirmed trend (Levin, 2012).

Carbon isotopic studies have shown the importance of the terrestrial biosphere and oceans as sources and sinks of CO₂ (Francey *et al.*, 1995; Keeling *et al.*, 1995; and Nakazawa *et al.*, 1993, 1997). In contrast, the atmospheric content of O₂ depends primarily on its removal by the burning of fossil fuels and on its release

from the terrestrial biosphere. Therefore, the uptake of carbon by the terrestrial biosphere and oceans can be estimated from the combination of measurements of O₂ (O₂/N₂) and CO₂ (Manning and Keeling, 2006). A quasi-equilibrium amount of CO₂ is expected to be retained in the atmosphere by the end of the millennium that is surprisingly large: typically 40% of the peak concentration enhancement over pre-industrial values (278 ppm) (Solomon *et al.*, 2009).

Large amounts of CO₂ are exchanged among the reservoirs in nature, and the global carbon cycle is coupled with the climate system on seasonal, yearly and decadal time scales. Complete understanding of the global carbon cycle is essential for estimating future CO₂ mole fractions in the atmosphere.

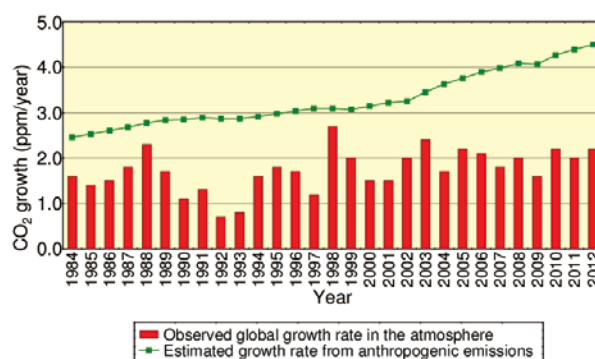


Fig. 3.1 Annual mean growth rates of CO₂ in the atmosphere, calculated from observational data (red columns) and from data for anthropogenic emissions (green curve). The estimated growth rates were calculated using CO₂ emissions as a proxy (from CDIAC, Boden *et al.*, 2013). The values from 2011 to 2012 are estimations of Carbon Dioxide Information Analysis Center (CDIAC), expressed as moles divided by the total mass of gas in the atmosphere (5.2 petatonnes) converted to moles based on the mean molar mass of dry air (about 29.0 g/mol). The observed growth rates were calculated by the WDCGG. The observed CO₂ abundance is expressed as mole fraction with respect to dry air, while the CO₂ amount calculated from anthropogenic emissions is based on the atmosphere, including water vapor, usually in a fraction less than 1%.

Mole fractions of CO₂ can be analyzed utilizing data submitted to the WDCGG from fixed stations and some ships. The observational sites from which data were used for the analysis are shown on the map at the beginning of this chapter. They include fixed stations performing continuous measurements as well as flask-sampling stations, including those in the NOAA/ESRL cooperative air sampling network. In

addition, mobile platforms such as ships and aircrafts and other stations observing on an event basis report their data to the WDCGG (see Appendix: LIST OF OBSERVATIONAL STATIONS), which are not used for global analysis.

Annual variation of CO₂ mole fraction in the atmosphere

The monthly mean mole fractions of CO₂ used in the analysis are shown in Plate 3.1, with mole fraction levels illustrated in different colours. Global, hemispheric and zonal mean mole fractions were analyzed based on data from selected stations under unpolluted conditions (see the caption on Plate 3.1). Zonally averaged mole fractions of atmospheric CO₂, together with their deseasonalized components and growth rates, are shown as three-dimensional representations in Plate 3.2. These plots show that the seasonal variations in mole fraction are large in northern high and mid-latitudes, but are indistinct in the Southern Hemisphere. The increases in the Northern Hemisphere precede those in the Southern Hemisphere by one or two years, and the interannual variations in growth rate are larger in the Northern Hemisphere.

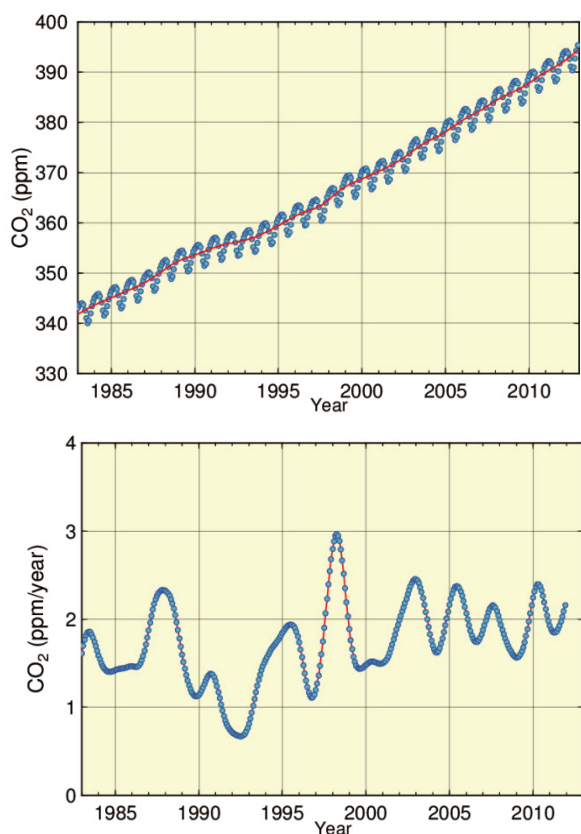


Fig. 3.2 Global monthly mean mole fraction of CO₂ from 1983 to 2012, including deseasonalized long-term trend shown as a red line (top) and annual growth rate (bottom).

Figure 3.2 shows global monthly mean CO₂ mole fractions and their growth rates from 1983 to 2012. The global average mole fraction reached a new high of 393.1 ± 0.1 ppm in 2012, which is 141% of the pre-industrial level of 278 ppm. The 2.2 ppm annual increase in 2011-2012 was greater than the average growth rate for the 1990s (about 1.5 ppm/year) and that of the past decade (about 2.0 ppm/year).

The global growth rate shows large interannual variations, with an instantaneous maximum of about 3 ppm/year in 1998 and a minimum below 1 ppm/year in 1992. There were short periods of high rates in 1987/1988, 1997/1998, 2002/2003, 2005/2006, 2007 and 2009/2010.

Figure 3.3 shows monthly mean mole fractions and long-term trends from 1983 to 2012 for each 30° latitudinal zone, indicating that there were clear long-term increases in both hemispheres and seasonal variations in the Northern Hemisphere.

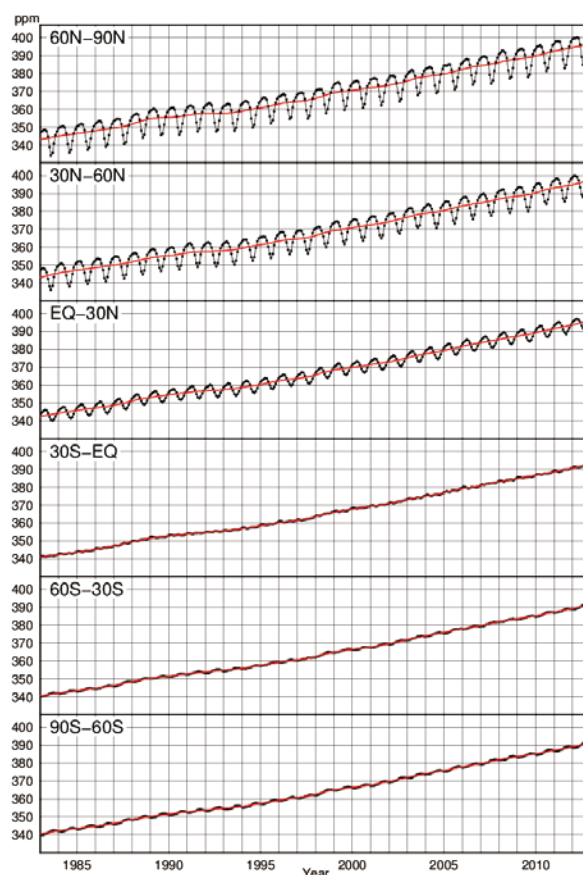


Fig. 3.3 Monthly mean mole fractions of CO₂ from 1983 to 2012 for each 30° latitudinal zone (dots) and their deseasonalized long-term trends (red lines).

As shown in Figure 3.4, the growth rates for each 30° latitudinal zone fluctuated between -0.3 and 3.6 ppm/year, with the largest interannual variability in northern high latitudes. High growth rates for all 30°

latitudinal zones were observed in 1987/1988, 1997/1998, 2002/2003, 2005, 2007, and 2010, with negative rates recorded in northern high latitudes in 1992.

Changes in growth rate are partly associated with ENSO. The El Niño events in 1982/1983, 1986–1988, 1991/1992, 1997/1998 and 2002/2003 coincided with high growth rates of CO₂, with an exception in 1992. The growth rates of CO₂ observed by aircraft at high altitudes (8–13 km) over the Pacific Ocean were also associated with ENSO (Matsueda *et al.*, 2002). The latest El Niño event occurred in 2009/2010.

During El Niño events, the up-welling of CO₂-rich ocean water in the eastern equatorial Pacific is suppressed, resulting in reduced CO₂ emissions from this area. El Niño events induce high temperature anomalies in many areas, particularly in the tropics, resulting in increased CO₂ emissions from the terrestrial biosphere due to the enhanced respiration of plants and activated decomposition of organic matter in soil (Keeling *et al.*, 1995). This effect is enhanced by the suppression of plant photosynthesis in areas of anomalously low precipitation, particularly in the tropics. These oceanic and terrestrial processes during El Niño events have opposing effects, but Heimann and Reichstein (2008) suggested that the latter was the main cause of the variation in the CO₂ growth rate.

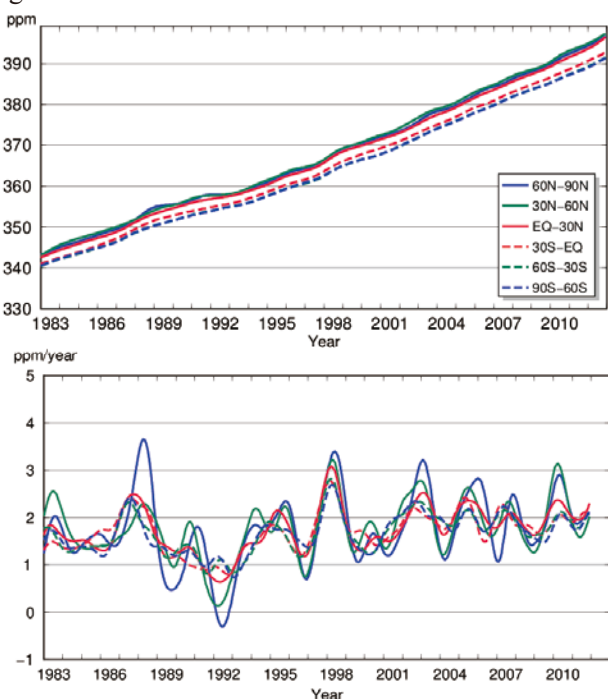


Fig. 3.4 Long-term trends in the mole fraction of CO₂ for each 30° latitudinal zone (top) and their growth rates (bottom) .

However, an exceptionally low CO₂ growth rate occurred during the El Niño event in 1991/1992. The injection of 14 - 20 megatonnes (Mt) of SO₂ aerosols into the stratosphere by the Mount Pinatubo eruption in June 1991 affected the radiation budget and atmospheric circulation (Hansen *et al.*, 1992; Stenchikov *et al.*, 2002), resulting in a drop in global temperature. Angert *et al.* (2004) suggested that the low CO₂ growth rate observed during this El Niño event was due to reduced CO₂ emissions caused by consequent changes in the respiration of terrestrial vegetation and the decomposition of organic matter (Conway *et al.*, 1994; Lambert *et al.*, 1995; Rayner *et al.*, 1999), and by enhanced CO₂ absorption due to intensive photosynthesis caused by an increase in diffuse radiation (Gu *et al.*, 2003).

Seasonal cycle of CO₂ mole fraction in the atmosphere

Figure 3.5 shows average seasonal cycles in the mole fraction of CO₂ for each 30° latitudinal zone. The seasonal cycles are clearly large in amplitude in northern high and mid-latitudes and small in the Southern Hemisphere. The seasonal cycle in the Northern Hemisphere is mainly dominated by the land biosphere (Nevison *et al.*, 2008), and it is characterized by rapid decreases from June to August and large returns from September to December.

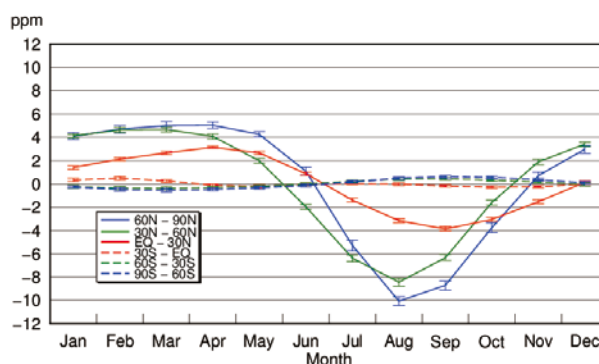


Fig. 3.5 Average seasonal cycles in the mole fraction of CO₂ for each 30° latitudinal zone obtained by subtracting long-term trends from the zonal mean time series. Vertical error bars represent the range of $\pm 1\sigma$ which is calculated for each month. (Average from 1983 to 2012)

The mole fractions of CO₂ in northern low latitudes lagged behind that in high latitudes by one or two months. Minimum values appeared in August in northern high and mid-latitudes and in September in northern low latitudes.

In the Southern Hemisphere, seasonal variations showed small amplitudes with a half-year delay due to small amounts of net emission and absorption by the terrestrial biosphere. Seasonal variations in both

northern and southern mid-latitudes were apparently superimposed in southern low latitudes (0–30°S). The direct influence of sources and sinks in the Southern Hemisphere may be partially cancelled by the propagation of an antiphase variation from the Northern Hemisphere.

Figure 3.6 shows latitudinal distributions of the mole fractions of CO₂ in January, April, July and October 2012, from sites marked with an asterisk in Plate 3.1. In latitudes north of 30°N, the mole fractions increased towards higher latitudes in January and April, and decreased towards higher latitudes in July, corresponding to the large seasonal variations in northern high and mid-latitudes, variations associated with activities of the terrestrial biosphere.

It can be noticed that during seasonal maximum many stations of the Northern Hemisphere observed monthly mean CO₂ mole fractions above the “symbolic” threshold of 400 ppm.

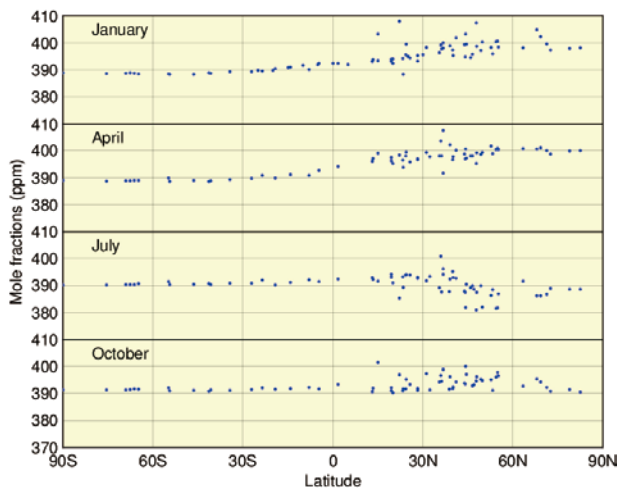


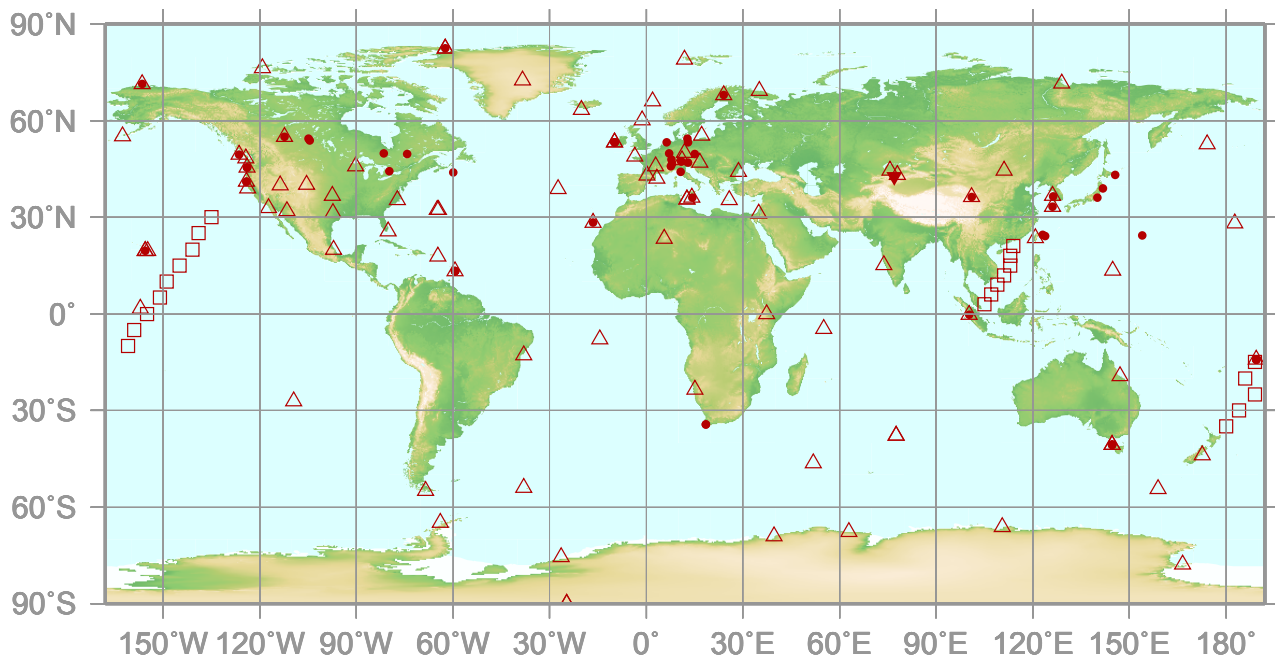
Fig. 3.6 Latitudinal distributions of the monthly mean mole fractions of CO₂ in January, April, July and October 2012.

4.

METHANE

(CH₄)

- : CONTINUOUS STATION
- △ : FLASK STATION
- : FLASK MOBILE (SHIP)
- ▼ : REMOTE SENSING STATION



This map shows locations of the stations that have submitted data for monthly mean mole fraction.

CH₄ Monthly Data

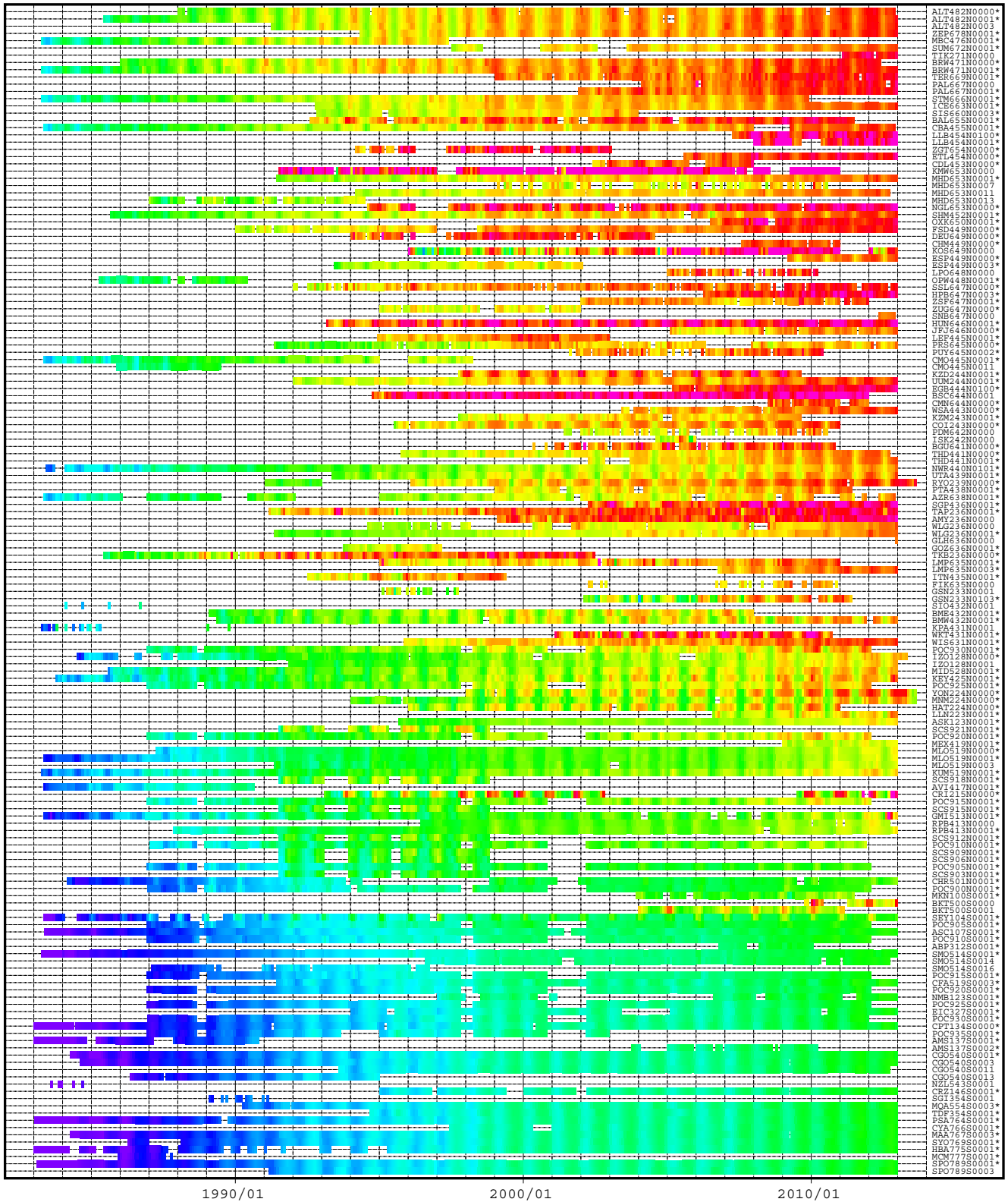
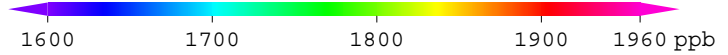
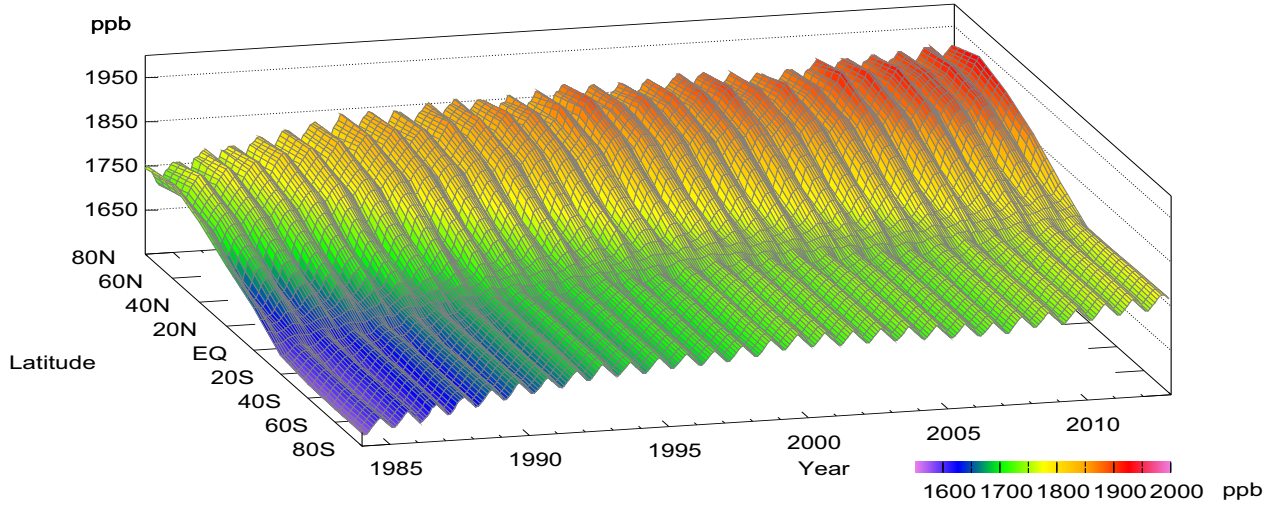
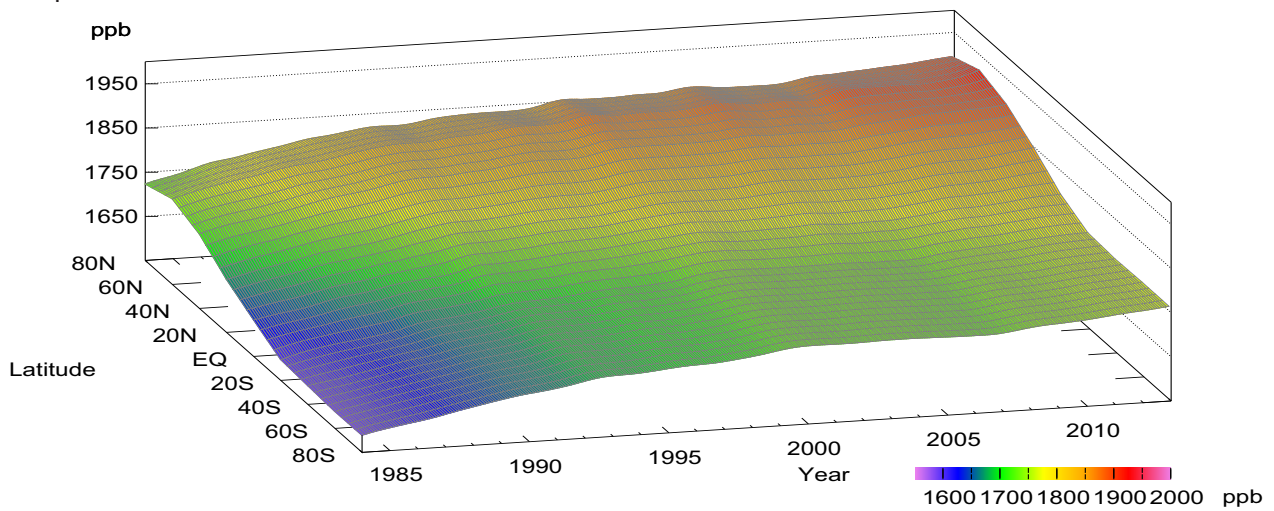


Plate 4.1 Monthly mean CH₄ mole fractions that have been reported to the WDCGG. The mole fractions are illustrated in different colours. The sites are listed in order from north to south. In the case where data are reported for two or three different altitudes, only the data at the highest altitudes are illustrated. In the case where monthly means are not reported, the WDCGG calculates them from hourly or other mole fractions reported to the WDCGG by simple arithmetic mean. The data from the sites with an asterisk at the end of the station index are used for the analysis shown in Plate 4.2. (see Chapter 2)

CH₄ mole fraction



CH₄ deseasonalized mole fraction



CH₄ growth rate

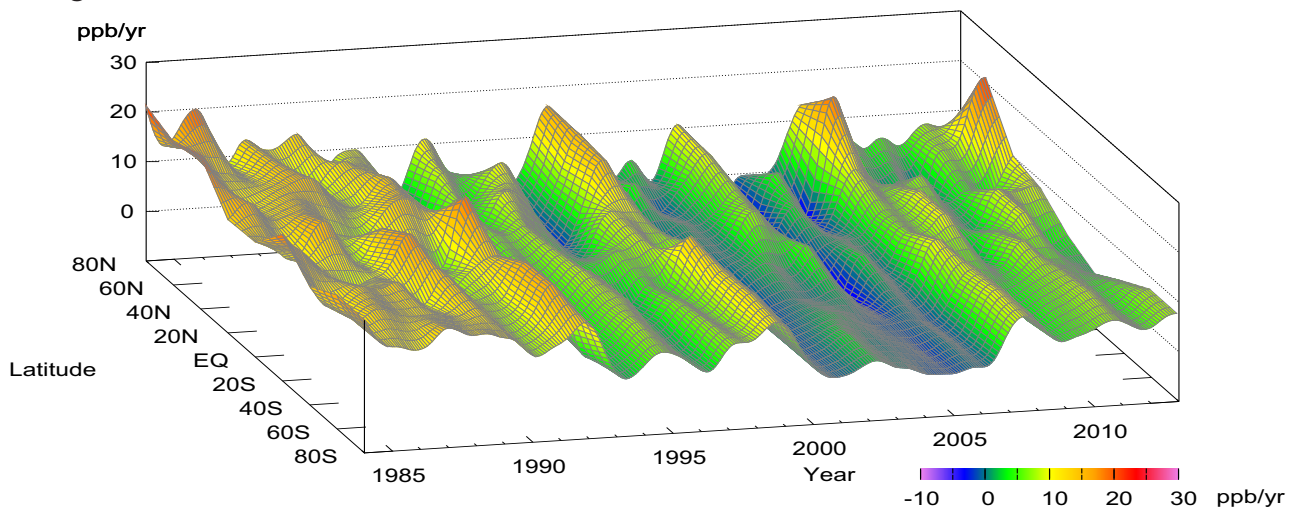


Plate 4.2 Variation of zonally averaged monthly mean CH₄ mole fractions (top), deseasonalized long-term trends (middle), and growth rates (bottom). The zonally averaged mole fractions are calculated for each 20° zone. The deseasonalized trends and growth rates are derived as described in Chapter 2.

4. METHANE (CH₄)

Basic information on CH₄ with regard to environmental issues

Methane (CH₄) is the second most important anthropogenic greenhouse gas, with an estimated global warming potential per molecule 25 times greater over a 100 year horizon and 72 times greater over a 20 years horizon than CO₂. Between 1750 and 2012, CH₄ accounted for about 18% of the total increase in radiative forcing due to long-lived greenhouse gases in the atmosphere (WMO, 2013).

Analyses of air trapped in ice cores from Antarctica and the Arctic revealed that the current atmospheric CH₄ mole fraction is the highest over the last 650,000 years (Solomon *et al.*, 2007). The mole fraction of CH₄ remained at about 700 ppb from 1000 A.D. until the start of the industrial era when it started increasing. Measurements in ice cores have shown that inter-polar differences in CH₄ mole fractions between Greenland and Antarctica ranged from 24 to 58 ppb between 1000 and 1800 A.D. (Etheridge *et al.*, 1998). However, the inter-hemispheric difference of the mole fractions (see Fig 4.3) averaged over the years 1984 to 2012 reached the value of 135 ppb. Increase in the inter-polar gradient reflects the dominant impact of the emissions from the Northern Hemisphere, where major anthropogenic and natural sources are situated.

CH₄ is emitted by both natural and anthropogenic sources, including natural wetlands, oceans, landfills, rice paddies, enteric fermentation, fossil fuel production and consumption and biomass burning. The global emission of CH₄ was 548 teragrams (Tg) CH₄ per year, with more than 60% related to anthropogenic activities (<http://www.globalcarbonproject.org/methanebudget/13/hl-compact.htm>). CH₄ is removed from the atmosphere by reaction with hydroxyl radicals (OH) in both the troposphere and stratosphere, and by reaction with chlorine atoms and O(¹D), an excited state of oxygen, in the stratosphere. CH₄ is one of the most important sources of water vapor in the stratosphere and has an atmospheric lifetime of about 10 years. More information regarding sources and sinks of CH₄ must be collected to better understand the budget of atmospheric CH₄.

Mole fractions of CH₄ are analyzed using data submitted to the WDCGG from fixed stations and some ships. These observational sites are shown on the map at the beginning of this chapter.

Annual variation of CH₄ mole fraction in the atmosphere

The monthly mean dry mole fractions of CH₄ used in this analysis are shown in Plate 4.1, with the mole fraction levels illustrated in different colours. Global,

hemispheric and zonal mean mole fractions have been calculated based on data from selected stations under unpolluted conditions (see the caption for Plate 4.1). Zonally averaged atmospheric CH₄ mole fractions, together with their deseasonalized components and growth rates, are shown as three-dimensional representations in Plate 4.2. These plots show that the seasonal variations in CH₄ mole fraction are larger in the Northern than in the Southern Hemisphere and that the increase in the Northern Hemisphere propagates to the Southern Hemisphere. The growth rates vary on a global scale with the patterns similar to those for CO₂ (see Section 3). There is a large latitudinal gradient in CH₄ mole fraction from the northern mid-latitudes to the tropics, suggesting major sinks in the tropics, where the mole fraction of OH radicals is higher.

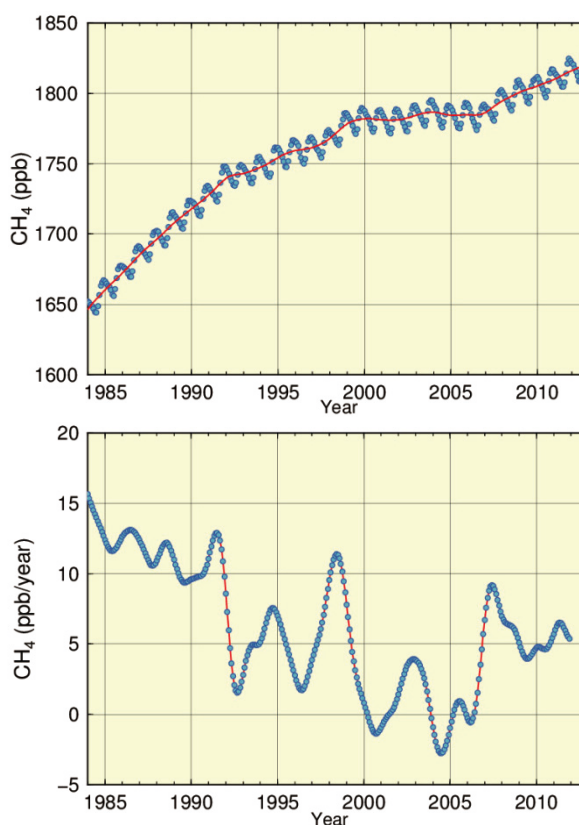


Fig. 4.1 Global monthly mean mole fraction of CH₄ from 1984 to 2012, including deseasonalized long-term trend in red line (top) and annual growth rate (bottom).

Figure 4.1 shows globally averaged monthly mean mole fractions and the growth rates for CH₄ from 1984

to 2012. The global average mole fraction was 1819 ± 1 ppb in 2012, an increase of 6 ppb from 2011. The mole fraction did not change much between 1999 and 2006. The mean annual absolute increase during last 10 years was 3.7 ppb/year. The current mole fraction is 260% of its pre-industrial level of 700 ppb.

Figure 4.2 shows monthly mean mole fractions from 1984 to 2012 for each 30° latitudinal zone. The smallest magnitude of the seasonal variations is observed in the latitudinal zone between the equator and 30° S.

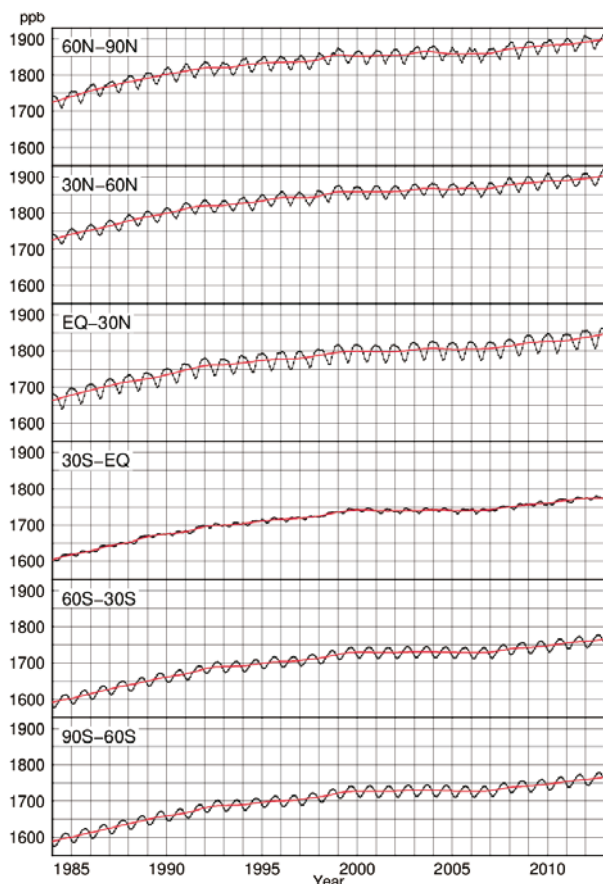


Fig. 4.2 Monthly mean mole fractions of CH_4 from 1984 to 2012 for each 30° latitudinal zone (dots) and their deseasonalized long-term trends (red lines).

Figure 4.3 summarizes deseasonalized long-term trends for each 30° latitudinal zone and their growth rates. A latitudinal gradient between the northern high and mid-latitudes with higher mole fractions and the southern latitudes with lower mole fractions is clearly seen in Fig. 4.3, while mole fractions in most latitudinal belts have similar tendencies. In the 1990s, the growth rates clearly decreased in all latitudinal zones, but remained positive. The declined growth rate was especially evident during the second half of 1992, in 1996, and almost even in 1999 and in 2004/2005, when growth rates were less than 5 ppb/year in all latitudes. During the year 1998, the

maximum global growth rate reached about 11 ppb/year (Fig. 4.1). Maximum increases occurred in northern high and mid-latitudes, where the growth rates were over 15 ppb/year. In 2000 and 2001, the global growth rate decreased to around -1 ppb/year. Around 2002/2003, the growth rates increased in the Northern Hemisphere, especially in northern high and mid-latitudes where they reached about 10 ppb/year. The global growth rate was -3 ppb/year in 2004 and 1 ppb/year in 2005. Despite the large growth rates in 1998 and 2002/2003, during El Niño events, the global mean mole fraction was relatively stable between 1999 and 2006. However, the global mean mole fraction increased by a total of 35 ppb in the five years since 2007.

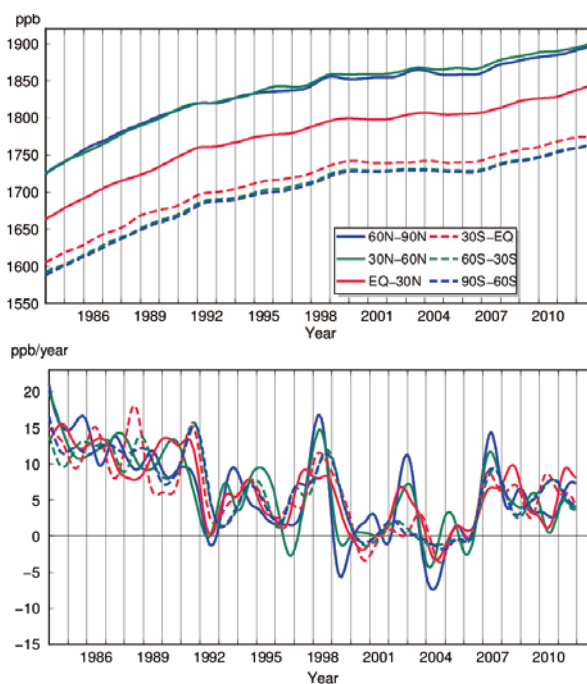


Fig. 4.3 Long-term trends in the mole fraction of CH_4 for each 30° latitudinal zone (top) and their growth rates (bottom).

The large increase in CH_4 growth rate in 1991 may have been caused by decreased levels of OH radicals in the atmosphere due to reduced UV radiation resulting from the eruption of Mount Pinatubo in 1991 (Dlugokencky *et al.*, 1996), and the subsequent decrease in 1992 may have been due to an increase in OH radicals resulting from the depletion of stratospheric ozone following this eruption (Bekki *et al.*, 1994).

In 1998, the growth rates were high in all latitudes, which may have been due to increased emissions in northern high latitudes and tropical wetlands caused by high temperatures and increased precipitation, as well as by biomass burning in boreal forests, mainly in Siberia (Dlugokencky *et al.*, 2001). In contrast,

Morimoto *et al.* (2006) estimated from isotope observations that the contribution of biomass burning to the increase in 1998 was about half that of wetlands. The growth rates were low from 1999 to 2006, with an exception during the El Niño event of 2002/2003. The causes of these decreases and increases in CH₄ growth rates are not yet determined.

Since 2007, atmospheric CH₄ has increased significantly throughout the entire monitoring network (Rigby *et al.*, 2008; Dlugokencky *et al.*, 2009). This is due to increased emissions in the tropical and mid-latitude Northern Hemisphere. The attribution of this increase to anthropogenic and natural sources is difficult because the current network is insufficient to characterize emissions by region and source process (Bergamaschi *et al.*, 2013).

The WMO/GAW observational network includes the observations of carbon stable isotopes in methane, with 20 datasets submitted to the WDCGG. Such observations can be useful for the identification of primary methane sources.

Seasonal cycle of CH₄ mole fraction in the atmosphere

Figure 4.4 shows seasonal cycles in the mole fraction of CH₄ for each 30° latitudinal zone. The seasonal cycles are driven mainly by reaction with OH radicals, a major CH₄ sink in the atmosphere. Seasonal cycles are also affected by the magnitude and timing of CH₄ emissions from sources such as wetlands and biomass burning as well as by its atmospheric transport. The seasonal cycles are large in amplitude in the Northern Hemisphere. Unlike CO₂, amplitudes were also large in southern high and mid-latitudes. Seasonally, the Northern Hemisphere shows minima in summer and maxima in winter, while the Southern Hemisphere shows a seasonal cycle lagging two-thirds to three-quarter years behind. The seasonal variations in the mole fraction of CH₄ were almost consistent with those of the OH radical that reacts with CH₄. Southern low latitudes have a distinct antiphase annual component with that of the seasonal cycle arising from southern mid-latitudes. The maximum in the former component occurs in boreal winter due to the interhemisphere transport of CH₄ from the Northern Hemisphere.

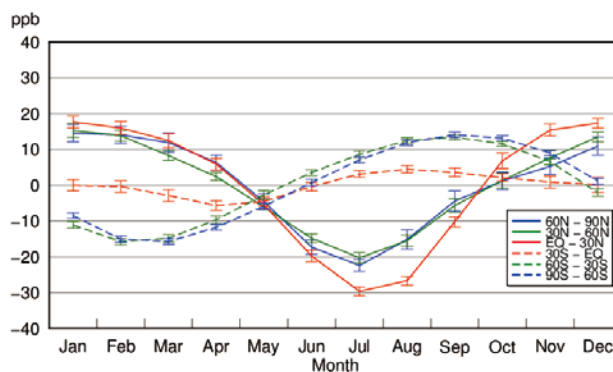


Fig. 4.4 Average seasonal cycles of CH₄ mole fraction for each 30° latitudinal zone obtained by subtracting long-term trends from the zonal mean time series. Vertical error bars represent the range of $\pm 1\sigma$ calculated for each month. (Average from 1984 to 2012)

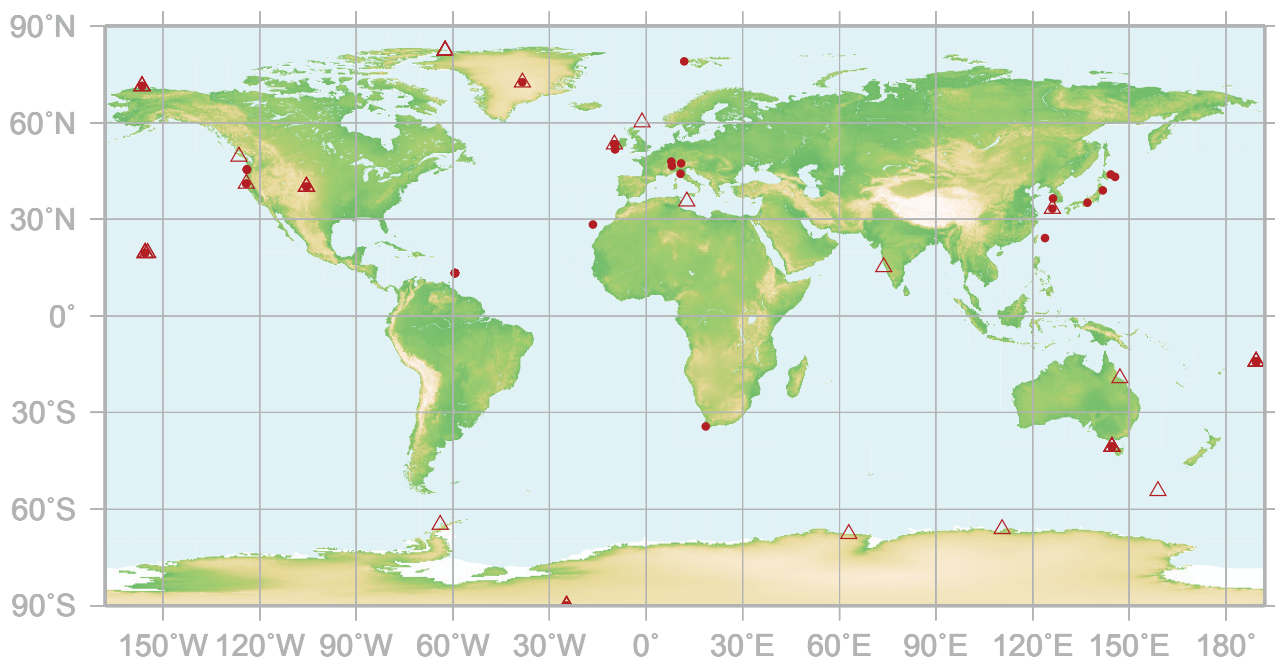
5.

NITROUS OXIDE

(N₂O)

● : CONTINUOUS STATION

△ : FLASK STATION



This map shows locations of the stations that have submitted data for monthly mean mole fraction.

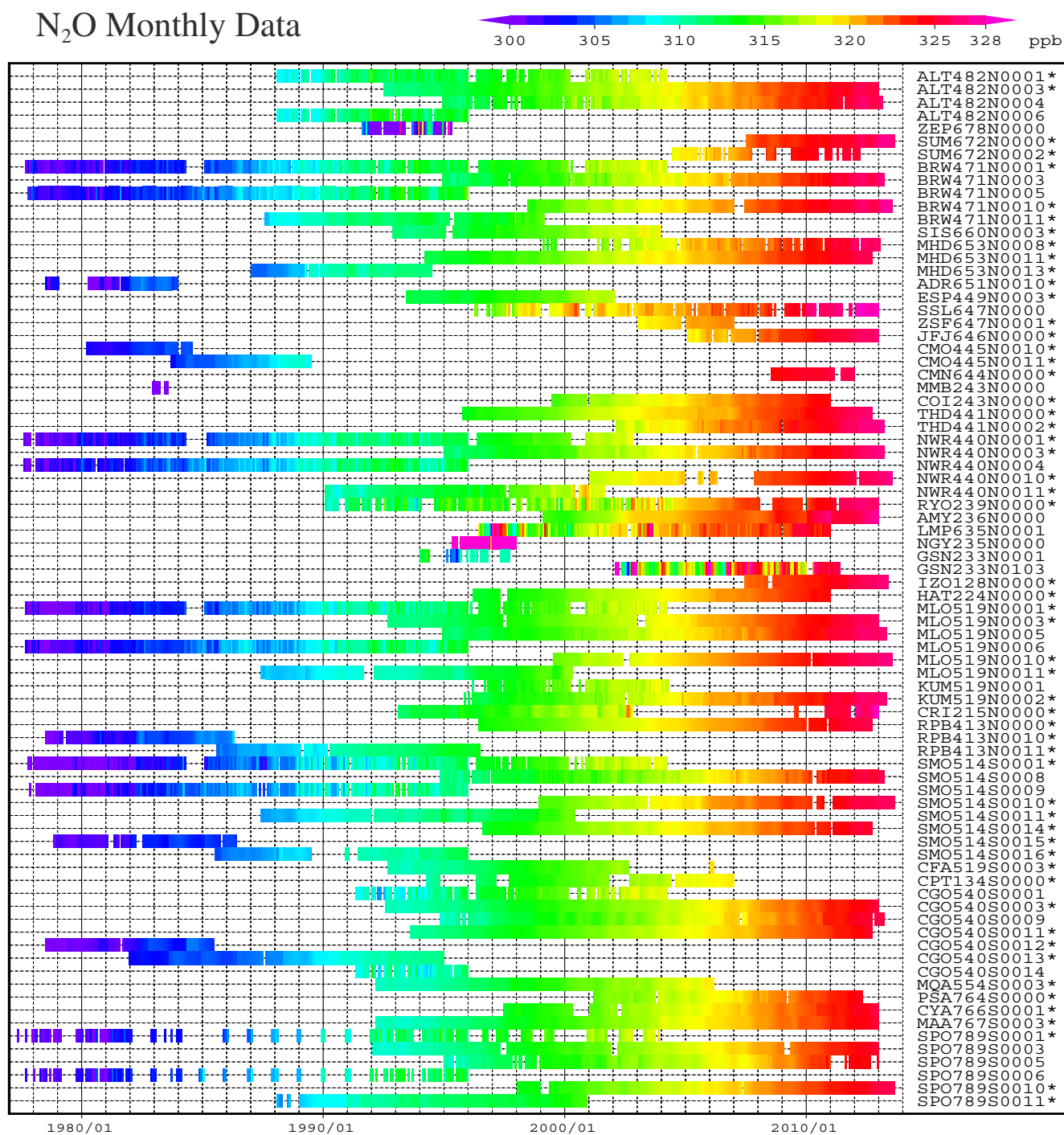
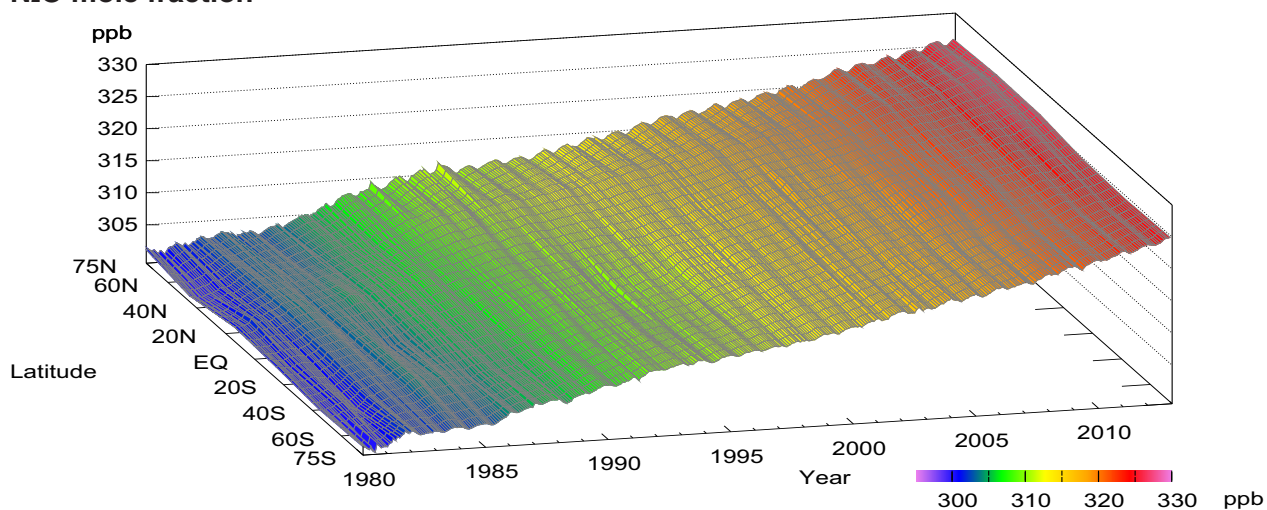
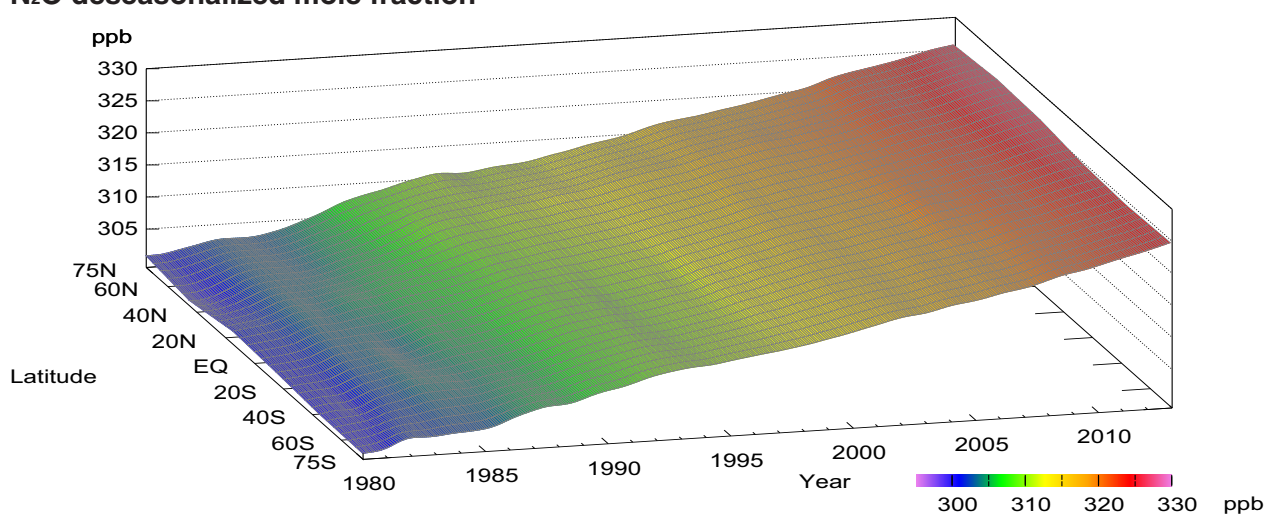


Plate 5.1 Monthly mean N₂O mole fractions that have been reported to the WDCGG. The mole fractions are illustrated in different colours. The sites are listed in order from north to south. The data from the sites with an asterisk at the end of the station index are used for the analysis shown in Plate 5.2. (see Chapter 2)

N₂O mole fraction



N₂O deseasonalized mole fraction



N₂O growth rate

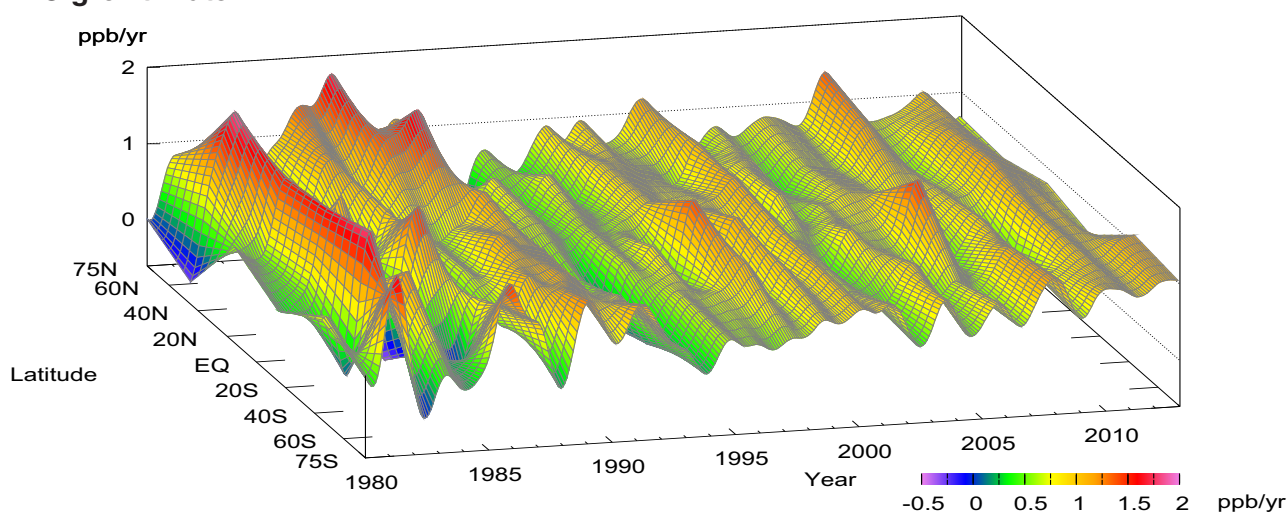


Plate 5.2 Variation of zonally averaged monthly mean N₂O mole fractions (top), deseasonalized long-term trends (middle), and growth rates (bottom). The zonally averaged mole fractions are calculated for each 30° zone. The deseasonalized trends and growth rates are derived as described in Chapter 2.

5. NITROUS OXIDE (N₂O)

Basic information on N₂O with regard to environmental issues

Nitrous oxide (N₂O) is a relatively stable greenhouse gas in the troposphere with a lifetime of 121 years (IPCC, 2013). Between 1750 and 2012, N₂O accounted for about 6 % of the total increase in radiative forcing due to long-lived greenhouse gases (WMO, 2013). N₂O is the third most important anthropogenic greenhouse gas in the atmosphere. The mole fraction of N₂O in the atmosphere has increased steadily from about 270 ppb during pre-industrial time to its current value, which is 20% higher. N₂O is emitted into the atmosphere from natural and anthropogenic sources, including the ocean, soil, combustion of fuels, biomass burning, use of fertilizer and various industrial processes. The amount of N₂O emitted into the atmosphere by human activities is approximately equal to that derived from natural systems (oceans, chemical oxidation of ammonia in the atmosphere, and soils). Most of the anthropogenic N₂O comes to the atmosphere from the transformation of fertilizer nitrogen into N₂O and its subsequent emission from agricultural soils. N₂O breaks down mainly by photo-dissociation in the stratosphere, forming nitrogen oxides that trigger ozone depleting reactions, so it may play important role in ozone depletion in the future when other ozone depleting substances are depleted or removed from the atmosphere (Ravishankara *et al.*, 2009). However, the complete cycling of N₂O in the global nitrogen cycle remains unclear.

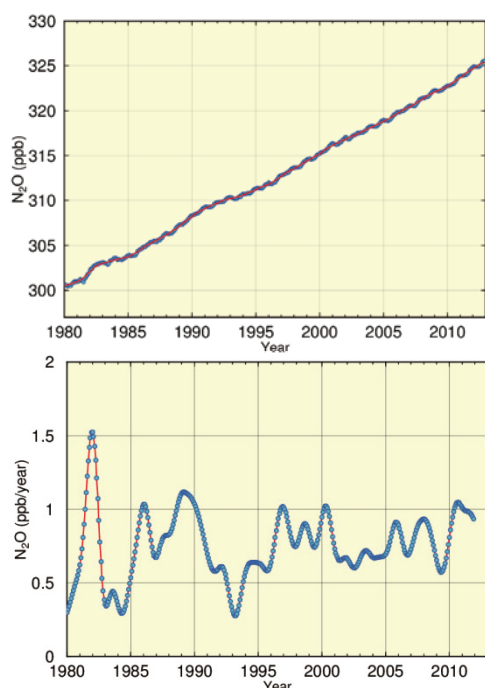


Fig. 5.1 Globally averaged monthly mean mole fraction of N₂O from 1980 to 2012, including deseasonalized long-term trend shown as a red line (top) and annual growth rate (bottom).

Long-term trend of N₂O mole fraction in the atmosphere

Dry mole fractions of N₂O are analyzed using the data submitted to the WDCGG from fixed stations and some ships. The observational sites that supplied data used for this analysis are shown on the map at the beginning of this chapter. The monthly mean mole fractions of N₂O, including the ones used in the global analysis, are shown in Plate 5.1, with the various mole fraction levels illustrated in different colours. The data submitted to the WDCGG show that N₂O mole fractions have increased at almost all stations. Zonally averaged atmospheric N₂O mole fractions, together with their deseasonalized components and growth rates, are shown as three-dimensional representations from 1980 to 2012 in Plate 5.2. Figure 5.1 shows globally averaged monthly mean N₂O mole fraction from 1980 to 2012 and its long-term trend. The global average mole fraction reached a new high of 325.1 ± 0.1 ppb in 2012, an increase of 0.9 ppb over the previous year. The mean annual absolute increase during last 10 years was 0.80 ppb/year. Atmospheric growth rate showed substantial variability (from 0.6 to 1.0 ppb/year) from the beginning of observations. The interhemispheric gradient in the yearly averaged mole fraction of N₂O changed from 0.2 ppb in 1980 to 1.4 ppb in 2012 (Figure 5.2 upper panel), indicating that the majority of N₂O sources are situated in the Northern Hemisphere (mostly agriculture).

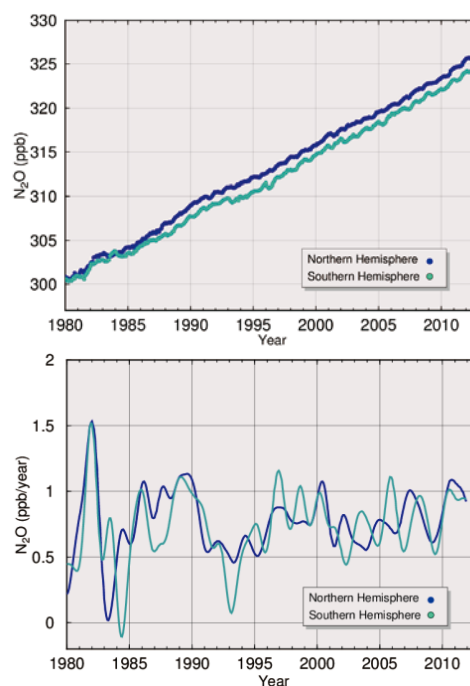


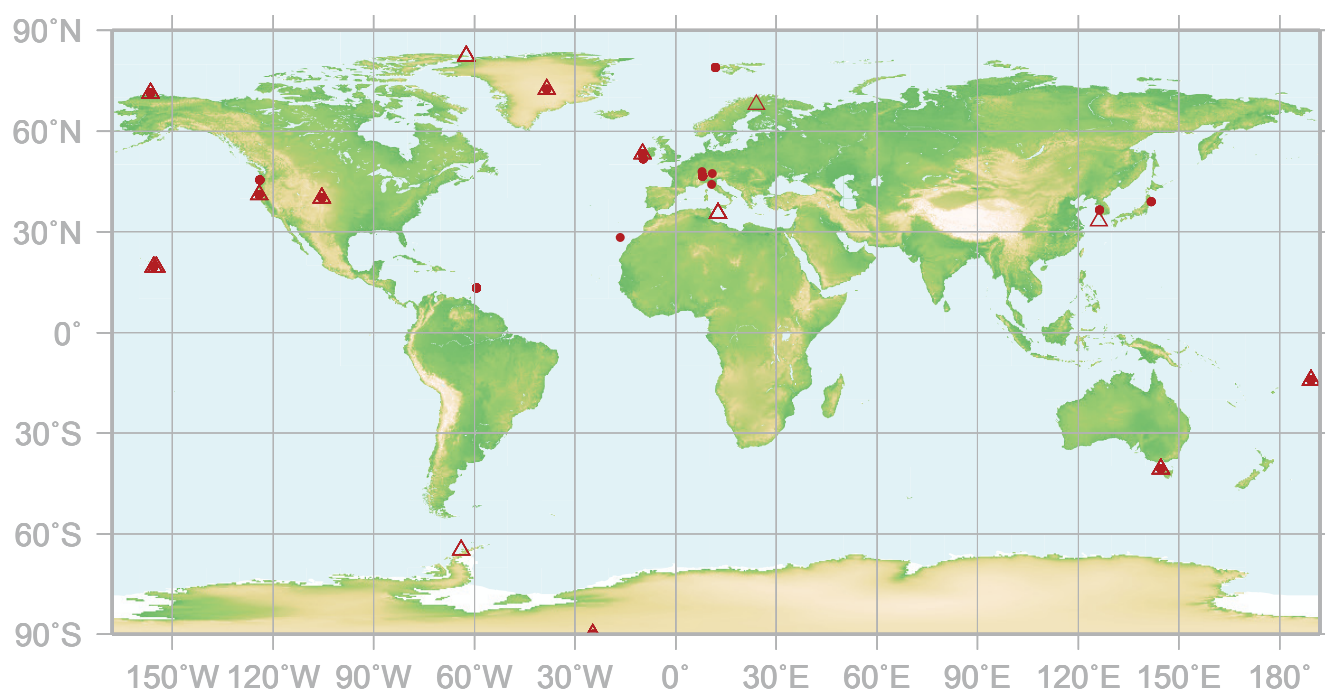
Fig. 5.2 Monthly mean mole fraction of N₂O from 1980 to 2012 (top) and annual growth rates (bottom), averaged over the Northern and Southern Hemispheres.

6.

HALOCARBONS AND OTHER HALOGENATED SPECIES

● : CONTINUOUS STATION

△ : FLASK STATION



This map shows locations of the stations that have submitted data for monthly mean mole fraction.

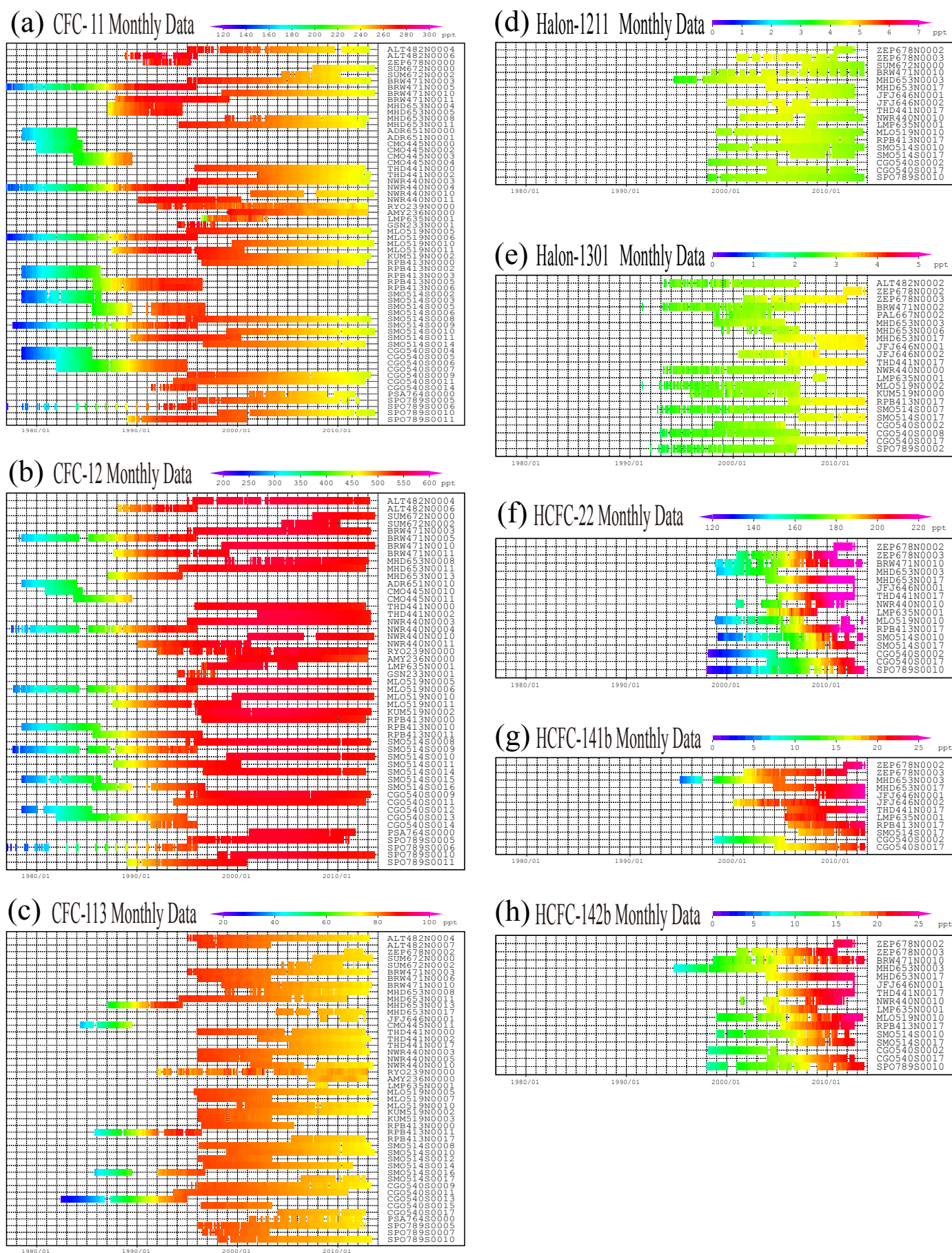


Plate 6.1 Monthly mean (a) CFC-11, (b) CFC-12, (c) CFC-113, (d) Halon-1211, (e) Halon-1301, (f) HCFC-22, (g) HCFC-141b, (h) HCFC-142b mole fractions that have been reported to the WDCGG. The mole fractions are illustrated in different colours. The sites are listed in order from north to south.

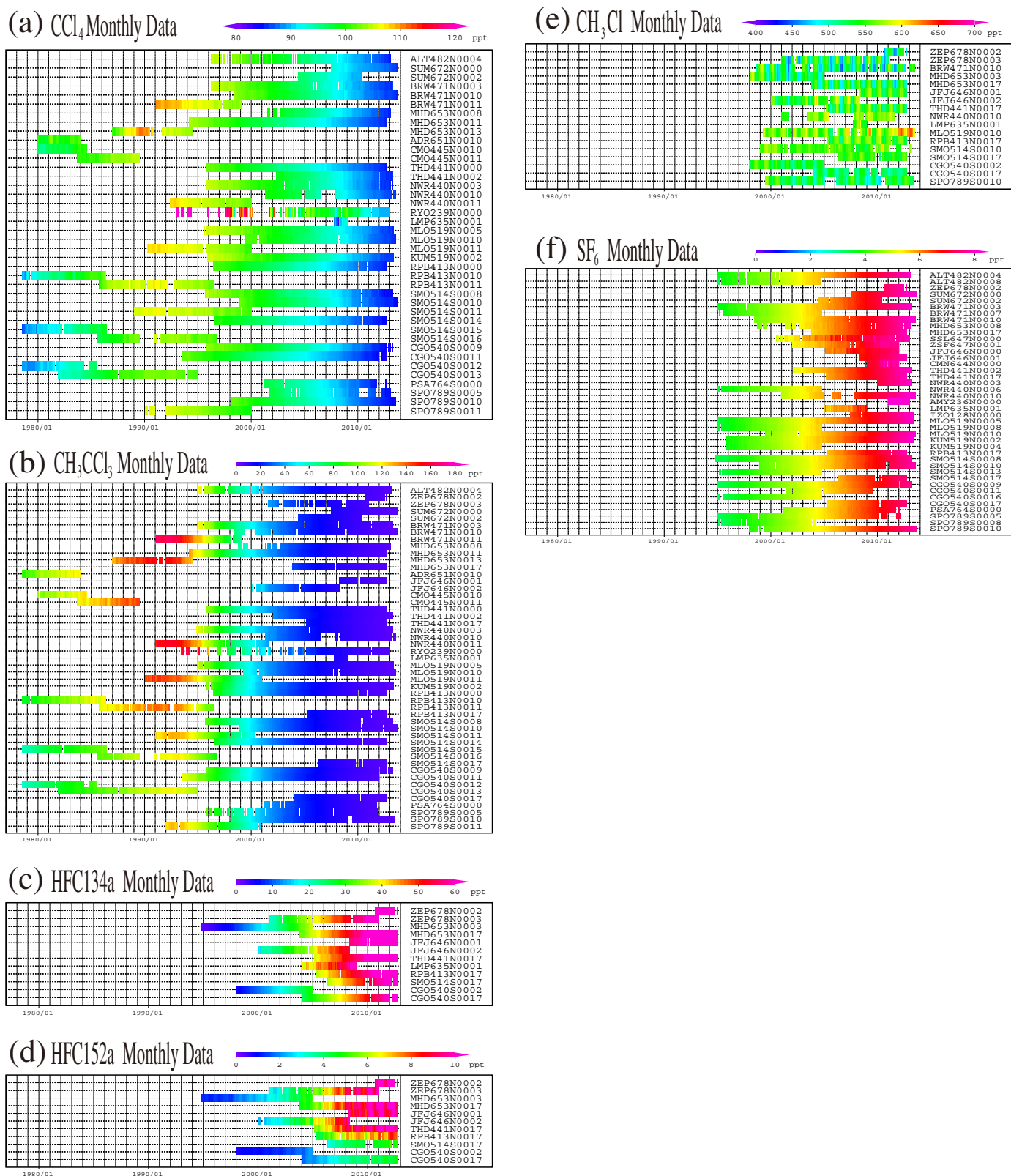


Plate 6.2 Monthly mean (a) CCl_4 , (b) CH_3CCl_3 , (c) HFC134a , (d) HFC152a , (e) CH_3Cl , (f) SF_6 mole fractions that have been reported to the WDCGG. The mole fractions are illustrated in different colours. The sites are listed in order from north to south.

6. HALOCARBONS AND OTHER HALOGENATED SPECIES

Basic information on halocarbons with regard to environmental issues

Halocarbons are carbon compounds containing one or more halogens, *i.e.*, fluorine, chlorine, bromine or iodine, with most being industrial products. Halocarbons are classified into chlorofluorocarbons (CFCs), which contain fluorine and chlorine; the hydrochlorofluorocarbons (HCFCs), which contain hydrogen in addition to fluorine and chlorine; and the halons, which contain bromine and other halogens. Perfluorocarbons (PFCs) are carbon compounds in which all hydrogen atoms are replaced by fluorine atoms, and hydrofluorocarbons (HFCs) are halocarbons that contain hydrogen and fluorine but no chlorine. Sulphur hexafluoride (SF_6), although not a halocarbon, behaves similarly to halocarbons and is a potent long-lived greenhouse gas. Carbon tetrachloride (CCl_4) and methyl chloroform (CH_3CCl_3) are produced industrially, whereas methyl chloride (CH_3Cl) has natural sources. Although the mole fractions of the halocarbons are relatively low in the atmosphere, they have high global warming potentials. The halocarbons have been shown to account for about 12% of the total increase in radiative forcing due to long-lived greenhouse gases from 1750 to 2012 (WMO, 2013).

The halocarbons are colourless, odourless and innocuous substances that can be readily gasified and liquefied and have low surface tension. Thus, they were commonly used as refrigerants, propellants and detergents for semiconductors, resulting in a rapid increase in their mole fractions in the atmosphere until the mid-1980s. Halocarbons containing chlorine and bromine led to the depletion of the ozone layer. Since the mid-1990s, the Montreal Protocol on Substances that Deplete the Ozone Layer and its subsequent Adjustments and Amendments have progressively tightened the regulations for the production, consumption and trade of ozone-depleting substances.

The CFCs are destroyed mainly by ultraviolet radiation in the stratosphere, and their lifetimes are generally long (*e.g.*, about 50 years for CFC-11). However, the HCFCs and CH_3CCl_3 , which contain hydrogen, react with hydroxyl radicals (OH) in the troposphere and have relatively short lifetimes (*e.g.*, about 5 years for CH_3CCl_3). As the reaction with OH in the troposphere is a major sink for CH_3CCl_3 , global measurements of CH_3CCl_3 provide an accurate estimate of the global mole fraction of OH (Prinn *et al.*, 2001). However, due to a substantial decrease of CH_3CCl_3 in the atmosphere, reconstruction of OH levels using this molecule is becoming increasingly difficult and other compounds are now used as reference tracers for OH mole fraction determination.

The Kyoto Protocol to the United Nations

Framework Convention on Climate Change (UNFCCC), which came into force on 16 February 2005, specifies HFCs, PFCs and SF_6 as targets for quantified emission limitation and reduction commitments.

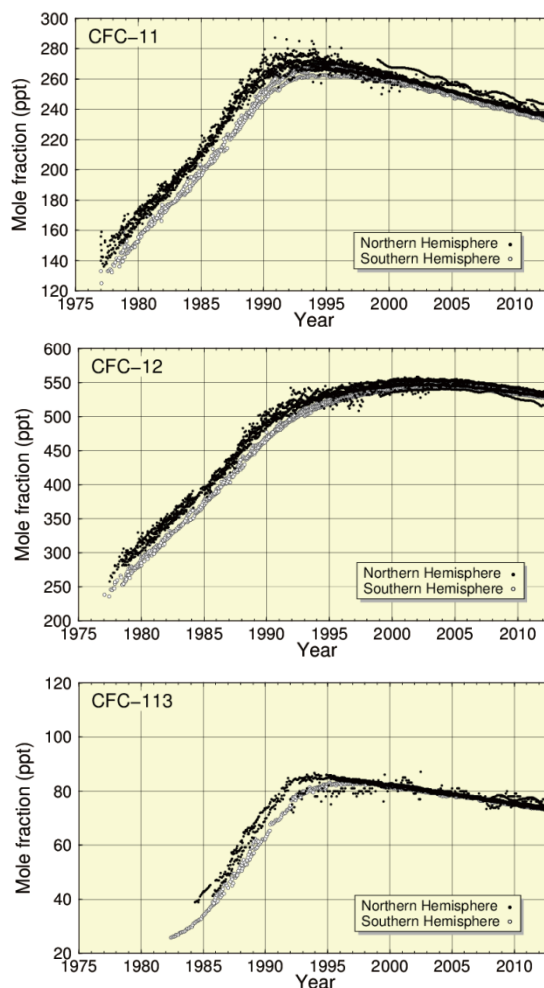


Fig. 6.1 Time series of the monthly mean mole fractions of CFC-11, CFC-12 and CFC-113 at individual stations. Solid circles show mole fractions measured in the Northern Hemisphere and open circles show mole fractions measured in the Southern Hemisphere.

Annual changes in the levels of halocarbons in the atmosphere

The cover map of this chapter shows observational sites that have submitted data on halocarbons and other halogenated species to the WDCGG. Although the number of stations measuring these species is rather limited, halocarbons are generally well mixed in the atmosphere and the data may be sufficient to reflect their global tendencies. Plates 6.1 and 6.2 show all the monthly mean mole fractions of these gases

submitted to the WDCGG. The figures (6.1 – 6.7) in this chapter show the monthly mean data reported to the WDCGG without spatial averaging. Some discrepancies in the absolute mole fractions were observed for several stations, suggesting that these stations may have adopted different standard scales. Observational data based on identical standard scales revealed that the differences in the mole fractions between the two hemispheres were large in the 1980s for CFCs, CCl_4 and CH_3CCl_3 but have since narrowed as the emissions have been suppressed and the existing constituents have been mixed across the hemispheres.

Figure 6.1 shows monthly mean mole fractions of CFC-11 (CCl_3F), CFC-12 (CCl_2F_2) and CFC-113 ($\text{CCl}_2\text{FCClF}_2$) over time. The mole fractions of CFC-11 were maximal around 1992 in the Northern Hemisphere, followed by a maximum about one year later in the Southern Hemisphere. The mole fractions of CFC-113 were maximal around 1992 in the Northern Hemisphere and around 1997 in the Southern Hemisphere. The mole fractions of these gases have since been decreasing slowly in both hemispheres. The mole fraction of CFC-12 increased until around 2005 and then started decreasing gradually.

Figure 6.2 shows time series of the monthly mean mole fractions of Halon-1211 (CBrClF_2) and Halon-1301 (CBrF_3). The mole fraction of Halon-1211 has not increased since 2005, whereas the mole fraction of Halon-1301 is increasing.

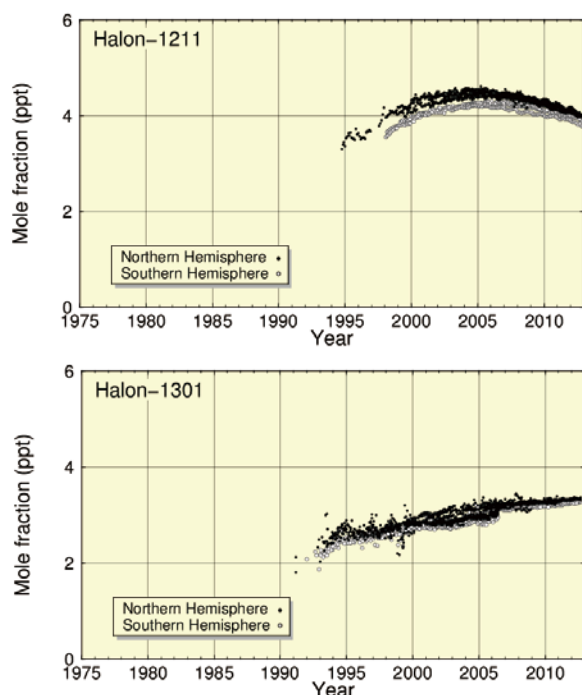


Fig. 6.2 Time series of the monthly mean mole fractions of Halon-1211 and Halon-1301 at individual stations. Solid circles show mole fractions measured in the Northern Hemisphere and open circles show mole fractions measured in the Southern Hemisphere.

Figure 6.3 shows time series of the mole fractions of HCFC-22 (CHClF_2), HCFC-141b ($\text{CH}_3\text{CCl}_2\text{F}$) and HCFC-142b (CH_3CClF_2). The mole fractions of these gases increased significantly during the last decade as a result of their continued use as substitutes for CFCs. However, the growth of HCFC-141b decelerated rapidly in the last decade.

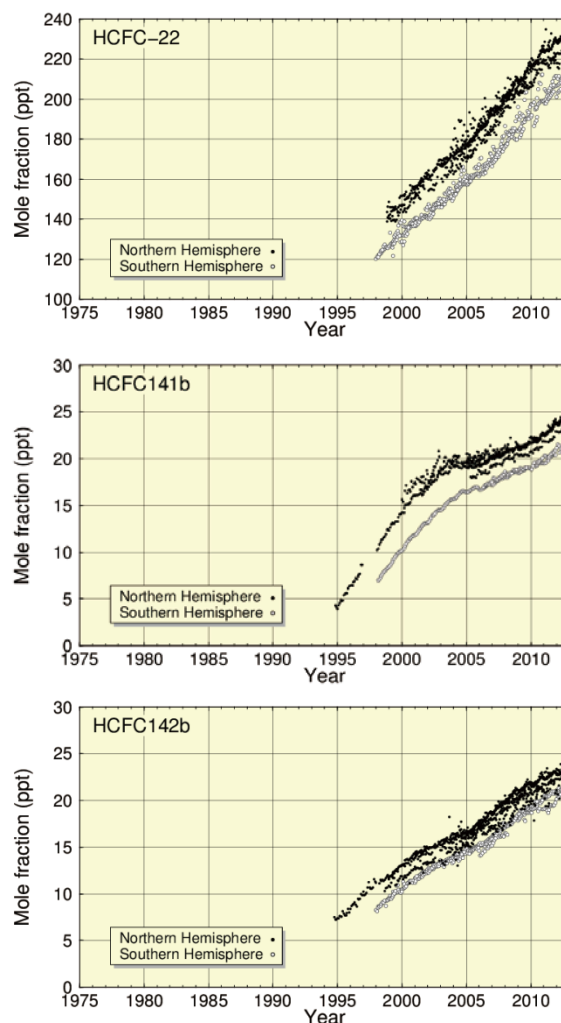


Fig. 6.3 Time series of the monthly mean mole fractions of HCFC-22, HCFC-141b, and HCFC-142b at individual stations. Solid circles show mole fractions measured in the Northern Hemisphere and open circles show mole fractions measured in the Southern Hemisphere.

Figure 6.4 shows time series of the mole fractions of CCl_4 and CH_3CCl_3 . The mole fractions of CCl_4 in both hemispheres were at a maximum around 1991. The mole fractions of CH_3CCl_3 were at a maximum around 1992 in the Northern Hemisphere and around 1993 in the Southern Hemisphere. The mole fractions of these gases have since been decreasing.

Figure 6.5 shows time series of the monthly mean mole fractions of HFC-134a (CH_2FCF_3) and HFC-152a (CH_3CHF_2). The mole fractions of HFC-134a and HFC-152a have increased by 4 to

5-fold over the last 10 years. These increases have been larger in the Northern than in the Southern Hemisphere, suggesting that predominant sources of this compound are located in the Northern Hemisphere.

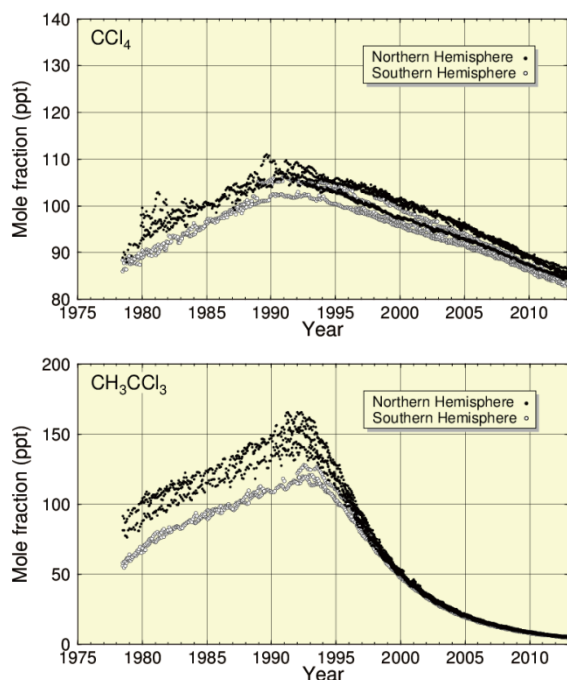


Fig. 6.4 Time series of the monthly mean mole fractions of CCl_4 and CH_3CCl_3 at individual stations. Solid circles show mole fractions measured in the Northern Hemisphere and open circles show mole fractions measured in the Southern Hemisphere.

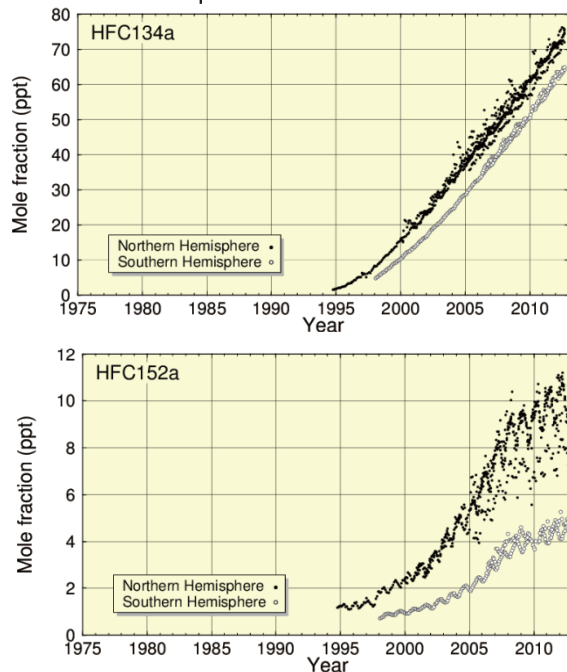


Fig. 6.5 Time series of the monthly mean mole fractions of HFC-134a and HFC-152a at individual stations. Solid circles show mole fractions measured in the Northern Hemisphere and open circles show mole fractions measured in the Southern Hemisphere.

Figure 6.6 shows time series of the monthly mean mole fractions of methyl chloride (CH_3Cl). The mole fraction of CH_3Cl does not show any particular long-term tendency although clear seasonal cycle can be seen in the dataset.

Figure 6.7 shows a time series of the monthly mean mole fractions of SF_6 . The mole fraction of SF_6 in 2012 was about twice of the level observed in the mid-1990s increasing nearly linearly at a rate of 0.24 ppt/year (WMO, 2013).

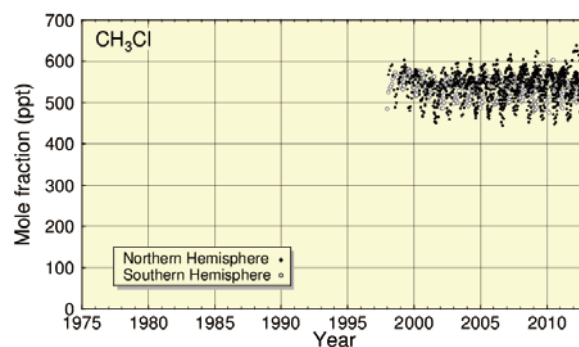


Fig. 6.6 Time series of the monthly mean mole fractions of CH_3Cl at individual stations. Solid circles show mole fractions measured in the Northern Hemisphere and open circles show mole fractions measured in the Southern Hemisphere.

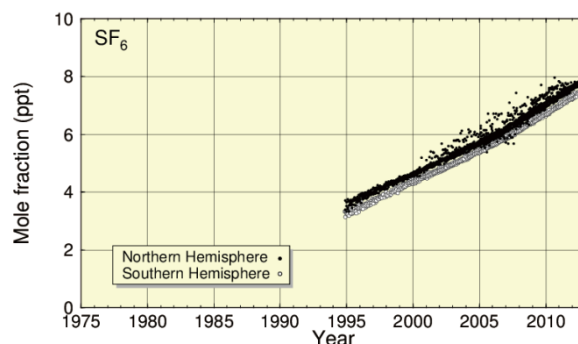
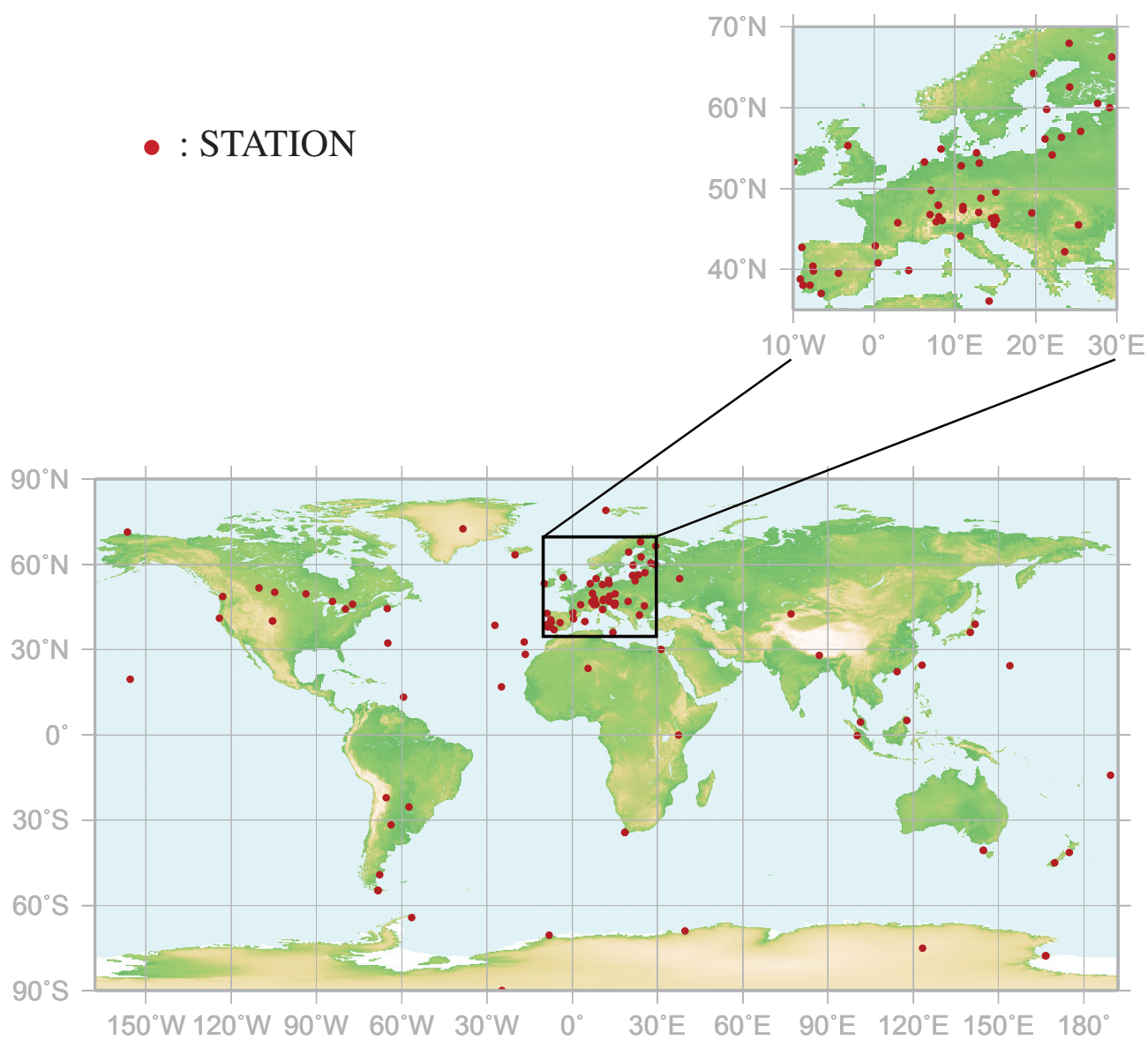


Fig. 6.7 Time series of the monthly mean mole fractions of SF_6 at individual stations. Solid circles show mole fractions measured in the Northern Hemisphere and open circles show mole fractions measured in the Southern Hemisphere.

7.

SURFACE OZONE

(O₃)



This map shows locations of the stations that have submitted data for monthly mean mole fraction.

O₃ Monthly Data

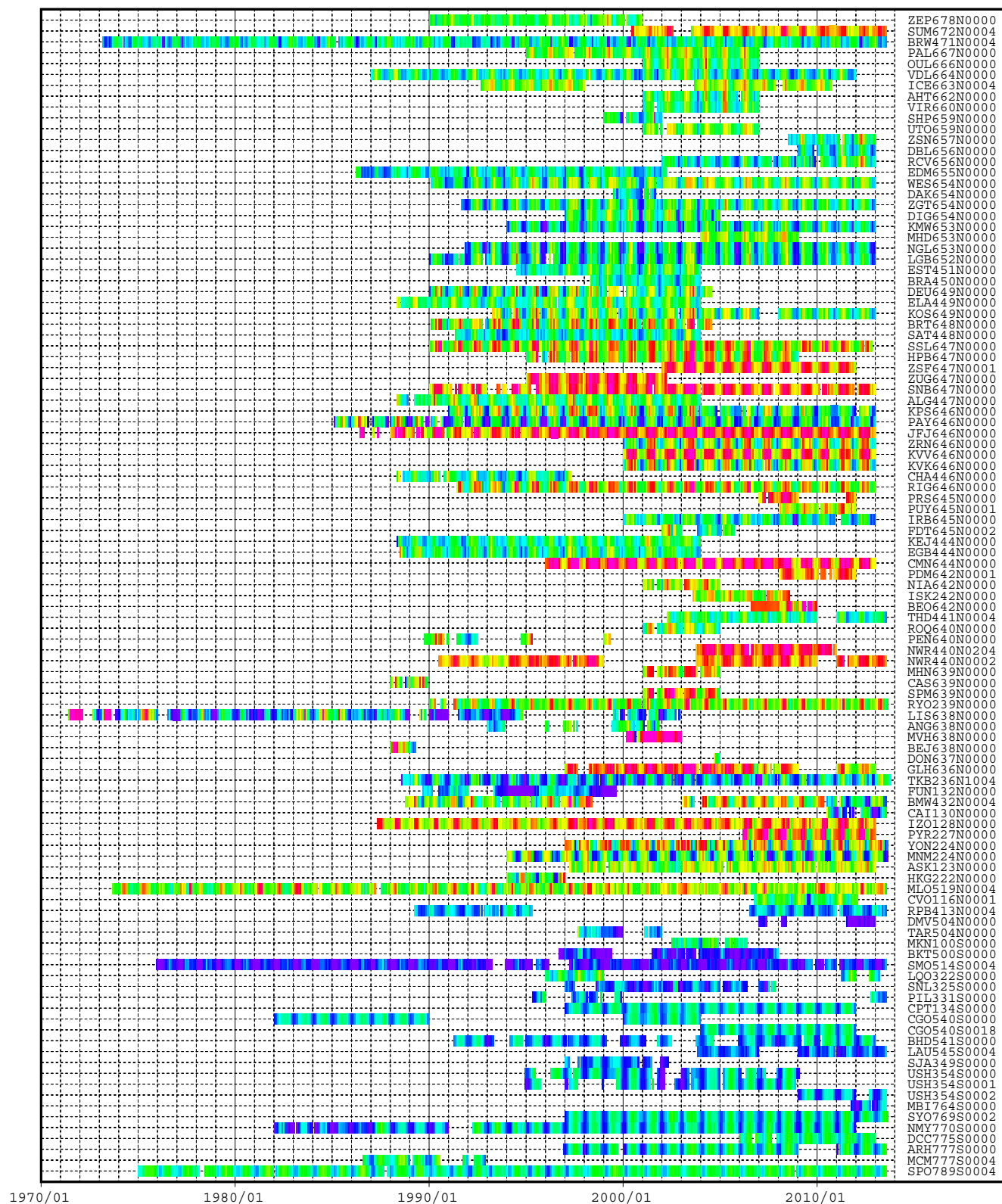
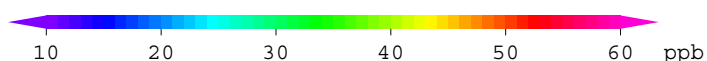


Plate 7.1 Monthly mean O₃ mole fractions that have been reported to the WDCGG. The mole fractions are illustrated in different colours. The sites are listed in order from north to south.

7. SURFACE OZONE (O₃)

Basic information on surface ozone (O₃) with regard to environmental issues

Ozone (O₃) in the atmosphere exists mostly in the stratosphere, with less than 10% in the troposphere. However, O₃ in the troposphere plays an important role in the atmospheric environment through its impact on radiative forcing and its involvement in the chemical processes. O₃ absorbs UV radiation in the stratosphere, thus influencing the vertical profile of temperature and circulation in the stratosphere. Moreover, as a greenhouse gas in the troposphere, O₃ absorbs IR radiation. The latter effect is more significant in the upper troposphere. Tropospheric O₃ in the northern extra tropics was the greatest contributor to global warming during the 20th century, and increases in tropospheric O₃ related to industrialization in developing countries was found to contribute to accelerated warming in the tropics during the latter half of the century (Shindell *et al.*, 2006). Furthermore, by reacting with water vapor in the presence of UV radiation, O₃ produces OH radicals, which control atmospheric mole fractions of many greenhouse gases, such as CH₄, through chemical reactions.

The observational results at high altitudes around 1990, compared with those from the end of the 19th century to the first half of the 20th century, show increases in tropospheric O₃, especially in urban areas (Staehelin *et al.*, 1994). However, ozonesonde measurements in the troposphere show stable or decreasing trends in northern mid-latitudes after 1980 (Oltmans *et al.*, 2006). There is as yet no consensus on the global trend of tropospheric O₃. Recently however, an attempt has been made to systematically review the observed trends. It was found that in most regions of the world — excluding East Asia — surface and free tropospheric ozone concentrations have not risen significantly since year 2000. Prior to the 1990s almost all records indicate a strong rise, while during the 1990s the picture is very diverse (http://www.igacproject.org/sites/all/themes/bluemasters/images/NewsletterArchives/Issue_45_Oct_2011.pdf).

Tropospheric O₃ originates from flux/mixing from the stratosphere and in-situ photochemical production. O₃ is destroyed in various processes, including chemical reactions with NO, the hydroperoxyl radical (HO₂) and OH, and deposition at different surfaces. The lifetime of tropospheric ozone varies from one or a few days in the boundary layer to a few tens of days or even a few months in the free troposphere.

In the troposphere, the mole fractions of O₃ are high in high and mid-latitudes in both hemispheres, and low in the tropics over the Atlantic (Marenco and Said, 1989) and Pacific (Tsutsumi *et al.*, 2003) oceans. The localised sources of ozone precursors and the generally short lifetime of surface O₃ make its distribution spatially non-uniform and time-variant.

Annual variation of surface O₃ mole fraction

The observational sites that have submitted data for surface O₃ to the WDCGG are shown on the map at the beginning of this chapter. The monthly mean wet mole fractions of O₃ that have been reported from these observational sites are shown in Plate 7.1, with different mole fraction levels illustrated in different colours. Data for the mole fractions of surface O₃ are reported in two different units, *i.e.*, mixing ratio (ppb) and concentration (µg/m³) at 25°C. The latter is converted to the former using the formula:

$$X_p [\text{ppb}] = (R \times T / M / P_0) \times 10 \times X_g [\mu\text{g}/\text{m}^3]$$

where R is the molar gas constant (8.31451 [J/K/mol]),

T is the absolute temperature reported from each station,

M is the molecular weight of O₃ (47.9982), and

P₀ is the standard pressure (1013.25 [hPa]).

If temperature is not reported by a station, T is taken to be 298K (25 °C).

The mole fraction of surface O₃ was found to vary from station to station, though many of these stations are located in Europe. Moreover, the seasonal and interannual variations were found to be relatively large at most stations, making it difficult to identify a global long-term trend in the mole fraction of surface O₃.

The seasonal cycles of monthly mean mole fraction of surface O₃ averaged for each 30° latitudinal zone for all available periods are shown in Figure 7.1. The latitudinal mean mole fractions were found to be elevated in spring in most latitudinal zones. However, several patterns of seasonal-diurnal cycles were observed at different locations, including a pronounced spring maximum, a spring maximum at night and a summer maximum during the day, a wide spring-summer maximum, and a pronounced winter maximum (Tarasova *et al.*, 2007).

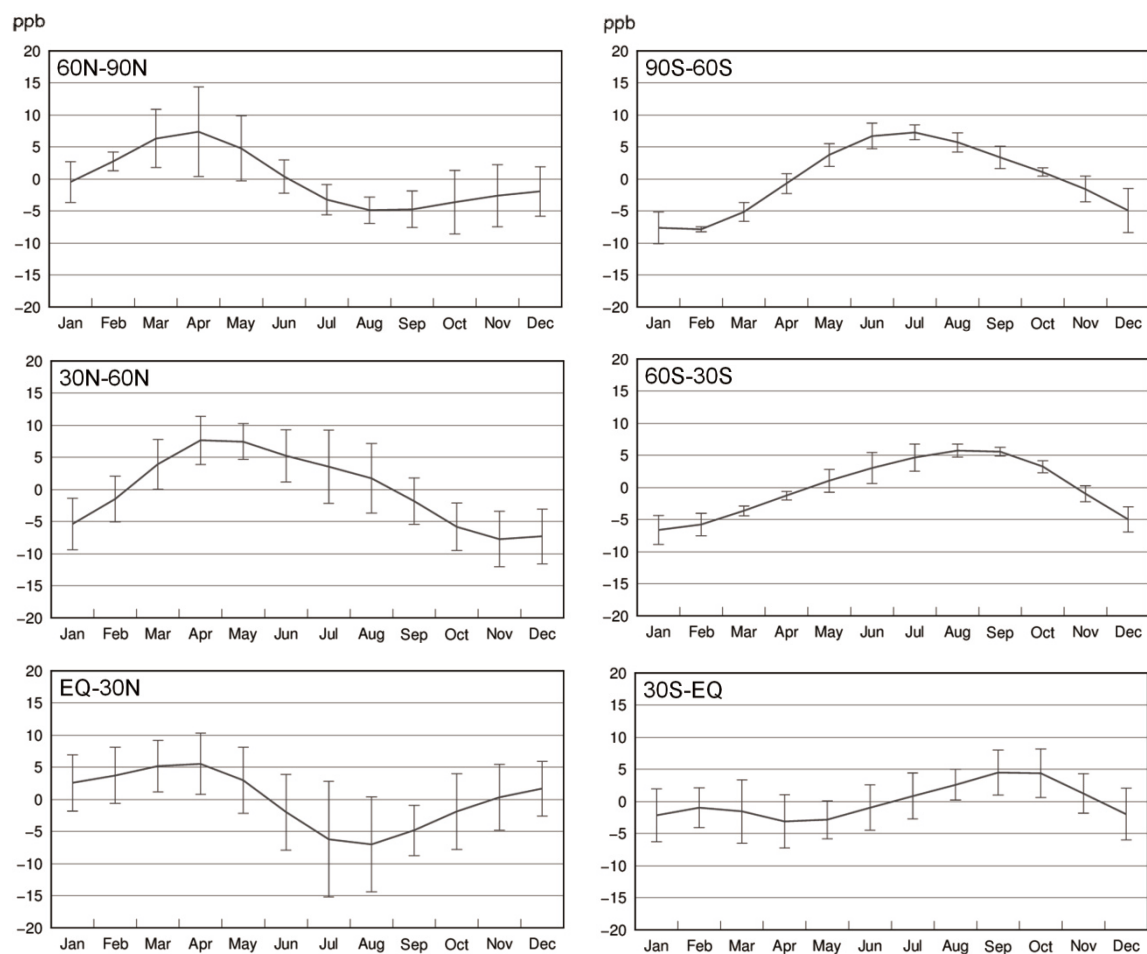
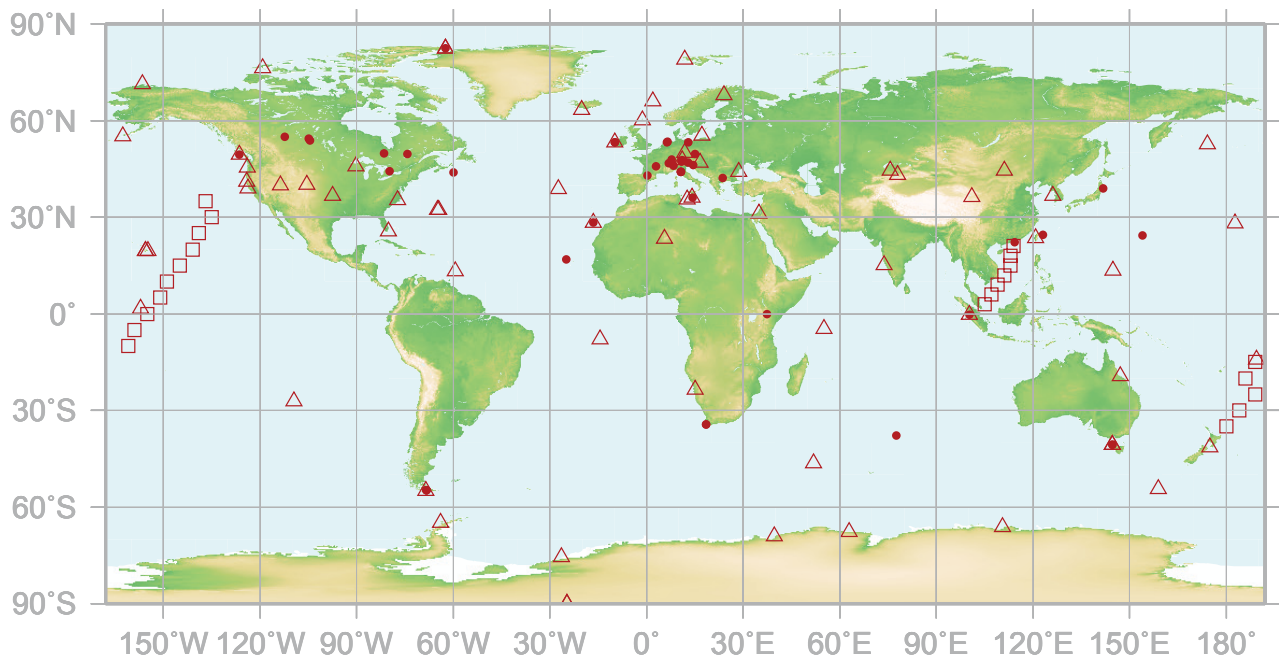


Fig. 7.1 Average seasonal cycles in the mole fraction of O_3 for each 30° latitudinal zone obtained from the seasonal cycle of each station for all available periods. Vertical error bars represent the range of $\pm 1\sigma$ calculated for each month.

8. CARBON MONOXIDE (CO)

- : CONTINUOUS STATION
- △ : FLASK STATION
- : FLASK MOBILE (SHIP)



This map shows locations of the stations that have submitted data for monthly mean mole fraction.

CO Monthly Data

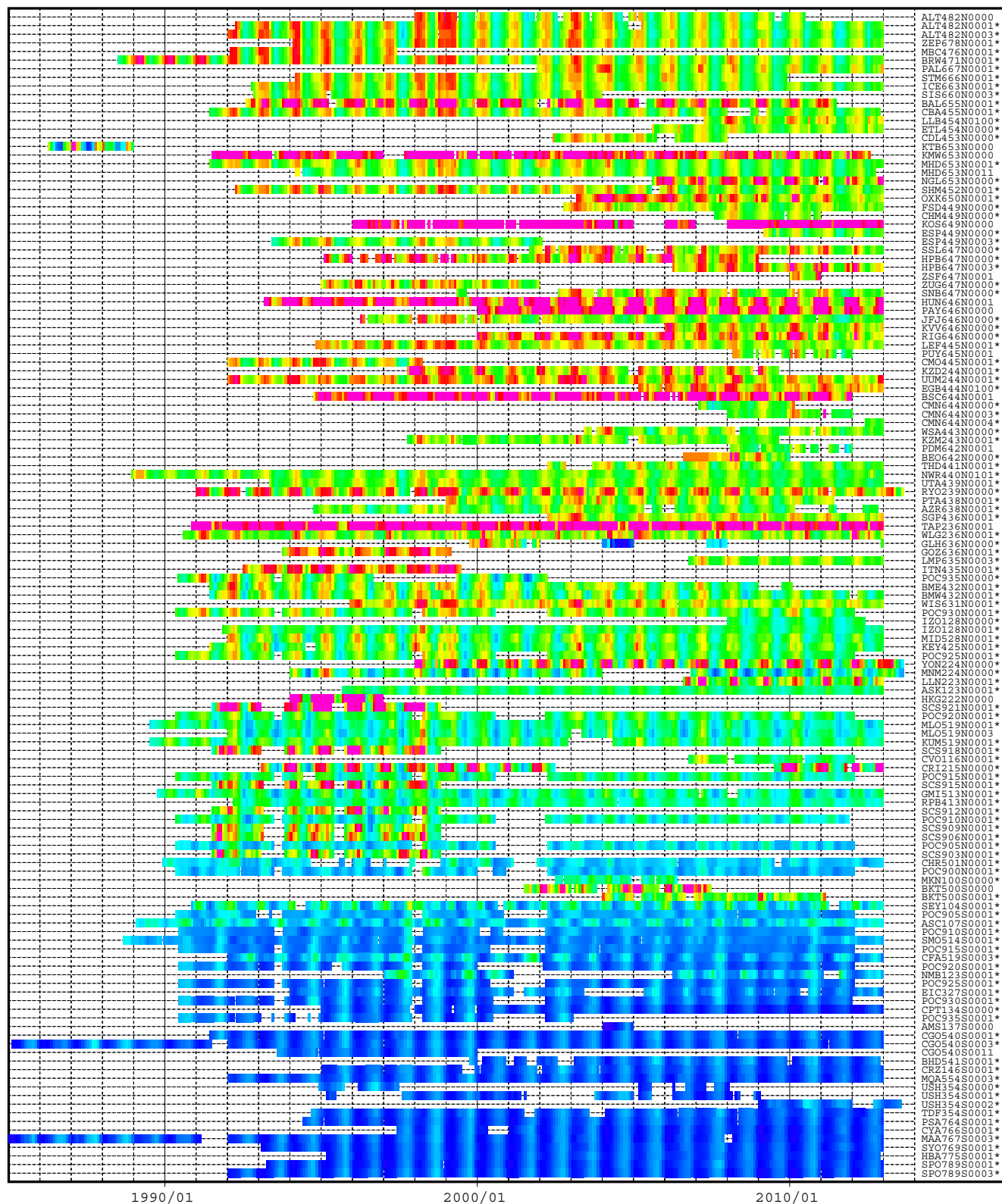
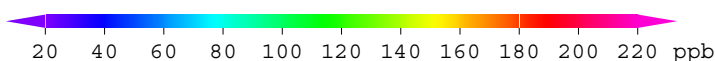
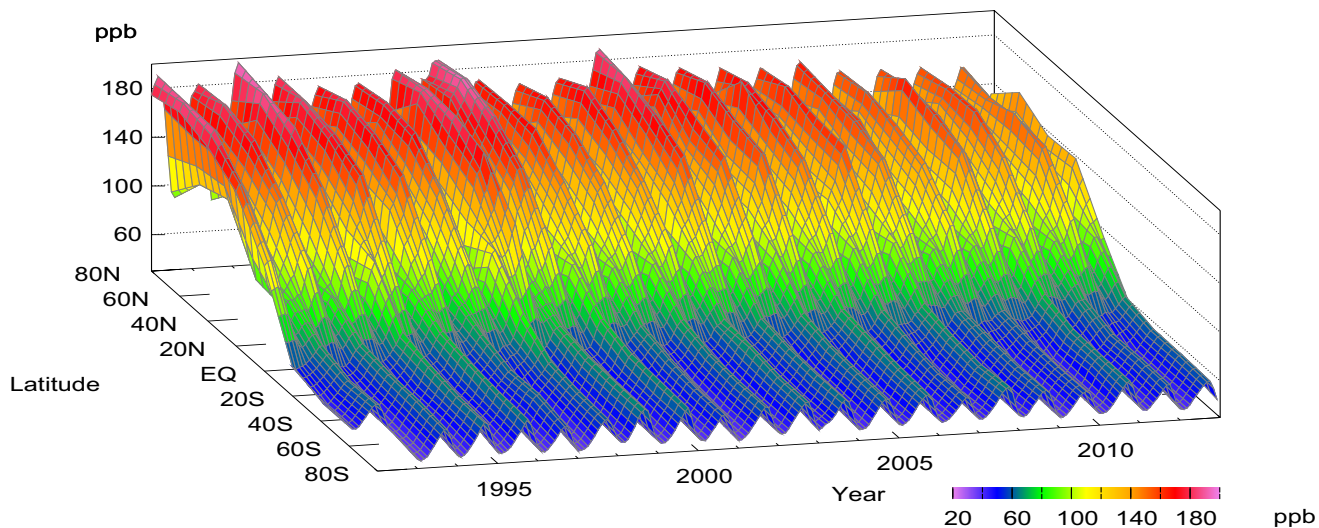
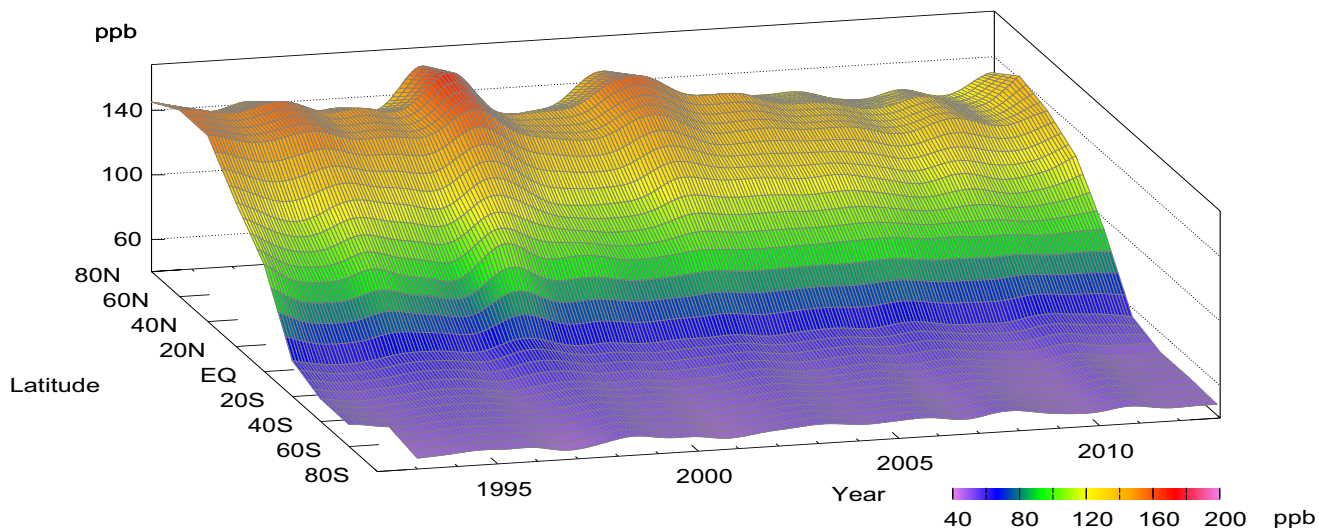


Plate 8.1 Monthly mean CO mole fractions that have been reported to the WDCGG. The mole fractions are illustrated in different colours. The sites are listed in order from north to south. The data from the sites with an asterisk at the end of the station index are used for the analysis shown in Plate 8.2. (see Chapter 2)

CO mole fraction



CO deseasonalized mole fraction



CO growth rate

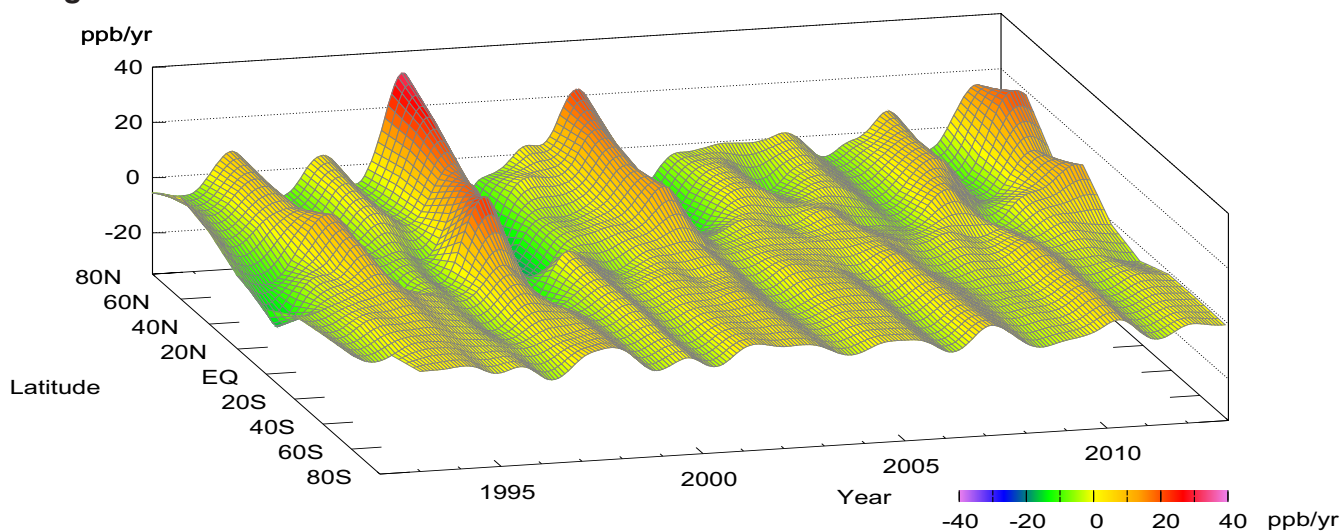


Plate 8.2 Variation of zonally averaged monthly mean CO mole fractions (top), deseasonalized long-term trends (middle), and growth rates (bottom). The zonally averaged mole fractions are calculated for each 20° zone. The deseasonalized trends and growth rates are derived as described in Chapter 2.

8. CARBON MONOXIDE (CO)

Basic information on CO with regard to environmental issues

Carbon monoxide (CO) is not a greenhouse gas; it absorbs hardly any infrared radiation from the Earth. However, CO influences the oxidation capacity of the atmosphere through its reaction with hydroxyl radicals (OH), which control the lifetimes of methane, halocarbons and tropospheric ozone. CO has been monitored due to its indirect influence on greenhouse gases through such reactions.

Sources of atmospheric CO include fossil fuel combustion and biomass burning, along with the oxidation of methane and non-methane hydrocarbons (NMHCs). Major sinks include reaction with OH and surface deposition; the reaction of CO and OH accounts for all of the chemical loss of CO in the troposphere (Seinfeld and Pandis, 1998). CO has a relatively long atmospheric lifetime, ranging from 10 days in summer in the tropics to more than a year over the polar regions in winter. Thus anthropogenic CO emissions do not lead to CO accumulation in the atmosphere. Furthermore, the uneven distribution of sources causes large spatial and temporal variations in CO mole fraction.

Measurements of trapped air in ice cores have shown that the pre-industrial CO mole fraction over central Antarctica during the last two millennia was about 50 ppb and the CO level increased to 110 ppb by 1950 in Greenland (Haan and Raynaud, 1998). Beginning in 1950, the global average CO mole fraction increased at a rate of 1% per year but started to decrease in the late 1980s (WMO, 1999). Between 1991 and 2001, the global average mole fraction of CO decreased at an annual rate of about 0.5 ppb, excluding temporal enhancements from large biomass burning events (Novelli *et al.*, 2003). Since 2001 global averaged CO mole fraction does not show any confirmed tendency.

Annual variation of CO mole fraction in the atmosphere

The monthly mean mole fractions of CO that have been reported from fixed stations and some ships to the WDCGG are shown in Plate 8.1, in which different mole fraction levels are plotted in different colours. The observational sites that provide data for global analysis are shown on the map at the beginning of this chapter.

Zonally averaged mole fractions of CO in the atmosphere, together with their deseasonalized mole fractions and growth rates, are shown in Plate 8.2 as three-dimensional representations.

Data for the mole fractions of CO are reported in various units, *i.e.*, ppb, $\mu\text{g}/\text{m}^3$ -25°C, $\mu\text{g}/\text{m}^3$ -20°C and

mg/m^3 -25°C. Units other than ppb were converted to ppb using the formulas:

$$X_p [\text{ppb}] = (R \times T / M / P_0) \times 10 \times X_g [\mu\text{g}/\text{m}^3]$$

$$X_p [\text{ppb}] = (R \times T / M / P_0) \times 10^4 \times X_g [\text{mg}/\text{m}^3]$$

where R is the molar gas constant (8.31451 [J/K/mol]),

T is the absolute temperature reported from each station,

M is the molecular weight of CO (28.0101), and

P_0 is the standard pressure (1013.25 [hPa]).

If temperature is not reported by a station, T is taken to be 298K(25 °C).

Plate 8.2 shows that the seasonal variations of CO were larger in the Northern Hemisphere and smaller in the Southern Hemisphere, and that the deseasonalized mole fractions were the highest in northern mid-latitudes and the lowest in the Southern Hemisphere, having a large latitudinal gradient from northern mid- to southern low latitudes. This is likely due to the presence of numerous anthropogenic sources of CO in the northern mid-latitudes, combined with the destruction of CO in the tropics, where OH radicals are abundant.

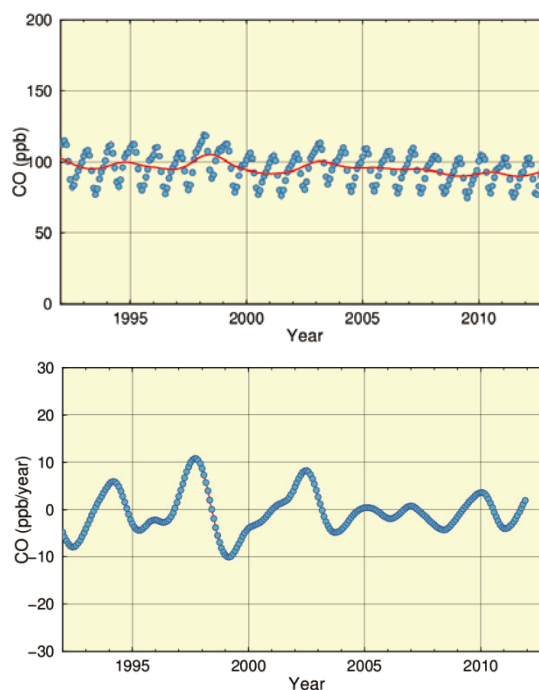


Fig. 8.1 Globally averaged monthly mean mole fraction of CO from 1992 to 2012, including deseasonalized long-term trend in red line (top) and annual growth rate (bottom).

Figure 8.1 shows globally averaged monthly mean CO mole fractions and their growth rates. Growth rates were high in 1993/1994, 1997/1998 and 2002, and low in 1992 and 1998/1999. The global annual mean mole fraction was about 92 ± 2 ppb in 2012, which was calculated irrespective of the difference in observation scales.

Figure 8.2 shows monthly mean mole fractions of CO for each 30° latitudinal zone. Seasonal variations were observed in both hemispheres, with mole fractions being higher in winter. Amplitudes of the seasonal cycle were larger in the Northern Hemisphere than in the Southern Hemisphere.

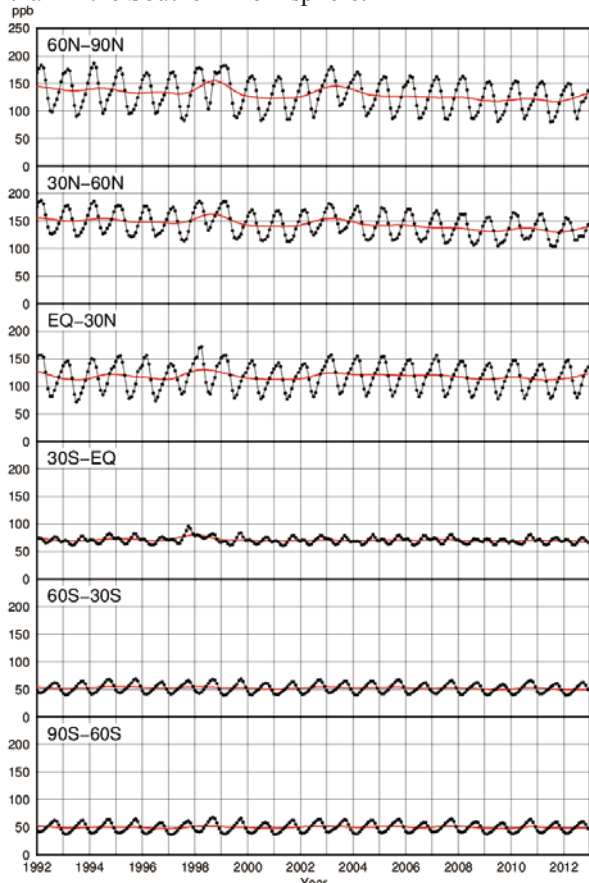


Fig. 8.2 Monthly mean mole fractions of CO from 1992 to 2012 for each 30° latitudinal zone (dots) and their deseasonalized long-term trends (red lines).

Figure 8.3 summarizes deseasonalized long-term trends for each 30° latitudinal zone and their growth rates. The CO mole fractions were the highest in northern mid-latitudes. There was a decline in CO mole fractions around 1992, almost coinciding with the decrease in the growth rate of CH_4 mole fractions, most likely due to variations in their common sink (reaction with OH). The enhanced stratospheric ozone depletion due to increased volcanic aerosols following the eruption of Mount Pinatubo in 1991 may have increased atmospheric OH radicals, which react with both CO and CH_4 (Dlugokencky *et al.*, 1996).

Increases in CO mole fractions were observed from 1997 to 1998 in northern latitudes and in southern low latitudes. These increases were attributed to large biomass burning events in Indonesia in late 1997 and in Siberia in the summer and autumn of 1998 (Novelli *et al.*, 1998).

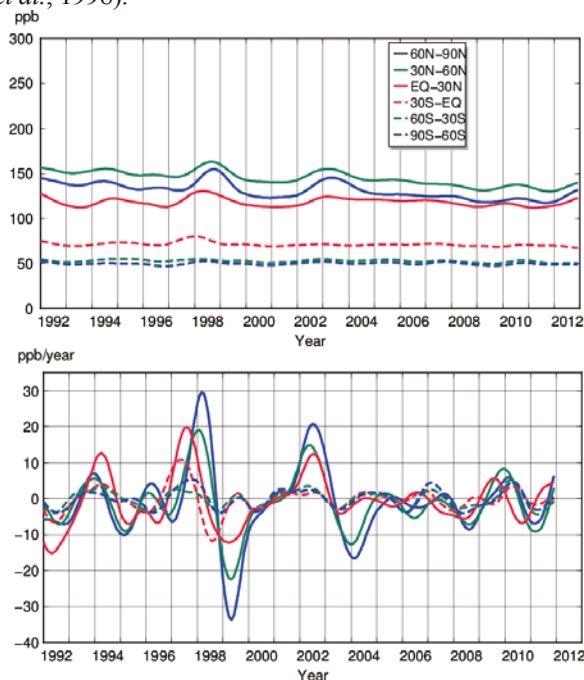


Fig. 8.3 Deseasonalized long-term trends of CO for each 30° latitudinal zone (top) and their growth rates (bottom).

The CO mole fractions returned to normal after 1999, but the growth rates in the Northern Hemisphere increased substantially again in 2002. The latter may have been due to large biomass burning. Large-scale boreal forest fires occurred in Siberia and North America from 2002 to 2003. Large forest fires also occurred in Russia in summer 2010 which is reflected in the data in the bottom panel of Figure 8.3.

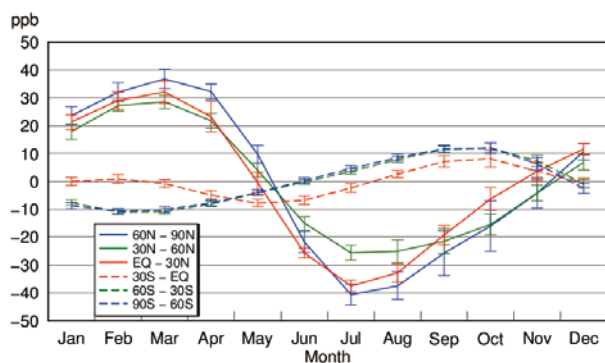


Fig. 8.4 Average seasonal cycles of CO mole fraction for each 30° latitudinal zone obtained by subtracting long-term trends from the zonal mean time series. Error bars represent the range of $\pm 1\sigma$ calculated for each month. (Average from 1992 to 2012)

Seasonal cycle of CO mole fraction in the atmosphere

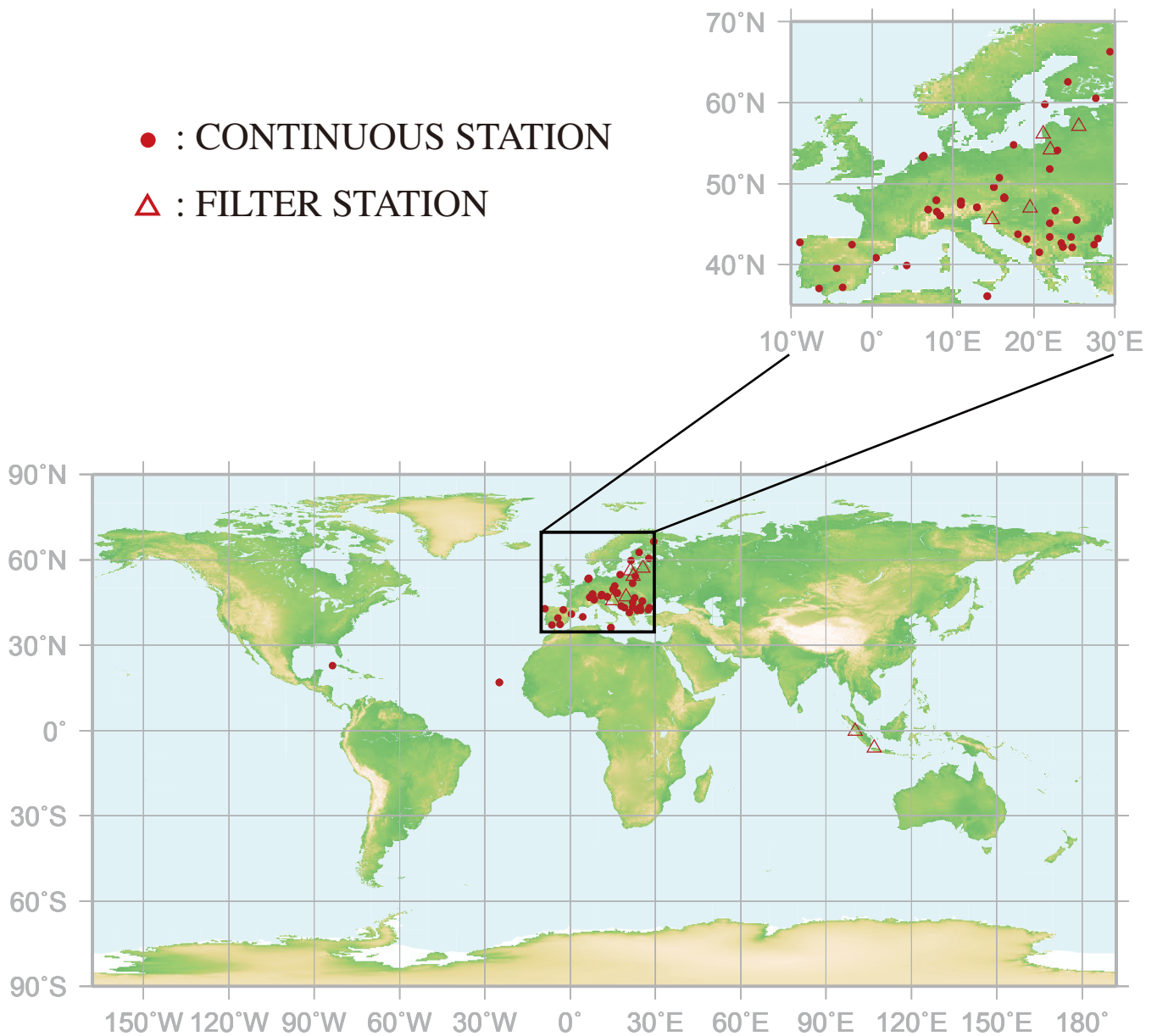
Figure 8.4 shows average seasonal cycles in the mole fraction of CO for each 30° latitudinal zone. The seasonal cycle is driven mainly by seasonal variations in OH abundance as a CO sink. This seasonality and a short lifetime of about a few months resulted in a sharp decrease in early summer followed by a relatively slow increase in autumn. The levelling-off in the beginning of the year observed in the southern low latitudes may be attributed to the transport of CO from the Northern Hemisphere.

9.

**NITROGEN MONOXIDE (NO)
AND
NITROGEN DIOXIDE (NO₂)**

● : CONTINUOUS STATION

△ : FILTER STATION



This map shows locations of the stations that have submitted data for monthly mean mole fraction.

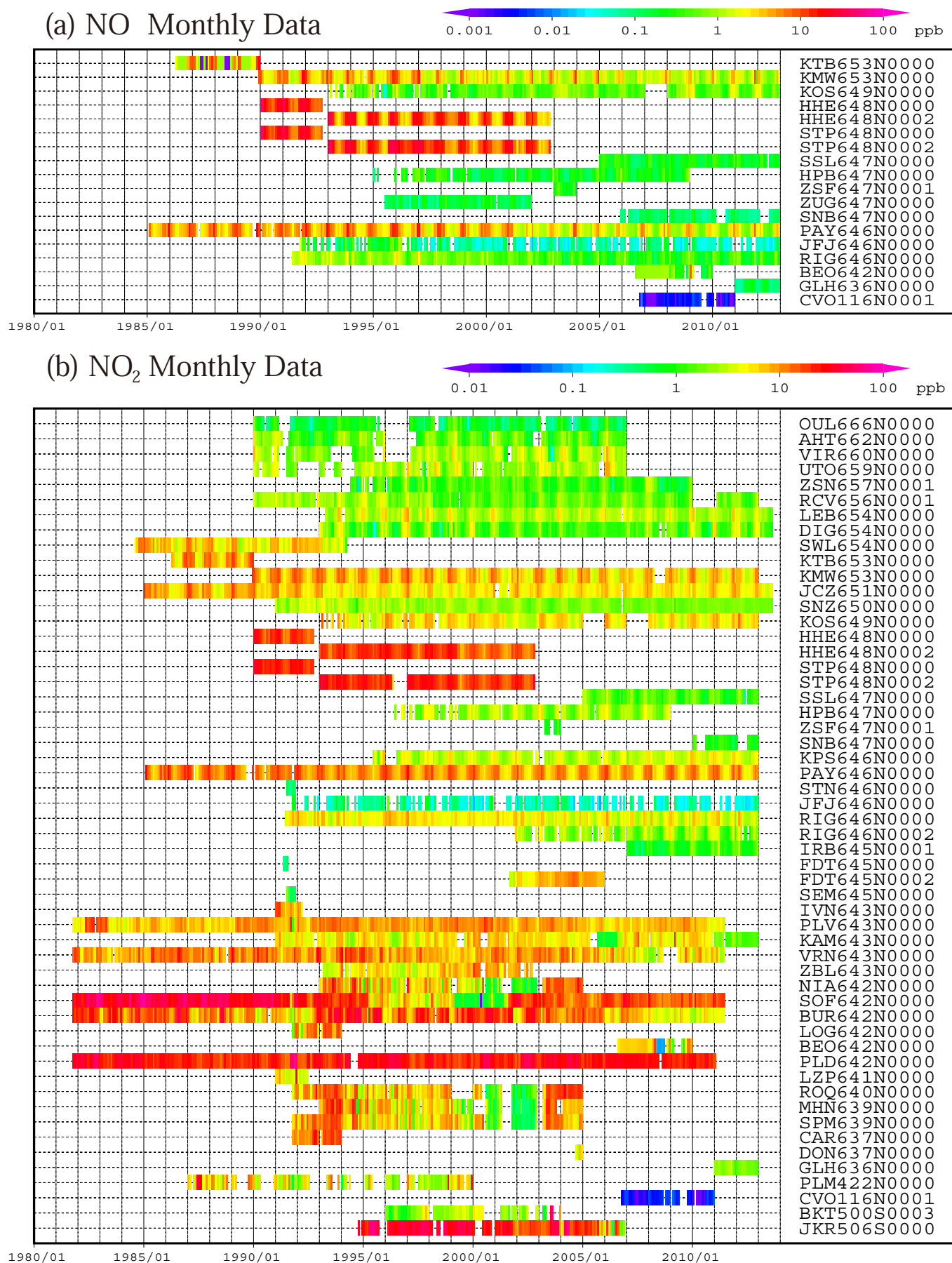


Plate 9.1 Monthly mean (a) NO and (b) NO₂ mole fractions that have been reported to the WDCGG. The mole fractions are illustrated in different colours. The sites are listed in order from north to south.

9. NITROGEN MONOXIDE (NO) AND NITROGEN DIOXIDE (NO₂)

Basic information on NO and NO₂ with regard to environmental issues

Nitrogen oxides (NO_x, *i.e.*, NO and NO₂) are not greenhouse gases. Nevertheless, these compounds have a central regulatory role in the free radical and oxidising chemistry of the troposphere. These chemical cycles regulate the lifetime of methane and the production of tropospheric O₃ and secondary aerosols, all of which have important roles in the natural and anthropogenic greenhouse effect. The O₃ produced in the atmosphere as a result of the availability of nitrogen oxides in the atmosphere can affect vegetation growth and human health.

Sources of NO_x include energy production, transport, lightning, soils and biomass burning (Reis *et al.*, 2009). They constitute major causes of acid rain and deposition. The dominant sink of NO_x in the atmosphere is its conversion into nitric acid (HNO₃) and peroxyacetyl nitrate (PAN), which are eventually removed by dry or wet deposition. In some cases, NO_x is removed from the atmosphere directly by dry deposition. NO_x abundance varies in both space and time because of their short lifetimes and uneven source distribution. Some regional assessments are done based on satellite information to clarify such variations and trends.

Annual variation of NO and NO₂ mole fractions in the atmosphere

The observational stations that have submitted data for NO and NO₂ to the WDCGG are shown on the map at the beginning of this chapter. Most of these stations are located in Europe.

The monthly mean mole fractions of NO and NO₂ reported to the WDCGG are shown in Plate 9.1, in which different mole fraction levels are plotted in different colours. Data for NO_x are reported in various units, *i.e.*, ppb, µg/m³-25°C, µg/m³-20°C, µgN/m³-25°C and mg/m³-25°C. Units other than ppb were converted to ppb using the formulas:

$$\begin{aligned} X_p [\text{ppb}] &= (R \times T / M / P_0) \times 10 \times X_g [\mu\text{g}/\text{m}^3] \\ X_p [\text{ppb}] &= (R \times T / M / P_0) \times 10^4 \times X_g [\text{mg}/\text{m}^3] \\ X_p [\text{ppb}] &= (R \times T / M_N / P_0) \times 10 \times X_g [\mu\text{gN}/\text{m}^3] \end{aligned}$$

where R is the molar gas constant (8.31451 [J/K/mol]),

T is the absolute temperature reported from each station,

M is the molecular weight of NO (30.00614) or NO₂ (46.00554),

M_N is the atomic weight of N (14.00674), and

P₀ is the standard pressure (1013.25 [hPa]).

If temperature is not reported by a station, T is taken to be 298K (25 °C).

The distributions of NO and NO₂ are spatially non-uniform and variable over time. Due to the high temporal variability in the mole fraction of NO₂ at each observational site, it was difficult to identify a long-term trend. A number of stations located in southern Europe showed higher mole fractions, and some stations reported increased NO₂ in winter.

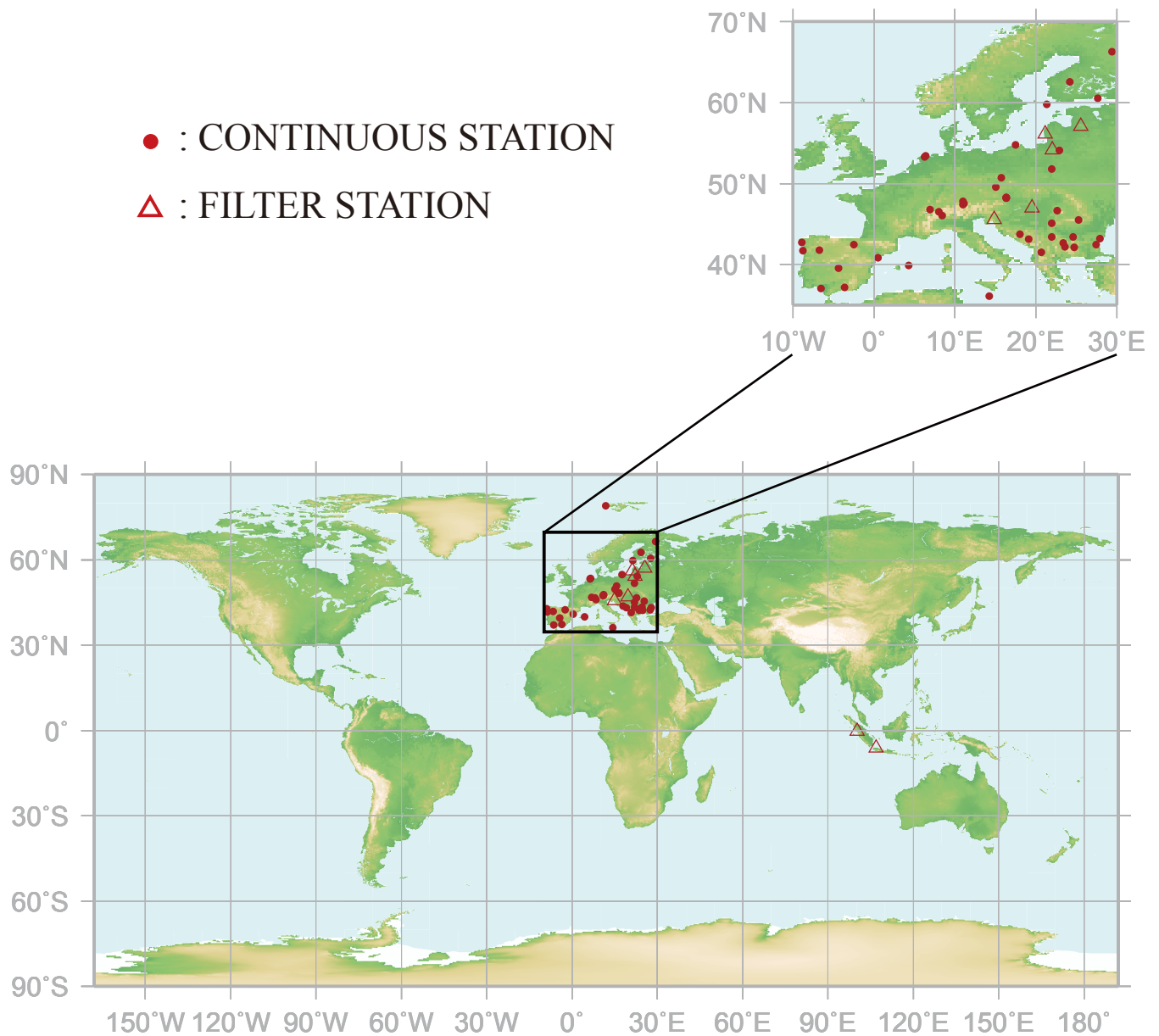
As there are few observational sites for NO, it was difficult to identify whether the global average NO mole fraction increases or decreases.

10.

SULPHUR DIOXIDE

(SO₂)

● : CONTINUOUS STATION
△ : FILTER STATION



This map shows locations of the stations that have submitted data for monthly mean mole fraction.

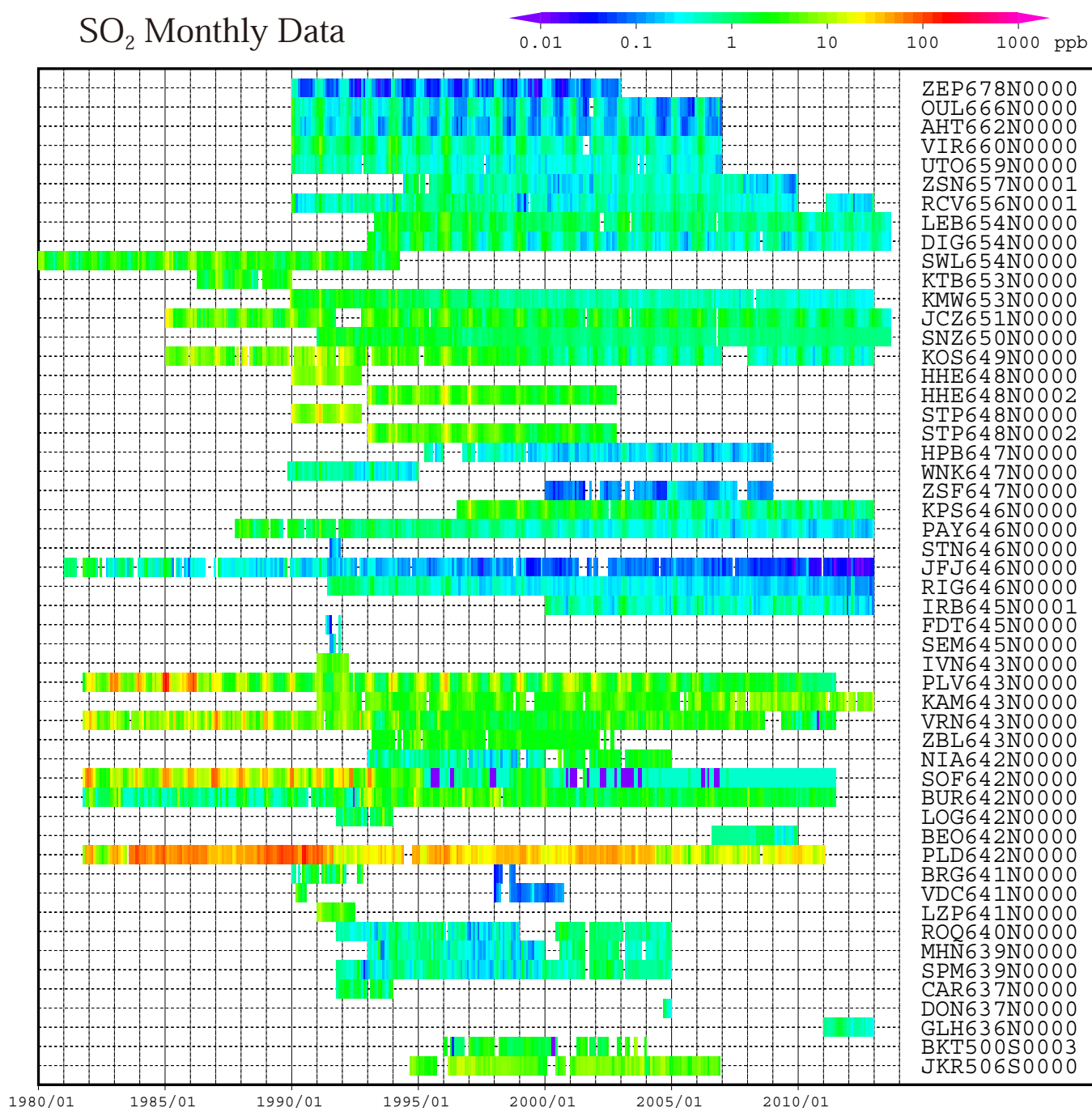


Plate 10.1 Monthly mean SO₂ mole fractions that have been reported to the WDCGG. The mole fractions are illustrated in different colours. The sites are listed in order from north to south.

10. SULPHUR DIOXIDE (SO₂)

Basic information on SO₂ with regard to environmental issues

Sulphur dioxide (SO₂) is not a greenhouse gas, but it is a precursor of atmospheric sulphuric acid (H₂SO₄) and sulphate aerosol. SO₂ is oxidized by hydroxyl radicals (OH) to form sulphuric acid, which then becomes aerosols through photochemical gas-to-particle conversion. While SO₂ reacts much more slowly with OH than does NO₂, SO₂ dissolves readily in suspended liquid droplets in the atmosphere. The global sulphur cycle affects atmospheric chemistry, including tropospheric ozone (Berglen *et al.*, 2004).

Sources of SO₂ include fossil fuel combustion by industry, biomass burning, volcanic release and the oxidation of dimethylsulphide (DMS) from the oceans (IPCC, 2007). Major SO₂ sinks are oxidation by OH and deposition onto wet surfaces. Anthropogenic SO₂ has caused acid rain and deposition throughout the industrial era. The mole fractions of SO₂ have shown large variations in both space and time because of the short lifetime and uneven anthropogenic source distribution of SO₂.

Annual variation of SO₂ mole fraction in the atmosphere

The observational sites that have submitted data for SO₂ to the WDCGG are shown on the map at the beginning of this chapter. Most of these stations are located in Europe.

The monthly mean mole fractions of SO₂ that have been reported to the WDCGG are shown in Plate 10.1, with different mole fraction levels illustrated in different colours. Data for SO₂ are reported in various units, *i.e.*, ppb, µg/m³, mg/m³ and µgS/m³. Units other than ppb were converted to ppb using the formulas:

$$\begin{aligned}X_p [\text{ppb}] &= (R \times T / M / P_0) \times 10 \times X_g [\mu\text{g}/\text{m}^3] \\X_p [\text{ppb}] &= (R \times T / M / P_0) \times 10^4 \times X_g [\text{mg}/\text{m}^3] \\X_p [\text{ppb}] &= (R \times T / M_S / P_0) \times 10 \times X_g [\mu\text{gS}/\text{m}^3]\end{aligned}$$

where R is the molar gas constant (8.31451 [J/K/mol]),

T is the absolute temperature reported from each station,

M is the molecular weight of SO₂ (64.0648),

M_S is the atomic weight of S (32.066), and

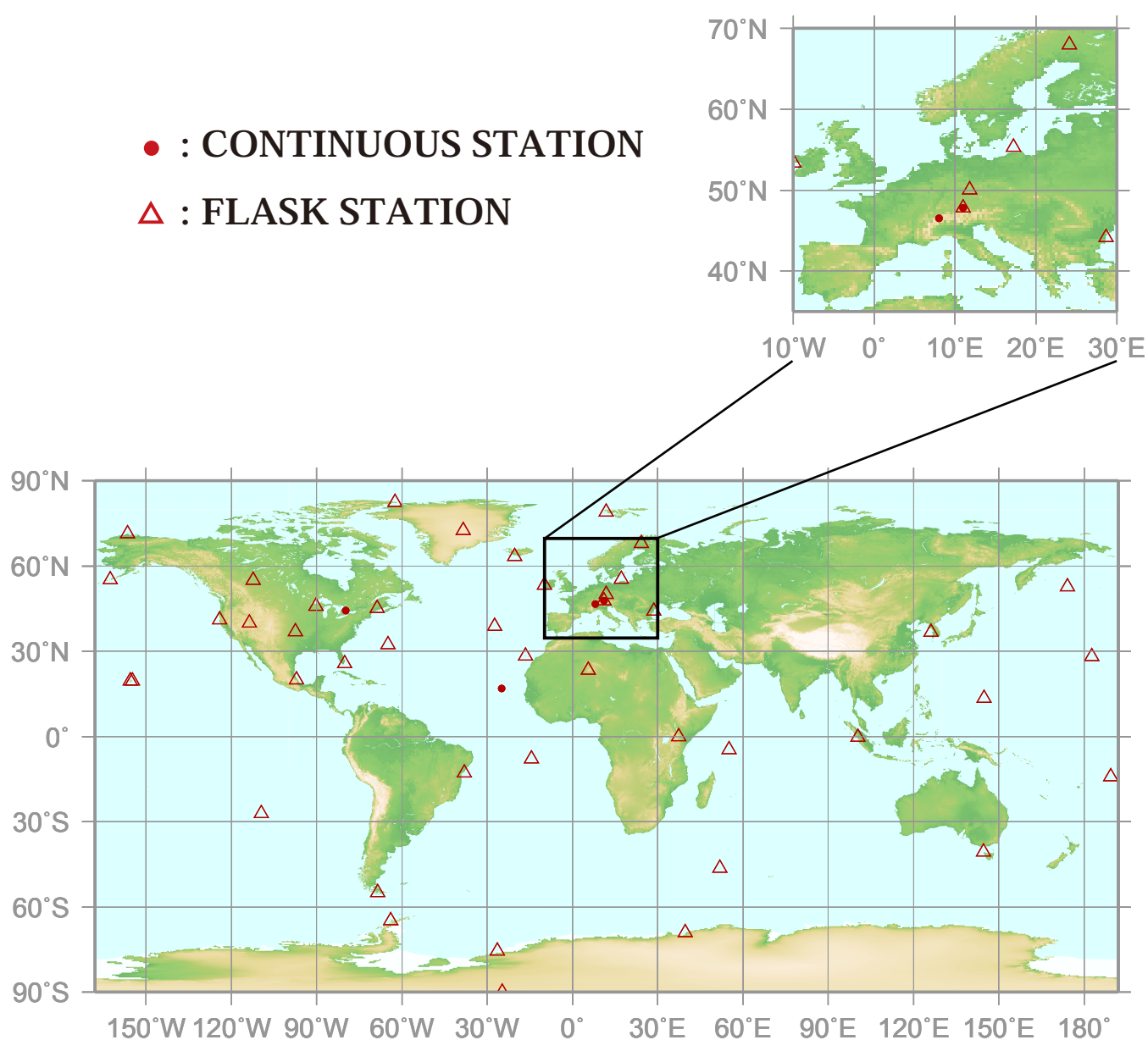
P₀ is the standard pressure (1013.25 [hPa]).

If temperature is not reported by a station, T is taken to be 298K (25 °C).

Although some stations in southern Europe have reported higher mole fractions, it has been difficult to identify the magnitude and direction of the trend for SO₂.

11.

VOLATILE ORGANIC COMPOUNDS (VOCs)



The map shows locations of the stations that have submitted data for mole fraction of VOCs (ethane and propane).

Ethane Data

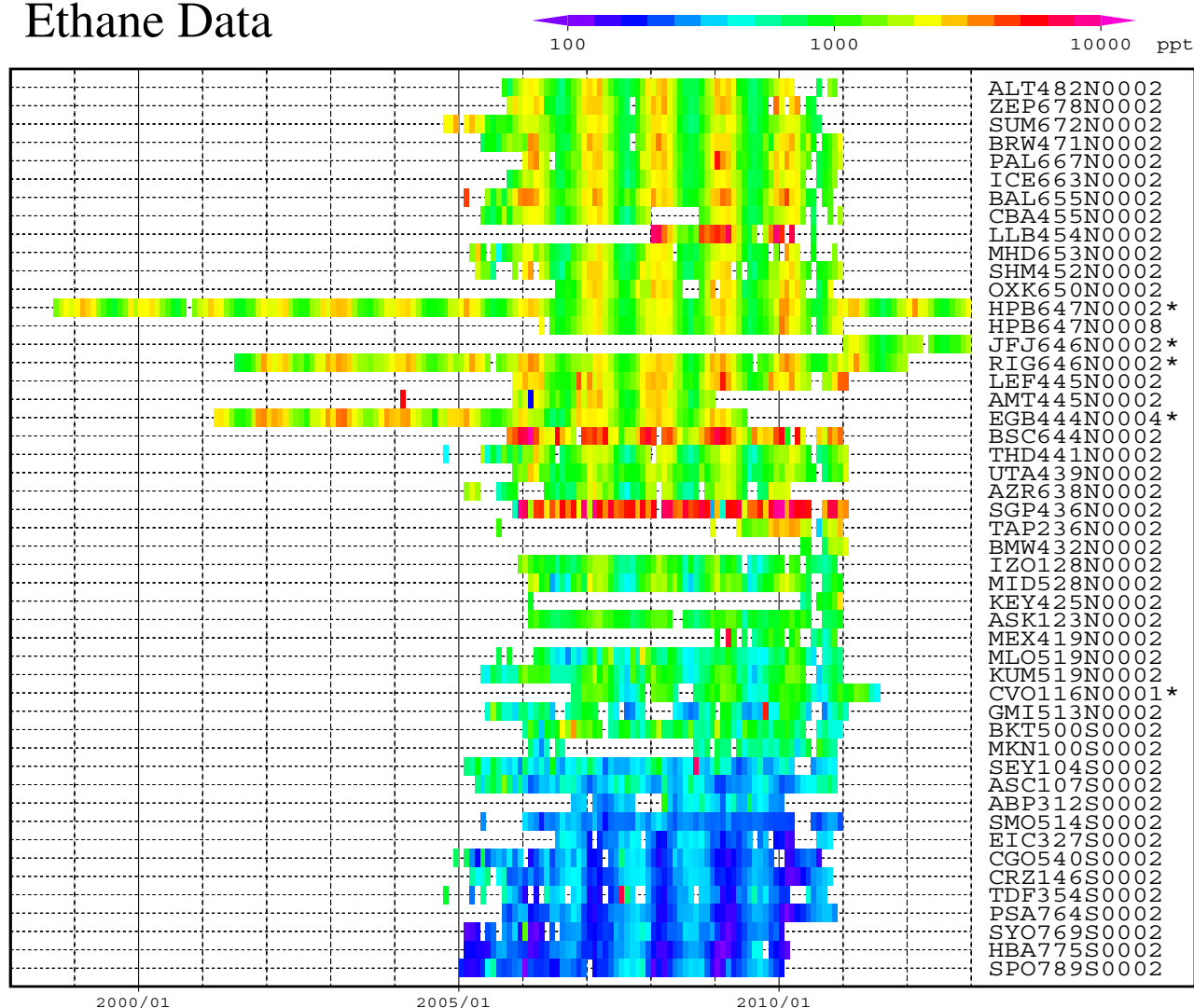


Plate 11.1 Monthly mean ethane mole fractions that have been reported to the WDCGG. The mole fractions are illustrated in different colours. The sites are listed in order from north to south. The data from the sites with an asterisk at the end of the station index are derived from continuous data.

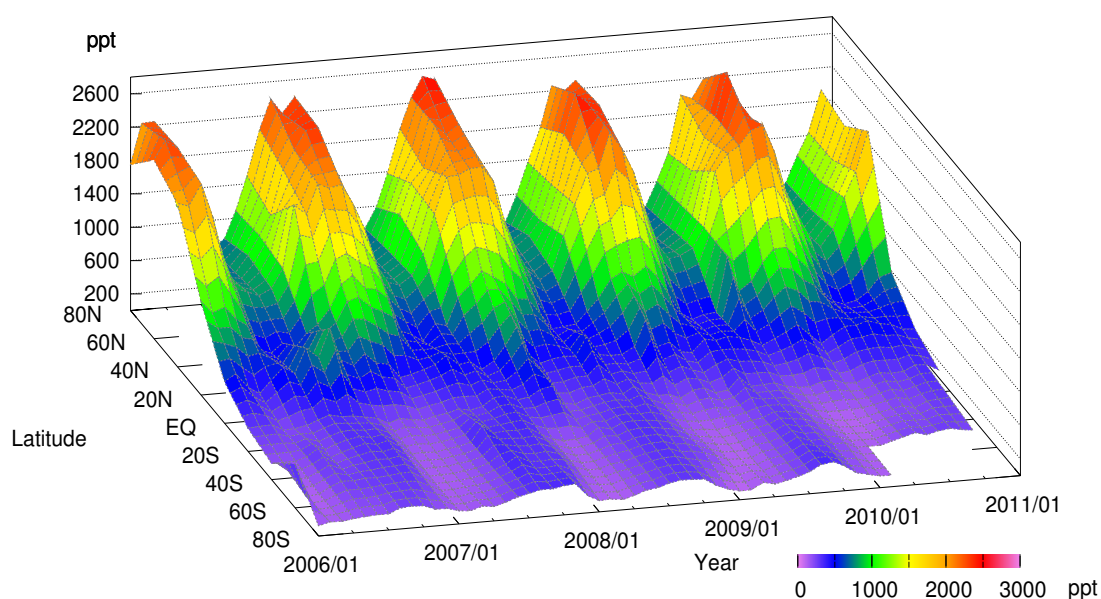


Plate 11.2 Variation of zonally averaged monthly mean ethane mole fractions. The zonally averaged mole fractions are calculated for each 20° zone.

Propane Data

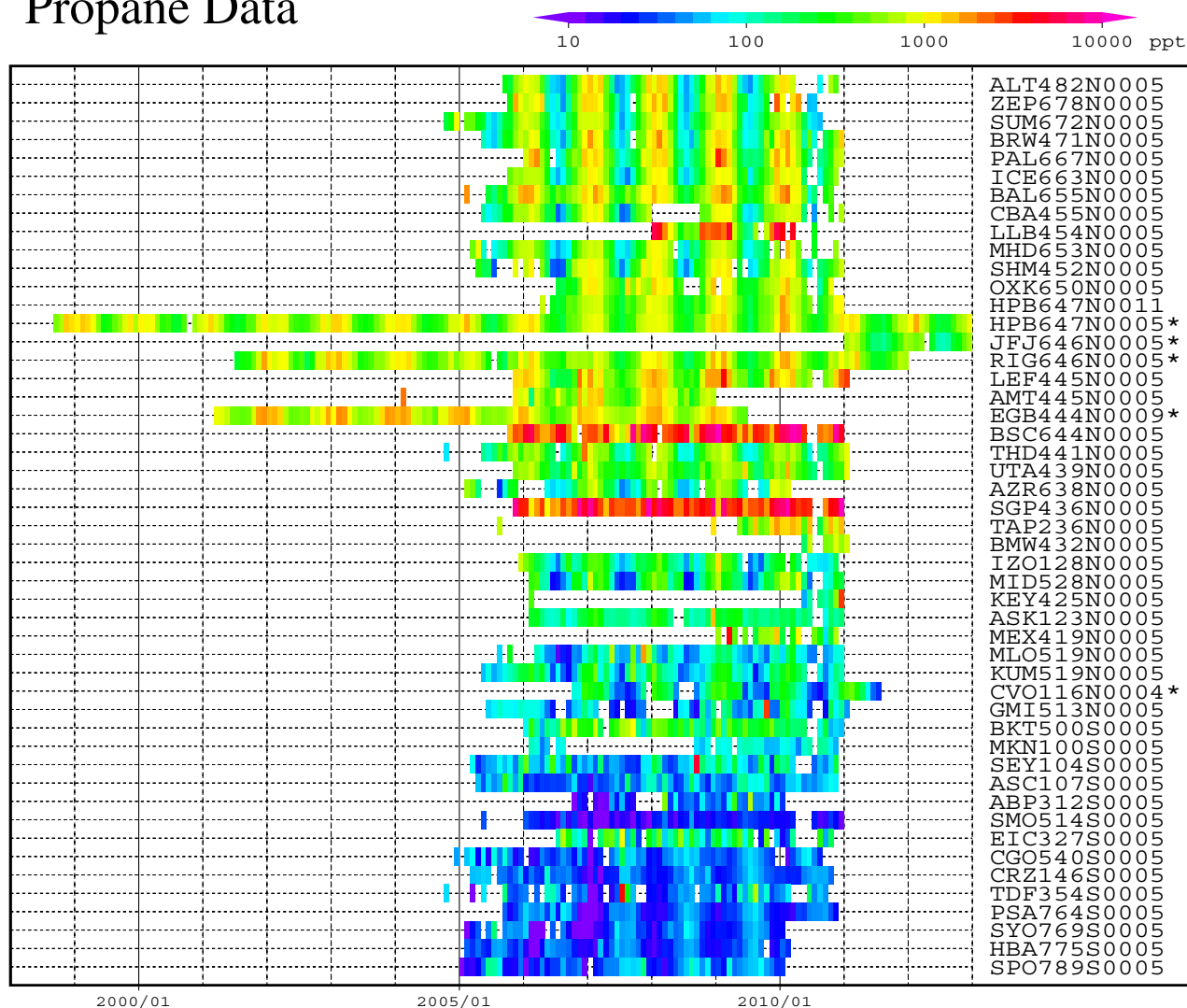


Plate 11.3 Monthly mean propane mole fractions that have been reported to the WDCGG. The mole fractions are illustrated in different colours. The sites are listed in order from north to south. The data from the sites with an asterisk at the end of the station index are derived from continuous data.

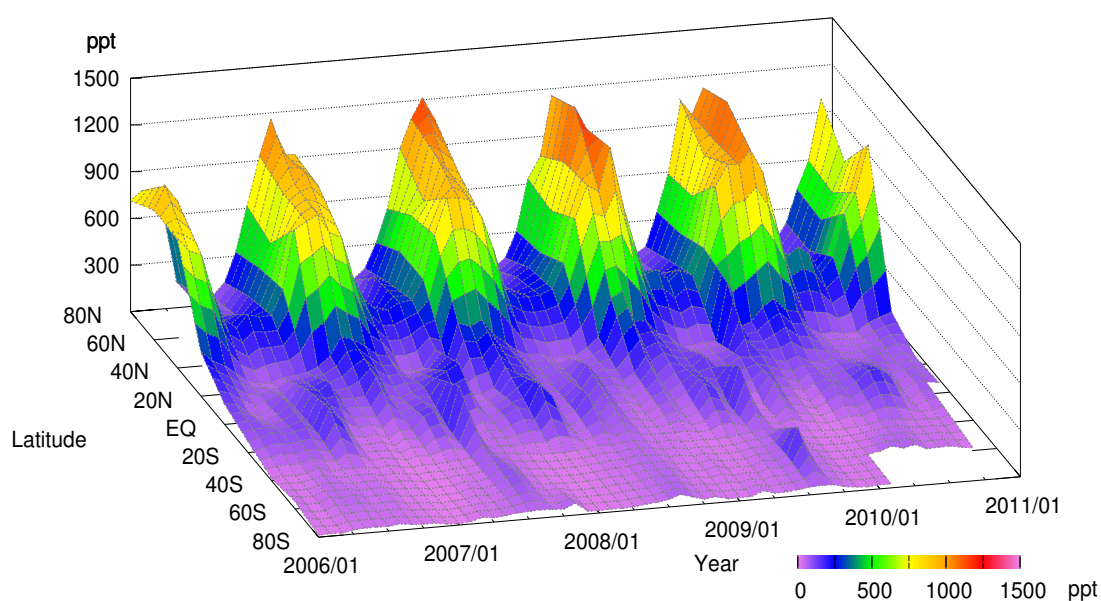


Plate 11.4 Variation of zonally averaged monthly mean propane mole fractions. The zonally averaged mole fractions are calculated for each 20° zone.

11. VOLATILE ORGANIC COMPOUNDS(VOCs)

Basic information on VOCs with regard to environmental issues

Volatile organic compounds (VOCs) have a variety of roles in atmospheric chemistry. They are major contributors to photochemical air pollution on both urban and regional scales and together with NO_x also impact ozone in the background troposphere. Some VOCs are injurious to human health at modest concentrations; they are also precursors to aerosols at quite low concentrations. VOCs can serve as tracers of many atmospheric processes and emission sources. VOC molecules occur in many forms and have both natural and anthropogenic sources. The GAW Programme has sought to focus measurements on species which help to provide understanding of a wide range of atmospheric properties, and which can mostly be measured using currently approved techniques. Table 11.1 shows the molecules selected for measurements within GAW, with reasons for their selection (WMO, 2007c). The main reasons for measurements within GAW are associated with their use as tracers of the source types of greenhouse gases such as methane, to provide quantitative information on the extent of atmospheric processing by hydroxyl radicals (OH) and other oxidants, and as precursors to organic acids and aerosols, particularly organic aerosols and sulphate aerosols. In addition, VOC measurements provide valuable information for air

mass characterisation at some stations to identify local sources of pollution. All of this information is of value to atmospheric modellers both in terms of input parameters and as constraints to model results.

Some GAW stations have the analytical capability to produce high quality measurements of an extended range of VOCs. Where appropriate calibration and quality assurance can be provided, these data should wherever possible be provided to WMO databases for wider scientific dissemination (WMO, 1996). The current global network of VOC measurements is shown on the front page of this chapter. In many respects it is identical to that set up by NOAA to measure stable greenhouse gases and carbon monoxide with flask samples collected every 1 or 2 weeks, but it does include a number of sites where selected gases are measured at a higher frequency. These include three sites in Europe, one in Greenland (Summit) and one on the Island of Cape Verde. Unfortunately, Egbert, a Canadian site which provided valuable information on the VOC composition over North America discontinued high frequency VOC measurements in 2010. It is anticipated, however, that the number of sites with high frequency measurements will increase significantly in the next few years in the European region and the adjacent Arctic.

Table 11.1 Molecules selected for measurements within the GAW VOC Programme with reasons for their selection

Molecule	Lifetime (OH=10 ⁶ cm ⁻³)	Importance to GAW	Analysis Method	Network Type
Ethane	1.5 months	• Source of methane • Natural sources • Biomass burning • Fossil fuel • Ocean production(S.hemisphere) • Trend in size of seasonal cycle	GC/FID	Global
Propane	11 days	• Source of methane • Natural sources • Biomass burning • Fossil fuel • Ocean production(S.hemisphere)	GC/FID	Global
Acetylene	15 days	• Motor vehicle tracer • Biomass burning tracer • Ratios to the other hydrocarbons • Trends	GC/FID	Global
Isoprene	3 hours	• Biosphere product • Sensitive to temperature/land use/climate change • O ₃ precursor • Oxidizing capacity • Precursor to formaldehyde	GC/FID PTR-MS	Mid latitudes and tropics
Formaldehyde	1 day	• Indicator of isoprene oxidation • Biomass burning • Comparison with satellites • Trends	DOAS	Small number of sites in Tropics for comparison with satellites
Terpenes	1.5 hours	• Precursors to organic aerosols	GC/MS PTR-MS	Selected sites in forested areas
Acetonitrile	0.5-1 year	• Biomass burning indicator • Biofuel burning indicator	GC/MS PTRMS	Global
Methanol	12 days	• Sources in the biosphere (methane oxidation) • Abundant oxidation product	GC/FID PTR-MS	Global
Ethanol	4 days	• Tracer of alternative fuel usage	GC/FID PTR-MS	Global
Acetone	1.7 months	• Abundant oxidation product • Free radical source in the upper troposphere	GC/FID PTR-MS	Global
DMS	2 days	• Major natural sulphur source • Sulphate aerosol precursor • Tracer of marine bioproductivity	GC/FID PTR-MS	Global Marine
Benzene	10 days	• Tracer of combustion • Biomass burning indicator	GC/FID GC-MS	Global
Toluene	2 days	• Ratio to benzene used for air massage • Precursor to particulates	GC/FID GC-MS	Global
Iso/normal Butane	5 days	• Chemical processing indicator • Lifetime/ozone production • Indication of halogen chemistry	GC/FID GC-MS	Global
Iso/normal Pentane	3 days	• Ratio provides impact of NO ₃ chemistry • Indication of halogen chemistry	GC/FID GC-MS	Global

GC/FID is Gas Chromatography – Flame Ionization Detection
DOAS is Different Optical Absorption Spectrometry

GC/MS is Gas Chromatography – Mass Spectrometry
PTRMS is Proton Transfer Reaction Mass Spectrometry

Seasonal Variation of VOCs and trends

As with all other measurements within GAW which are designed to study atmospheric composition, an important use of the data is to evaluate trends over time. Sufficient data are available for many individual molecules, particularly the non-methane hydrocarbons (NMHCs) both from the American and the European continents. The current database also contains much information on the seasonal variation of both natural and anthropogenic hydrocarbons. The map on the front page of this chapter shows the network currently reporting VOC data, differentiating between sites where flask samples are collected for a limited set of NMHC measurements, and sites with measurements of a wider range of VOCs collected in a semi-continuous manner.

The review of the data record from the different sites reveals some interesting and important features, both concerning the characteristics of the measurement sites and the global and seasonal atmospheric behavior of VOCs. The data from many of the sites illustrate a regular seasonal behavior, with concentration maxima being observed in the winter and minima in the summer. Sites in Plate 11.1 and 11.3 are sorted by the station latitude, with the northern most site at the top, and South Pole at the bottom. NMHCs show a latitudinal concentration gradient, with concentrations decreasing from the poles to the equator, most notable during the winter. A number of stations, however, deviate from this behavior, showing higher VOC levels than other sites at similar latitude, with these deviations being more pronounced during the winter. Examples for this behavior are the stations: Lac La Biche (LLB), Constanta, Black Sea (BSC), and the Southern Great Plains (SGP). These elevated NMHC data are clear evidence of these sites being influenced by nearby local sources of these gases, elevating atmospheric NMHC mole fractions significantly above their latitudinal background. Consequently, while these data are of value for assessing the burden and long-term change of VOCs in these regions, they are not suitable for consideration in the reconstruction of the latitudinal background of these gases.

Focusing on ethane and propane, which have rather longer lifetimes amongst VOCs, Plate 11.1 and 11.3 show monthly mean mole fractions of ethane and propane, respectively, from 1998 to 2012 at each of the stations reporting to the database at WDCGG; and Plate 11.2 and 11.4 show the global ethane and propane distribution, respectively, from the flask network between 2006 to 2010 projected on to a continuous surface, with vertical and horizontal coordinates, respectively, of the mole fraction in ppt and latitude in degrees.

The ethane mole fraction shows a large seasonal variation which is mostly associated with its removal from the atmosphere by hydroxyl radical chemistry, leading to lower values in the summer months. The stronger latitudinal gradients in winter in the Northern Hemisphere compared to summer are also associated to corresponding larger OH gradients in winter. Absolute mole fractions are much larger in the Northern Hemisphere reflecting the preponderance of sources, mostly from oil and gas extraction in this region; Southern Hemisphere sources are dominated by emissions from biomass burning with a much smaller contribution associated with fossil fuel usage. A consideration of the budget of ethane in the atmosphere and its relevance to understanding the sources of methane was published by Simpson *et al.* (2012). This study concluded that the slow-down in the growth of atmospheric methane observed in 1999 to 2006 was predominantly associated with improved containment of methane emissions from processes such as flaring during oil and gas extraction. This result is a good example of how VOCs can be used as tracers to understand the behavior of important greenhouse gases.

The 38th edition of the *Data Summary* includes results for the global distribution of propane. Mole fractions of propane are generally lower than for ethane owing to shorter lifetimes of propane by about a factor of 5 and roughly 50% lower sources for propane (mole-base). The propane distribution shows similar general features as for ethane, however, stronger seasonality and more pronounced latitudinal gradients in the Northern Hemisphere (NH) during winter. Due to its shorter atmospheric lifetime compared to ethane, the NH seasonal cycle shows a maximum close to mid-winter, whereas the ethane maximum is shifted by about 2 months towards spring owing to ethane reacting slower to the changes in OH. The seasonal cycle for propane is more pronounced than for ethane mainly due to its higher reactivity. The higher rate decay of propane in summertime during transport of the air masses to remote observation sites compared to ethane contributes to the relatively lower summer minima seen in the representation of the network data.

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APPENDICES

CALIBRATION AND STANDARD SCALES

1. Calibration System in the GAW Programme

Under the Global Atmosphere Watch (GAW) Programme, the Central Calibration Laboratories (CCLs) are assigned to host a Primary (Reference) Standard/scale, while the World Calibration Centres (WCCs) are responsible for the scale propagation to the stations via distribution of calibration standards for certain compounds, conducting instrument calibrations, comparison campaigns, station audits and providing

training to the station personnel. A Reference Standard/scale is designated for each variable to be used for all GAW measurements of that variable. Table 1 lists the organizations that serve as WCCs and CCLs for GAW (WMO, 2011). For CFCs and SO₂, no central facilities or quality control systems have so far been established within the GAW Programme.

Table 1. Overview of the GAW Central Calibration Laboratories (GAW-CCL, Reference Standard) and World Calibration Centres for greenhouse and other related gases. The World Calibration Centres have assumed global responsibilities, except where indicated (Am, Americas; E/A, Europe and Africa; A/O, Asia and the South-West Pacific)

Compounds	Central Calibration Laboratory (Host of Primary Standard)	World Calibration Centre
Carbon Dioxide (CO ₂)	NOAA/ESRL	NOAA/ESRL (round robin) Empa (audits)
carbon isotopes	MPI-BGC	
Methane (CH ₄)	NOAA/ESRL	Empa (Am, E/A) JMA (A/O)
Nitrous Oxide (N ₂ O)	NOAA/ESRL	IMK-IFU
Chlorofluorocarbons (CFCs)		
Sulphur Hexafluoride (SF ₆)	NOAA/ESRL	KMA
Molecular Hydrogen (H ₂)	MPI-BGC	
Surface Ozone (O ₃)	NIST	Empa
Carbon Monoxide (CO)	NOAA/ESRL	Empa
Volatile Organic Compounds (VOCs)	NPL (8 components)	KIT/IMK-IFU
Sulphur Dioxide (SO ₂)		
Nitrogen Oxides (NO _x)		FZ-Jurich

2. Carbon Dioxide (CO₂)

In 1995, the National Oceanic and Atmospheric Administration's Earth System Research Laboratory (NOAA/ESRL, formerly CMDL; Climate Monitoring and Diagnostics Laboratory) in Boulder, Colorado, USA, took over the role of the Central Calibration Laboratory (CCL) from the Scripps Institution of Oceanography (SIO) in San Diego, California, USA. Since then, NOAA/ESRL has served as the CCL responsible for the maintenance of the GAW Primary Standard for CO₂. As CCL for CO₂, NOAA/ESRL maintains a high-precision manometric system for absolute calibration of CO₂ as the reference for GAW measurements throughout the world (Zhao *et al.*, 1997), as well as carrying out Round-Robin in the function of

WCC. It has been recommended that the standards of the GAW measurement laboratories be calibrated at least every three years at the CCL (WMO, 2012).

Under the WMO calibration system, there have been several calibration scales for CO₂, *e.g.*, SIO-based X74, X85, X87, X93 and X2002 scales and the NOAA/ESRL-based WMO Mole Fraction Scale partially based on previous SIO scales. The CCL adopted the WMO X2005 scale, reflecting historical manometric calibrations of the CCL's set of cylinders and the possible small differences between SIO and NOAA/ESRL calibrations. The most current WMO Mole Fraction Scale is the WMO-X2007 scale.

To assess the differences in standard scales among

measuring laboratories, NOAA/ESRL organizes intercomparisons or Round Robin experiments endorsed by WMO. It is recommended that round-robins are repeated at least once every three years. Many laboratories participated in the experiments organized in 1991–1992, 1995–1997, 1999–2000, 2002–2006, and 2009–2012. Table 2 shows the results of the experiments performed in 2009–2012, in which the mole fractions measured by various laboratories are compared

with the mole fractions measured by NOAA/ESRL (<http://www.esrl.noaa.gov/gmd/ccgg/wmorr/results.php>). In addition, many laboratories compare their standards bilaterally or multilaterally.

Table 3 lists laboratories and sites that contributed to the present issue of the *Data Summary* with standard scales of reported data and history of participation in WMO intercomparison experiments.

Table 2. Round robin results for the mole fraction of carbon dioxide. Differences between the mole fractions measured by various laboratories and the mole fractions measured by NOAA (Laboratory minus NOAA, ppm).

Laboratory	Measurement Date	Mole fraction Difference (ppm)		
		Low 365-375 ppm	Medium 380-390 ppm	High 395-405 ppm
JMA	Jun-09	−0.00	−0.06	−0.04
MRI	Jul-09	−0.13	−0.05	−0.05
AIST	Sep-09	0.08	0.13	0.16
NIES	Oct-09	−0.03	−0.01	0.03
TU	Jan-10	0.15	0.21	0.28
CMA	Apr-10	−0.21	−0.14	−0.04
KMA/KGAWO	Jul-10	−0.54	0.01	0.05
KRISS	Jul-10	0.14	0.20	0.24
SNU	Nov-10	0.01	0.03	0.14
IPEN	Jul/Aug-11	−0.09	−0.03	−0.02
AEMET	Sep-12	0.03	−0.03	0.02
EC	Apr-09	−0.02	−0.02	−0.03
SIO	May-09	−0.20	−0.14	−0.11
AMERIFLUX	May-09	−0.02	0.03	0.02
SAWS	Aug-09	0.04	0.06	−0.04
NIWA	Nov-09	0.02	−0.02	−0.07
CSIRO	Jan-10	0.00	0.00	−0.10
PSU	Sep-10	0.06	0.08	−0.01
HU	Nov-10	−0.04	−0.07	−0.09
NCAR	Apr-11	−0.02	0.01	−0.02
OSU	Aug-11	0.02	0.08	0.00
ITM	May-09	−0.86	−0.73	−0.20
FMI	Jun-09	0.02	0.03	0.03
UBA/SCH	Aug-09	−0.00	0.01	0.04
UHEI-IUP	Aug-09	−0.04	−0.04	−0.08
MRI-BGC	Sep/Oct-09	−0.02	0.00	−0.00
UBA/ZUG	Dec-09	−0.07	−0.24	−0.48
HMS	Mar-10	0.07	−0.02	−0.03
CESI/RICERCA	Apr-10	−0.01	0.06	0.08
ENEA	Jun-10	−0.18	−0.30	−0.37
JRC	Oct-10	−0.04	0.02	0.00
WCC-EMPA	Nov-10	0.07	0.11	0.13

EMPA	Dec-10	-0.02	0.04	0.06
UEA	Dec-10	-0.26	-0.17	-0.12
KUP	Jan-11	-0.15	-0.11	-0.04
ECN	Mar-11	0.01	0.03	0.04
LSCE	Apr-11	0.06	0.01	0.12
RHUL	Jun-11	-0.08	-0.07	-0.11

Table 3. Status of standard scales and calibration/intercomparison for CO₂ at laboratories.

Laboratory	WDCGG Filename Code	Calibration Scale	WMO Inter-comparison
AEMET	IZO128N0000	WMO	91/92, 96/97, 99/00, 09/12
Aichi	MKW234N0000	WMO	
AIST	TKY236N0000	AIST	96/97, 99/00, 02/06, 09/12
BMKG & Empa	BKT500S0000	WMO	
BoM & CSIRO	CGO540S0000, CGO540S0010	WMO	
CMA	WLG236N0000	WMO	96/97, 99/00, 02/06, 09/12
CNR-ICES & DNA-IAA	JBN762S0000	WMO	
CSIRO	ALT482N0003, CFA519S0003, CGO540S0003, CRI215N0000, CYA766S0001, ESP449N0003, MAA767S0003, MLO519N0003, MQA554S0003, SIS660N0003, SPO789S0003	WMO	91/92, 96/97, 99/00, 02/06, 09/12
EC	ALT482N0000, ALT482N0005, CDL453N0000, CHM449N0000, CSJ451N0000, EGB444N0100, ESP449N0000, ETL454N0000, FSD449N0000, LLB454N0100, WSA443N0000, WSA443N0001	WMO	91/92, 96/97, 99/00, 02/06, 09/12
EMA	CAI130N0000		
Empa	JFJ646N0000	WMO	09/12
ENEA	LMP635N0001	WMO	91/92, 96/97, 99/00, 02/06, 09/12
FMI	PAL667N0000	WMO	02/06, 09/12
HKO	HKG222N0001, HKO222N0001	WMO	
	HKO222N0000	NIST	
HMS	HUN646N0000, KPS646N0000	WMO	91/92, 96/97, 99/00, 02/06, 09/12
IAFMS	CMN644N0001, CMN644N0002	WMO	91/92, 96/97, 02/06
IGP	HUA312S0000	WMO	
IMK-IFU	WNK647N0000, ZUG647N0014	WMO	99/00
INRNE	BEO642N0000	WMO	
IOEP	DIG654N0000	IOEP	
ITM	ZEP678N0000	WMO	96/97, 99/00, 09/12

JMA	MNM224N0000, RYO239N0000, YON224N0000	WMO	91/92, 96/97, 99/00, 02/06, 09/12
KMA	AMY236N0000, KSG762S0000	KRISS	02/06, 09/12
KSNU	ISK242N0000		
KUP	JFJ646N0003	WMO	
LSCE	AMS137S0000, BGU641N0000, LPO648N0000, MHD653N0002, PDM642N0000, PUY645N0000	WMO	91/92, 96/97, 99/00, 02/06, 09/12
	FIK635N0000		
MGO	BER255N0001, KOT276N0001, KYZ240N0001, STC652N0001, TER669N0001, TIK271N0000	WMO	
MMD	DMV504N0000	WMO	
MRI	TKB236N0002		91/92, 96/97, 99/00, 02/06, 09/12
NIER	GSN233N0103	WMO	
NIES	COI243N0000, HAT224N0000	NIES 95**	96/97, 99/00, 02/06, 09/12
NIMR	GSN233N0001	WMO	96/97
NIPR & Tohoku Univ.	SYO769S0000		Tohoku Univ.: 91/92, 96/97, 99/00, 02/06, 09/12
NIWA	BHD541S0000	WMO	91/92, 96/97, 99/00, 02/06, 09/12
NMA	FDT645N0002		
NOAA/ESRL	BRW471N0000, MLO519N0000, SMO514S0000, SPO789S0000, NOAA/ESRL flask network*	WMO	91/92, 96/97, 99/00, 02/06, 09/12
Osaka Univ.	SUI234N0000		
RIVM	KMW653N0000	NIST	
RSE	PRS645N0000	WMO	99/00, 02/06
Saitama	DDR236N0000, KIS236N0000, URW235N0000	WMO	
SAWS	CPT134S0000	WMO	99/00, 02/06, 09/12
Shizuoka Univ.	HMM234N0000		
UBA	BRT648N0000, DEU649N0000, LGB652N0000, NGL653N0000, SNB647N0000, SSL647N0000, SSL647N0002, WES654N0000, ZGT654N0000, ZSF647N0010, ZUG647N0000	WMO	91/92, 96/97, 99/00, 02/06, 09/12
Univ. Malta	GLH636N0000		

* NOAA/ESRL flask network:

ABP312S0001,ALT482N0001,AMS137S0001,ASC107S0001,ASK123N0001,AVI417N0001,AZR638N0001,BAL655N0001,BHD541S0001,
BKT500S0001,BME432N0001,BMW432N0001,BRW471N0001,BSC644N0001,CBA455N0001,CGO540S0001,CHR501N0001,CMO445N0001,
CRZ146S0001,EIC327S0001,GMI513N0001,GOZ636N0001,HBA775S0001,HPB647N0003,HUN646N0001,ICE663N0001,ITN435N0001,
IZO128N0001,KCO204N0001,KEY425N0001,KUM519N0001,KZD244N0001,KZM243N0001,LEF445N0001,LLB454N0001,LLN223N0001,
LMP635N0003,MBG476N0001,MEX419N0001,MHD653N0001,MID528N0001,MKN100S0001,MLQ519N0001,NMB123S0001,NWR440N0101,
OPW448N0001,PAL667N0001,POC900N0001,POC905N0001,POC905S0001,POC910N0001,POC910S0001,POC915N0001,POC915S0001,
POC920N0001,POC920S0001,POC925N0001,POC925S0001,POC930N0001,POC930S0001,POC935S0001,PSA764S0001,PTA438N0001,
RPB413N0001,SCS903N0001,SCS906N0001,SCS909N0001,SCS912N0001,SCS915N0001,SCS918N0001,SCS921N0001,SEY104S0001,
SGP436N0001,SHM452N0001,SMO514S0001,SPO789S0001,STC654N0001,STM666N0001,SUM672N0001,SYO769S0001,TAP236N0001,
TDF354S0001,THD441N0001,UTA439N0001,UUM244N0001,WIS631N0001,WLG236N0001,ZEP678N0001

**NIES 95 CO₂ scale is 0.10 to 0.14 ppm lower than WMO in a range between 355 to 385 ppm.

(Machida *et al.*, WMO/GAW Report No. 186, 26-29, 2009.)

3. Methane (CH₄)

The GAW Programme has established two WCCs for CH₄, the Swiss Federal Laboratory for Materials Testing and Research (Empa), Dübendorf, Switzerland; and the Japan Meteorological Agency (JMA), Tokyo, Japan (WMO, 2007). In addition, the Central Calibration Laboratory for CH₄ has been established at NOAA/ESRL (Dlugokencky *et al.*, 2005; WMO, 2007).

The NOAA04 scale has been designated as the Primary scale of the GAW Programme. This scale results in CH₄ mole fractions that are a factor of 1.0124 higher than the previous NOAA scale (Dlugokencky *et al.*, 2005).

Table 4 summarizes the CH₄ standard scales used by laboratories contributing to the WDCGG and lists tentative multiplying conversion factors applied for analysis in this issue of the *Data Summary*. The

standard is the NOAA04 scale, and conversion factors were calculated from the results of comparisons with other laboratories performed bilaterally or multilaterally before the establishment of the GAW Standard.

The former CMDL scale is lower than an absolute gravimetric scale (Aoki *et al.*, 1992) by ~1.5% (Dlugokencky *et al.*, 1994) and lower than the AES (Atmospheric Environment Service, currently EC) scale by a factor of 1.0151 (Worthy *et al.*, 1998). The former CMDL scale can be converted to the Tohoku University standard by multiplying by 1.0121 (Dlugokencky *et al.*, 2005). The conversion factors $1.0124 / 1.0151 = 0.9973$ and $1.0124 / 1.0121 = 1.0003$ have been adopted for comparisons with the NOAA04 scale.

Table 4. Status of the standard scales of CH₄ at laboratories with conversion factors.

Laboratory	WDCGG Filename Code	Calibration Scale	Conversion Factor
AEMET	IZO128N0000	NOAA04	1
AGAGE	CGO540S0011, CGO540S0013, CMO445N0011, MHD653N0011, MHD653N0013, RPB413N0000, RPB413N0011, SMO514S0014, SMO514S0016, THD441N0000	Tohoku Univ.	1.0003
BMKG & Empa	BKT500S0000	NOAA04	1
CHMI	KOS649N0000	CHMI	
CMA	WLG236N0000	NOAA04	1
CSIRO	ALT482N0003, CFA519S0003, CGO540S0003, CRI215N0000, CYA766S0001, ESP449N0003, MAA767S0003, MLO519N0003, MQA554S0003, SIS660N0003, SPO789S0003	NOAA04	1
EC	ALT482N0000, CDL453N0000, CHM449N0000, EGB444N0100, ESP449N0000, ETL454N0000, FSD449N0000, LLB454N0100, WSA443N0000	NOAA04	1
Empa	JFJ646N0000	NOAA04	1
ENEA	LMP635N0001	NOAA04	1
FMI	PAL667N0000	NOAA04	1
ISAC	CMN644N0000	NOAA04	1
JMA	MNM224N0000, RYO239N0000, YON224N0000	NOAA04	1
KMA	AMY236N0000	KRISS	
KSNU	ISK242N0000		
LSCE	AMS137S0002, BGU641N0000, LPO648N0000, PDM642N0000, PUY645N0002	NOAA83	1.0124
	FIK635N0000, MHD653N0007		
MGO	TER669N0001, TIK271N0000	NOAA04	1
MRI	TKB236N0000		0.9973
NIER	GSN233N0103	NOAA04	1
NIES	COI243N0000, HAT224N0000	NIES	0.9973

NIMR	GSN233N0001	SIO X97	
NOAA/ESRL	BRW471N0000, MLO519N0000, NOAA/ESRL flask network*	NOAA04	1
	KPA431N0001, LEF445N0001, MCM777S0001, NZL543S0001, POC935S0001, SGI354S0001, SIO432N0001	NOAA/CMDL	1.0124
RIVM	KMW653N0000	NIST	0.9973
RSE	PRS645N0000	NOAA04	1
SAWS	CPT134S0000	NOAA04	1
UBA	DEU649N0000, NGL653N0000, SSL647N0000, ZGT654N0000, ZSF647N0010, ZUG647N0000	NOAA04	1
	SNB647N0000		
Univ. Malta	GLH636N0000		

* NOAA/ESRL flask network:

ABP312S0001,ALT482N0001,AMS137S0001,ASC107S0001,ASK123N0001,AVI417N0001,AZR638N0001,BAL655N0001,BKT500S0001,
BME432N0001,BMW432N0001,BRW471N0001,BSC644N0001,CBA455N0001,CGO540S0001,CHR501N0001,CMO445N0001,CRZ146S0001,
EIC327S0001,GMI513N0001,GOZ636N0001,HBA775S0001,HPB647N0003,HUN646N0001,ICE663N0001,ITN435N0001,IZO128N0001,
KEY425N0001,KUM519N0001,KZD244N0001,KZM243N0001,LLB454N0001,LLN223N0001,LMP635N0003,MBG476N0001,MEX419N0001,
MHD653N0001,MID528N0001,MKN100S0001,MLO519N0001,NMB123S0001,NWR440N0101,OPW448N0001,OXK650N0001,PAL667N0001,
POC900N0001,POC905N0001,POC905S0001,POC910N0001,POC910S0001,POC915N0001,POC915S0001,POC920N0001,POC920S0001,
POC925N0001,POC925S0001,POC930N0001,POC930S0001,PSA764S0001,PTA438N0001,RPB413N0001,SCS903N0001,SCS906N0001,
SCS909N0001,SCS912N0001,SCS915N0001,SCS918N0001,SCS921N0001,SEY104S0001,SGP436N0001,SHM452N0001,SMO514S0001,
SPO789S0001,STM666N0001,SUM672N0001,SYO769S0001,TAP236N0001,TDF354S0001,THD441N0001,UTA439N0001,UUM244N0001,
WIS631N0001,WKT431N0001,WLG236N0001,ZEP678N0001

4. Nitrous Oxide (N₂O)

The Halocarbons and other Atmospheric Trace Species (HATS) Group of NOAA/ESRL maintains a set of standards for N₂O (Hall *et al.*, 2001) and serves as a CCL for N₂O. The NOAA-2006 N₂O scale (Hall *et al.*, 2007), revised and updated to NOAA-2006A in 2011 to deal with drifting in secondary standards, has been designated as the Primary scale for the GAW Programme. CCL compares its standards with the ones of other laboratories, including those of Environment Canada (EC) and the Australian Commonwealth Scientific and Industrial Research Organisation (CSIRO). Karlsruhe Institute of

Technology, Institute for Meteorology and Climate Research, Germany, serves as the GAW WCC for N₂O.

The SIO-98 scale is essentially equivalent to the NOAA-2006 scale, with an average difference of 0.01% over the range of 299–319 ppb; the NOAA-2000 scale can be converted to the 2006 scale by using the factor 0.999402 (Hall *et al.*, 2007). A constant ratio of 1.0017 between CSIRO and AGAGE data was used by Huang *et al.* (2008), and a factor of $1 / 1.0017 = 0.9983$ has been used in this report to convert CSIRO scale to the NOAA-2006 scale.

Table 5. Status of the standard scales of N₂O at laboratories.

Laboratory	WDCGG Filename Code	Calibration Scale	Conversion Factor
AEMET	IZO128N0000	NOAA-2006	1
AGAGE	ADR651N0010, CGO540S0011, CGO540S0012, CGO540S0013, CMO445N0010, CMO445N0011, MHD653N0011, MHD653N0013, RPB413N0000, RPB413N0010, RPB413N0011, SMO514S0014, SMO514S0015, SMO514S0016, THD441N0000	SIO 1998	1
CSIRO	ALT482N0003, CFA519S0003, CGO540S0003, CRI215N0000, CYA766S0001, ESP449N0003, MAA767S0003, MLO519N0003, MQA554S0003, SIS660N0003, SPO789S0003	NOAA-2006A	1
Empa	JFJ646N0000	SIO 1998	1
ENEA	LMP635N0001	NOAA-2006	1

ISAC	CMN644N0000	NOAA-2006	1
JMA	RYO239N0000	NOAA-2006A	1
KMA	AMY236N0000	KRISS	
MRI	MMB243N0000		
Nagoya Univ.	NGY235N0000		
NIER	GSN233N0103	NOAA-2006	1
NIES	COI243N0000, HAT224N0000	NIES 96*	1
NILU	ZEP678N0000		
NIMR	GSN233N0001	WMO X97	
NOAA/ESRL	ALT482N0001, BRW471N0001, CGO540S0001, KUM519N0001, MLO519N0001, NWR440N0001, SMO514S0001, SPO789S0001	NOAA/CMDL	0.999402
	ALT482N0004, ALT482N0006, BRW471N0003, BRW471N0005, BRW471N0011, CGO540S0009, CGO540S0014, KUM519N0002, MHD653N0008, MLO519N0005, MLO519N0006, MLO519N0011, NWR440N0003, NWR440N0004, NWR440N0011, PSA764S0000, SMO514S0008, SMO514S0009, SMO514S0011, SPO789S0005, SPO789S0006, SPO789S0011, SUM672N0002, THD441N0002	NOAA-2006	1
	BRW471N0010, MLO519N0010, NWR440N0010, SMO514S0010, SPO789S0010, SUM672N0000	NOAA-2006A	1
SAWS	CPT134S0000	NOAA/CMDL	0.999402
UBA	SSL647N0000, ZSF647N0010	SIO 1998	1

*NIES 96 N₂O scale is approximately 0.6 ppb lower than NOAA-2006 in a range between 317 to 321 ppb.
(<http://www.esrl.noaa.gov/gmd/ccgg/wmorr/results.php?rr=rr5¶m=n2o>)

5. Surface Ozone (O₃)

The National Institute of Standards and Technology (NIST) has developed and deployed Standard Reference Photometers (SRPs) in the USA and other countries. The GAW has designated SRP #2 maintained by NIST as the Primary Standard for the GAW Programme, making NIST the CCL for O₃. The Swiss Federal Laboratory for Materials Testing

and Research (Empa) maintains NIST SRP #15 as the reference and is the GAW WCC for surface ozone (Hofer *et al.*, 1998). The traceability and uncertainty of O₃ within the GAW network were reported by Klausen *et al.*, (2003). Regional Calibration Centre has been established at Observatorio Central Buenos Aires, Argentina (WMO, 2007).

Table 6. Status of surface ozone standard scales at laboratories

Laboratory	WDCGG Filename Code	Calibration Scale	Audit Empa-WCC
AEMET	IZO128N0000	WMO (NIST & Empa)	96, 98, 00, 04, 09
	DON637N0000, MHN639N0000, NIA642N0000, ROQ640N0000, SPM639N0000	NPL (U. K.)	
AQRB	ALG447N0000, BRA450N0000, CHA446N0000, EGB444N0000, ELA449N0000, EST451N0000, KEJ444N0000, SAT448N0000		
ARSO	IRB645N0000, KVK646N0000, KVV646N0000, ZRN646N0000	WMO (NIST & Empa)	
AWI	NMY770S0000		

BMKG & Empa	BKT500S0000	WMO (NIST & Empa)	99, 01, 04, 07, 08, 11
BoM & CSIRO	CGO540S0000		
	CGO540S0018	WMO (NIST & Empa)	02
CHMI	KOS649N0000	WMO (NIST & Empa)	
DEFRA	EDM655N0000		
DWD	HPB647N0000	WMO (NIST & Empa)	97, 06, 11
EMA	CAI130N0000		
Empa	JFJ646N0000, PAY646N0000, RIG646N0000	WMO (NIST & Empa)	Jungfraujoch: 99, 06
Empa & KMD	MKN100S0000	WMO (NIST & Empa)	00, 02, 05, 06, 08, 10
FMI	AHT662N0000, OUL666N0000, PAL667N0000, UTO659N0000, VIR660N0000		Pallas-Sammaltunturi: 97, 03, 07, 12
HMS	KPS646N0000	WMO (NIST & Empa)	
IM	ANG638N0000, BEJ638N0000, CAS639N0000, FUN132N0000, LIS638N0000, MVH638N0000, PEN640N0000		
INRNE	BEO642N0000	WMO (NIST & Empa)	
IOEP	DIG654N0000	WMO (NIST & Empa)	
ISAC	CMN644N0000, DCC775S0000, PYR227N0000	WMO (NIST & Empa)	Monte Cimone: 12
IVL	VDL664N0000	WMO (NIST & Empa)	
JMA	MNM224N0000, RYO239N0000, SYO769S0002, TKB236N1004, YON224N0000	WMO (NIST & Empa)	Ryori: 05
KSNU	ISK242N0000		
LA	PDM642N0001	EMD (France)	
LAMP	PUY645N0001	EMD (France)	
LVGMC	DBL656N0000, RCV656N0000, ZSN657N0000	WMO (NIST & Empa)	
MMD	DMV504N0000, TAR504N0000	WMO (NIST & Empa)	Datum Valley GAW Baseline Station: 08
NILU	ZEP678N0000	WMO (NIST & Empa)	97, 01, 05, 12
NIWA	BHD541S0000	WMO (NIST & Empa)	

NMA	FDT645N0002		
NOAA/ESRL	ARH777S0000, BMW432N0004, BRW471N0004, ICE663N0004, LAU545S0004, MLO519N0004, NWR440N0002, NWR440N0204, RPB413N0004, SMO514S0004, SPO789S0004, SUM672N0004, THD441N0004	WMO (NIST & Empa)	Mauna Loa: 03 Barrow: 08 Lauder: 10
	MCM777S0004		
NUI	MHD653N0000	NPL (U. K.)	96, 98, 02, 05, 09
ONM	ASK123N0000	WMO (NIST & Empa)	03, 07
PolyU	HKG222N0000		
RIVM	KMW653N0000		
Roshydromet	DAK654N0000, SHP659N0000		
RSE	PRS645N0000	INRIM (Italy)	
SAWS	CPT134S0000	WMO (NIST & Empa)	97, 98, 02, 06, 11
SMN	USH354S0000, USH354S0001	WMO (NIST & Empa)	98, 03, 08
SMNA	LQO322S0000, MBI764S0000, PIL331S0000, SJA349S0000, USH354S0002	WMO (NIST & Empa)	
UBA	BRT648N0000, DEU649N0000, LGB652N0000, NGL653N0000, SNB647N0000, SSL647N0000, WES654N0000, ZGT654N0000, ZSF647N0010, ZUG647N0000	WMO (NIST & Empa)	Zugspitze: 96, 97, 01 Sonnblick: 98 Zugspitze/Schneefern erhaus: 06, 11
UNA	SNL325S0000		
Univ. Malta	GLH636N0000	Tohoku Univ.	
Univ. York	CVO116N0001	NPL (U. K.)	

6. Carbon Monoxide (CO)

The Swiss Federal Laboratory for Materials Testing and Research (Empa) serves as the WCC under GAW based on its secondary standards calibrated against the

standard at NOAA/ESRL designated as the Primary Standard for GAW. The most recent scale is WMO CO X2004.

Table 7. Status of CO standard scales at laboratories

Laboratory	WDCGG Filename Code	Calibration Scale	Audit Empa-WCC
AEMET	IZO128N0000	WMO 2004 (NOAA/ESRL & Empa)	00, 04, 09
AGAGE	CGO540S0011, MHD653N0011	CSIRO	
ARSO	KVV646N0000	CHMI	
BMKG & Empa	BKT500S0000	WMO 2000 (NOAA/ESRL & Empa)	04, 07, 08, 11
CHMI	KOS649N0000	CHMI	

CSIRO	ALT482N0003, CFA519S0003, CGO540S0003, CRI215N0000, CYA766S0001, ESP449N0003, MAA767S0003, MLO519N0003, MQA554S0003, SIS660N0003, SPO789S0003	CSIRO	Cape Grim: 02
DWD	HPB647N0000	WMO (NOAA/ESRL & Empa)	97, 06, 11
EC	ALT482N0000, CDL453N0000, CHM449N0000, EGB444N0100, ESP449N0000, ETL454N0000, FSD449N0000, LLB454N0100, WSA443N0000	WMO (NOAA/ESRL & Empa)	Alert: 04
Empa	JFJ646N0000, PAY646N0000, RIG646N0000	WMO 2000 (NOAA/ESRL & Empa)	Jungfraujoch: 99,06
Empa & KMD	MKN100S0000	WMO 2000 (NOAA/ESRL & Empa)	05, 06, 08, 10
INRNE	BEO642N0000	WMO (NOAA/ESRL & Empa)	
ISAC	CMN644N0000, CMN644N0003, CMN644N0004	WMO 2004 (NOAA/ESRL & Empa)	12
JMA	MNM224N0000, RYO239N0000, YON224N0000	WMO 2000 (NOAA/ESRL)	Ryori: 05
LA	PDM642N0001	EMD (France)	
LAMP	PUY645N0001	EMD (France)	
LSCE	AMS137S0000	WMO 2004 (NOAA/ESRL & Empa)	08
NOAA/ESRL	NOAA/ESRL flask network*	WMO (NOAA/ESRL & Empa)	Mauna Loa: 03 Barrow: 08 Mt. Waliguan: 00, 04, 09
PolyU	HKG222N0000		
RIVM	KMW653N0000, KTB653N0000	NMI	
SAWS	CPT134S0000	WMO (NOAA/CMDL)	98, 02, 06, 11
SMN	USH354S0000, USH354S0001	WMO (NOAA/ESRL & Empa)	98, 03, 08
SMNA	USH354S0002	WMO (NOAA/ESRL & Empa)	
UBA	NGL653N0000, SSL647N0000, ZSF647N0001, ZUG647N0000	WMO (NOAA/CMDL)	Zugspitze: 01 Sonnblick: 98 Zugspitze/Schneefernerhaus: 06, 11
	SNB647N0000		
Univ. Malta	GLH636N0000		
Univ. York	CVO116N0001	WMO 2000 (NOAA/ESRL & Empa)	

*NOAA/ESRL flask network:

ALT482N0001, ASC107S0001, ASK123N0001, AZR638N0001, BAL655N0001, BHD541S0001, BKT500S0001, BME432N0001, BMW432N0001, BRW471N0001, BSC644N0001, CBA455N0001, CGO540S0001, CHR501N0001, CMO445N0001, CRZ146S0001, EIC327S0001, GMI513N0001, GOZ636N0001, HBA775S0001, HPB647N0003, HUN646N0001, ICE663N0001, ITN435N0001, IZO128N0001, KEY425N0001, KUM519N0001, KZD244N0001, KZM243N0001, LEF445N0001, LLN223N0001, LMP635N0003, MBC476N0001, MHD653N0001, MID528N0001, MLO519N0001, NMB123S0001, NWR440N0101, OXK650N0001, PAL667N0001, POC900N0001, POC905N0001, POC905S0001, POC910N0001, POC910S0001, POC915N0001, POC915S0001, POC920N0001, POC920S0001, POC925N0001, POC925S0001, POC930N0001, POC930S0001, POC935N0000, POC935S0001, PSA764S0001, PTA438N0001, RPB413N0001, SCS903N0001, SCS906N0001, SCS909N0001, SCS912N0001, SCS915N0001, SCS918N0001, SCS921N0001, SEY104S0001, SGP436N0001, SHM452N0001, SMO514S0001, SPO789S0001, STM666N0001, SYO769S0001, TAP236N0001, TDF354S0001, THD441N0001, UTA439N0001, UUM244N0001, WIS631N0001, WLG236N0001, ZEP678N0001

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LIST OF ABBREVIATIONS IN THE CALIBRATION AND STANDARD SCALES

AEMET	Agencia Estatal de Meteorología (Spain)
AGAGE	Advanced Global Atmospheric Gases Experiment
Aichi	Aichi Prefecture (Japan)
AIST	National Institute of Advanced Industrial Science and Technology (Japan)
AQRB	Air Quality Research Branch, Meteorological Service of Canada (Canada)
ARSO	Agencija Republike Slovenije za Okolje (Slovenia)
AWI	Alfred Wegener Institute for Polar and Marine Research (Germany)
BMKG	Agency for Meteorology, Climatology and Geophysics (Indonesia)
BoM	Commonwealth Bureau of Meteorology (Australia)
CHMI	Czech Hydrometeorological Institute (Czech Republic)
CMA	China Meteorological Administration (China)
CNR-ICES	International Centre for Earth Sciences, Consiglio Nazionale delle Ricerche (Italy)
CSIRO	Commonwealth Scientific and Industrial Research Organisation (Australia)
DEFRA	Department for Environment, Food and Rural Affairs (United Kingdom)
DNA-IAA	Dirección Nacional del Antártico-Instituto Antártico Argentino (Argentina)
DWD	Deutscher Wetterdienst (German Meteorological Service, Germany)
EARS	Environmental Agency of the Republic of Slovenia
EC	Environment Canada (Canada)
EMA	Egyptian Meteorological Authority (Egypt)
EMD	Ecole des Mines de Douai (France)
Empa	Swiss Federal Laboratories for Material Testing and Research (Switzerland)
ENEA	Italian National Agency for New Technology, Energy and the Environment (Italy)
FMI	Finnish Meteorological Institute
GAGE	Global Atmospheric Gases Experiment
GAW	Global Atmosphere Watch (WMO)
HATS	Halocarbons and other Atmospheric Trace Species Group, NOAA/ESRL
HKO	Hong Kong Observatory (Hong Kong, China)
HMS	Hungarian Meteorological Service (Hungary)
IAFMS	Italian Air Force Meteorological Service (Italy)
IEK-8	Institute for Energy and Climate Research: Troposphere (IEK-8), Research Center Juelich GmbH (Germany)
IGP	Instituto Geofísico del Perú (Peru)
IM	Instituto de Meteorologia (Portugal)
IMK-IFU	Institut für Meteorologie und Klimatologie, Atmosphärische Umweltforschung, Forschungszentrum Karlsruhe (Germany)
INRIM	Istituto Nazionale di Ricerca Metrologica (Italy)
INRNE	Institute for Nuclear Research and Nuclear Energy (Bulgaria)
IOEP	Institute of Environmental Protection (Poland)
ISAC	Istituto di Scienze dell'Atmosfera e del Clima, Consiglio Nazionale delle Ricerche (Italy)
ITM	Department of Applied Environmental Science, Stockholm University, (Sweden)
IVL	Swedish Environmental Research Institute, Göteborg (Sweden)
JMA	Japan Meteorological Agency (Japan)
KMA	Korea Meteorological Administration (Republic of Korea)
KMD	Kenya Meteorological Department (Kenya)

KRISS	Korea Research Institute of Standards and Science (Republic of Korea)
KSNU	Kyrgyz State National University (Kyrgyzstan)
KUP	Physics Institute, Climate and Environmental Physics, University of Bern (Switzerland)
LA	Laboratoire d'Aérodynamique (France)
LAMP	Laboratoire de Météorologie Physique (France)
LEGMA	Latvian Environment, Geology and Meteorology Agency (Latvia)
LSCE	Laboratoire des Sciences du Climat et de l'Environnement (France)
LVGMC	Latvian Environment, Geology and Meteorology Centre (Latvia)
MGO	Main Geophysical Observatory, Roshydromet (Russian Federation)
MPI-BGC	Max-Planck Institute (MPI) for Biogeochemistry in Jena (Germany)
MMD	Malaysian Meteorological Department
MRI	Meteorological Research Institute, JMA (Japan)
Nagoya Univ.	Nagoya University (Japan)
NIER	National Institute of Environmental Research (Republic of Korea)
NIES	National Institute for Environmental Studies (Japan)
NILU	Norwegian Institute for Air Research (Norway)
NIMR	National Institute of Meteorological Research, KMA (Republic of Korea)
NIPR	National Institute of Polar Research (Japan)
NIST	National Institute of Standards and Technology (USA)
NIWA	National Institute of Water & Atmospheric Research (New Zealand)
NMA	National Meteorological Administration (Romania)
NMI	Nederlands Meetinstituut
NOAA/ESRL	Earth System Research Laboratory, NOAA (USA)
NPL	National Physical Laboratory (United Kingdom)
NUI	National University of Ireland, Galway (Ireland)
ONM	Office National de la Météorologie (Algeria)
Osaka Univ.	Osaka University (Japan)
PolyU	Hong Kong Polytechnic University (Hong Kong, China)
RIVM	National Institute for Health and Environment (Netherlands)
Roshydromet	Federal Service for Hydrometeorology and Environmental Monitoring (Russian Federation)
RSE	Ricerca sul Sistema Elettrico (Italy)
Saitama	Saitama Prefecture (Japan)
SAWS	South African Weather Service (South Africa)
Shizuoka Univ.	Shizuoka University (Japan)
SMN(SMNA)	Servicio Meteorológico Nacional (Argentina)
Tohoku Univ.	Tohoku University (Japan)
UBA	Umweltbundesamt (Germany)
UNA	Universidad Nacional de Asunción (Paraguay)
Univ. Malta	University of Malta (Malta)
Univ. York	University of York (United Kingdom of Great Britain and Northern Ireland)
WDCGG	World Data Centre for Greenhouse Gases, operated by JMA, Japan (WMO)
WMO	World Meteorological Organization

LIST OF OBSERVATIONAL STATIONS

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
REGION I (Africa)						
Amsterdam Island	France	AMS137S00	37 47 S	77 31 E	55	CH ₄ , CO ₂
Amsterdam Island	France	AMS137S00	37 47 S	77 31 E	55	CH ₄ , CO, CO ₂ , O ₃ , VOCs
Ascension Island	United Kingdom of Great Britain and Northern Ireland	ASC107S00	7 55 S	14 25 W	54	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Assekrem	Algeria	ASK123N00	23 16 N	5 37 E	2710	O ₃
Assekrem	Algeria	ASK123N00	23 16 N	5 37 E	2710	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Cairo	Egypt	CAI130N00	30 04 N	31 16 E	35	CO ₂ , O ₃
Cape Point	South Africa	CPT134S00	34 21 S	18 28 E	230	CH ₄ , CO ₂
Cape Point	South Africa	CPT134S00	34 21 S	18 28 E	230	CH ₄ , CO, CO ₂ , N ₂ O, O ₃
Cape Verde Observatory	Cape Verde	CVO116N00	16 50 N	24 52 W	10	CO, NO, NO ₂ , NO _x , NO _y , O ₃ , VOCs
Crozet	France	CRZ146S00	46 27 S	51 51 E	120	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Funchal	Portugal	FUN132N00	32 38 N	16 52 W	58	O ₃
Gobabeb	Namibia	NMB123S00	23 34 S	15 01 E	461	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂
Izaña (Tenerife)	Spain	IZO128N00	28 18 N	16 30 W	2367	CH ₄ , CO, CO ₂ , N ₂ O, O ₃ , SF ₆
Izaña (Tenerife)	Spain	IZO128N00	28 18 N	16 30 W	2367	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Mahe Island	Seychelles	SEY104S00	4 40 S	55 10 E	7	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Mt. Kenya	Kenya	MKN100S00	0 03 S	37 17 E	3678	CO, O ₃
Mt. Kenya	Kenya	MKN100S00	0 03 S	37 17 E	3678	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO ₂ , VOCs
REGION II (Asia)						
Anmyeon-do	Republic of Korea	AMY236N00	36 31 N	126 19 E	47	CFCs, CH ₄ , CO ₂ , N ₂ O, SF ₆
Bering Island	Russian Federation	BER255N00	55 12 N	165 58 E	13	CO ₂
Cape Ochi-ishi	Japan	COI243N00	43 09 N	145 29 E	42.5	CH ₄ , CO ₂ , HCFCs, HFCs, N ₂ O
Cape Rama	India	CRI215N00	15 04 N	73 49 E	60	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Everest - Pyramid	Nepal	PYR227N00	27 57 N	86 48 E	5079	O ₃
Gosan	Republic of Korea	GSN233N00	33 16 N	126 10 E	72	CFCs, CH ₄ , CO ₂ , N ₂ O
Gosan	Republic of Korea	GSN233N01	33 10 N	126 05 E	72	CFCs, CH ₄ , CO ₂ , N ₂ O
Hamamatsu	Japan	HMM234N00	34 43 N	137 43 E	35	CO ₂
Hateruma	Japan	HAT224N00	24 03 N	123 48 E	10.8	CH ₄ , CO ₂ , HCFCs, HFCs, N ₂ O
Hok Tsui	Hong Kong, China	HKG222N00	22 12 N	114 15 E	60	CO ₂
Hok Tsui	Hong Kong, China	HKG222N00	22 12 N	114 15 E	60	CO, O ₃
Issyk-Kul	Kyrgyzstan	ISK242N00	42 37 N	76 58 E	1640	CH ₄ , CO ₂ , O ₃
Kaashidhoo	Maldives	KCO204N00	4 58 N	73 28 E	1	¹³ CO ₂ , CH ₄ , CO ₂
King's Park	Hong Kong, China	HKO222N00	22 18 N	114 10 E	65	CO ₂
Kisai	Japan	KIS236N00	36 04 N	139 33 E	13	CO ₂
Kotelny Island	Russian Federation	KOT276N00	76 00 N	137 52 E	5	CO ₂
Kyzylcha	Uzbekistan	KYZ240N00	40 52 N	66 09 E	340	CO ₂
Lulin	China	LLN223N00	23 28 N	120 52 E	2867	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂
Memambetsu	Japan	MMB243N00	43 55 N	144 11 E	32.9	N ₂ O
Mikawa-Ichinomiya	Japan	MKW234N00	34 51 N	137 25 E	50	CO ₂
Minamitorishima	Japan	MNM224N00	24 16 N	153 58 E	8	CH ₄ , CO, CO ₂ , O ₃
Mt. Dodaira	Japan	DDR236N00	36 00 N	139 10 E	840	CO ₂
Mt. Waliguan	China	WLG236N00	36 16 N	100 54 E	3810	CH ₄ , CO ₂
Mt. Waliguan	China	WLG236N00	36 16 N	100 54 E	3810	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂

LIST OF OBSERVATIONAL STATIONS (continued)

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
Nagoya	Japan	NGY235N00	35 08 N	136 58 E	35	N ₂ O
Plateau Assy	Kazakhstan	KZM243N00	43 15 N	77 52 E	2519	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Ryori	Japan	RYO239N00	39 01 N	141 49 E	260	CCl ₄ , CFCs, CH ₃ CCl ₃ , CH ₄ , CO, CO ₂ , N ₂ O, O ₃
Sary Taukum	Kazakhstan	KZD244N00	44 27 N	75 34 E	412	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Shangdianzi	China	SDZ240N00	40 38 N	117 06 E	287	CH ₄ , CO ₂
Ship between Ishigaki Island and Hateruma Island	Japan	SIH224N00	24 07 N	123 49 E	5	CO ₂
South China Sea (03N)	N/A	SCS903N00	3 00 N	105 00 E	15	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
South China Sea (06N)	N/A	SCS906N00	6 00 N	107 00 E	15	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
South China Sea (09N)	N/A	SCS909N00	9 00 N	109 00 E	15	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
South China Sea (12N)	N/A	SCS912N00	12 00 N	111 00 E	15	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
South China Sea (15N)	N/A	SCS915N00	15 00 N	113 00 E	15	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
South China Sea (18N)	N/A	SCS918N00	18 00 N	113 00 E	15	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
South China Sea (21N)	N/A	SCS921N00	21 00 N	114 00 E	15	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Suita	Japan	SUI234N00	34 49 N	135 31 E	63	CO ₂
Tae-ahn Peninsula	Republic of Korea	TAP236N00	36 43 N	126 07 E	20	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Takayama	Japan	TKY236N00	36 08 N	137 25 E	1420	CO ₂
Tiksi	Russian Federation	TIK271N00	71 35 N	128 55 E	8	CH ₄ , CO ₂
Tsukuba	Japan	TKB236N00	36 02 N	140 07 E	26	CH ₄ , CO ₂
Tsukuba	Japan	TKB236N10	36 02 N	140 07 E	25	O ₃
Ulaan Uul	Mongolia	UUM244N00	44 27 N	111 04 E	914	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Urawa	Japan	URW235N00	35 52 N	139 35 E	10	CO ₂
Yonagunijima	Japan	YON224N00	24 28 N	123 01 E	30	CH ₄ , CO, CO ₂ , O ₃

REGION III (South America)

Arembepe	Brazil	ABP312S00	12 46 S	38 10 W	0	CH ₄ , CO, CO ₂ , N ₂ O
Arembepe	Brazil	ABP312S00	12 46 S	38 10 W	0	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO ₂ , VOCs
Bird Island	United Kingdom of Great Britain and Northern Ireland	SGI354S00	54 00 S	38 02 W	30	CH ₄ , CO ₂
Easter Island	Chile	EIC327S00	27 07 S	109 27 W	50	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Huancayo	Peru	HUA312S00	12 04 S	75 31 W	3313	CO ₂
La Quiaca Observatorio	Argentina	LQO322S00	22 06 S	65 35 W	3459	O ₃
Pilar Observatorio	Argentina	PIL331S00	31 40 S	63 53 W	338	O ₃
San Julian Aero	Argentina	SJA349S00	49 18 S	67 49 W	58	O ₃
San Lorenzo	Paraguay	SNL325S00	25 22 S	57 32 W	133	O ₃
Tierra del Fuego	Argentina	TDF354S00	54 52 S	68 28 W	20	¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , H ₂ , HCFCs, HFCs, VOCs
Ushuaia	Argentina	USH354S00	54 49 S	68 17 W	18	CO, O ₃
Ushuaia	Argentina	USH354S00	54 49 S	68 17 W	18	CO, O ₃

REGION IV (North and Central America)

Alert	Canada	ALT482N00	82 27 N	62 31 W	210	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Alert	Canada	ALT482N00	82 27 N	62 31 W	210	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , N ₂ O, SF ₆

LIST OF OBSERVATIONAL STATIONS (continued)

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
Alert	Canada	ALT482N00	82 27 N	62 31 W	210	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , H ₂ , HCFCs, HFCs, N ₂ O, SF ₆ , VOCs
Algoma	Canada	ALG447N00	47 01 N	84 22 W	411	O ₃
Argyle	United States of America	AMT445N00	45 01 N	68 40 W	50	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , VOCs
Barrow	United States of America	BRW471N00	71 19 N	156 35 W	11	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , H ₂ , HCFCs, HFCs, N ₂ O, O ₃ , SF ₆ , VOCs
Bratt's Lake	Canada	BRA450N00	50 12 N	104 42 W	595	O ₃
Candle Lake	Canada	CDL453N00	53 52 N	104 39 W	489	CH ₄ , CO, CO ₂
Cape Meares	United States of America	CMO445N00	45 28 N	123 58 W	30	CCl ₄ , CFCs, CH ₃ CCl ₃ , CH ₄ , N ₂ O
Cape Meares	United States of America	CMO445N00	45 28 N	123 58 W	30	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Cape St. James	Canada	CSJ451N00	51 55 N	131 01 W	89	CO ₂
Chalk River	Canada	CHA446N00	46 04 N	77 24 W	184	O ₃
Chibougamau	Canada	CHM449N00	49 40 N	74 20 W	393	CH ₄ , CO, CO ₂
Churchill	Canada	CHL458N00	58 45 N	94 04 W	35	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO ₂ , N ₂ O
Cold Bay	United States of America	CBA455N00	55 12 N	162 43 W	25	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
East Trout Lake	Canada	ETL454N00	54 21 N	104 59 W	492	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂
Egbert	Canada	EGB444N00	44 13 N	79 46 W	253	O ₃
Egbert	Canada	EGB444N01	44 13 N	79 46 W	253	CH ₄ , CO, CO ₂ , VOCs
Estevan Point	Canada	ESP449N00	49 22 N	126 32 W	39	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Estevan Point	Canada	ESP449N00	49 22 N	126 32 W	39	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , N ₂ O, SF ₆
Esther	Canada	EST451N00	51 40 N	110 12 W	707	O ₃
Experimental Lakes Area	Canada	ELA449N00	49 40 N	93 43 W	369	O ₃
Fraserdale	Canada	FSD449N00	49 52 N	81 34 W	210	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂
Grifton	United States of America	ITN435N00	35 21 N	77 22 W	505	¹³ CO ₂ , C ¹⁸ O ₂ , CCl ₄ , CFCs, CH ₄ , CO, CO ₂ , H ₂ , N ₂ O, SF ₆
Harvard Forest	United States of America	HFM442N00	42 53 N	72 17 W	340	C ₂ Cl ₄ , CBrClF ₂ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, HCFCs, HFCs, N ₂ O, SF ₆
Kejimikujik	Canada	KEJ444N00	44 25 N	65 12 W	127	O ₃
Key Biscayne	United States of America	KEY425N00	25 40 N	80 12 W	3	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Kitt Peak	United States of America	KPA431N00	31 58 N	111 35 W	2083	CH ₄
La Jolla	United States of America	SIO432N00	32 49 N	117 16 W	14	CH ₄
La Palma	Cuba	PLM422N00	22 45 N	83 31 W	47	NO ₂
Lac La Biche	Canada	LLB454N00	54 57 N	112 27 W	540	CH ₄ , CO ₂ , VOCs
Lac La Biche (Alberta)	Canada	LLB454N01	54 57 N	112 27 W	540	CH ₄ , CO, CO ₂

LIST OF OBSERVATIONAL STATIONS (continued)

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
Mex High Altitude Global Climate Observation Center, Mexico	Mexico	MEX419N00	19 58 N	97 10 W	4560	CH ₄ , CO ₂ , VOCs
Moody	United States of America	WKT431N00	31 19 N	97 19 W	708	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄
Mould Bay	Canada	MBC476N00	76 15 N	119 19 W	58	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Niwot Ridge (C-1)	United States of America	NWR440N00	40 02 N	105 32 W	3021	C ₂ Cl ₄ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, HCFCs, HFCs, N ₂ O, O ₃ , SF ₆
Niwot Ridge (Saddle)	United States of America	NWR440N02	40 03 N	105 35 W	3528	O ₃
Niwot Ridge (T-van)	United States of America	NWR440N01	40 03 N	105 35 W	3523	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Olympic Peninsula	United States of America	OPW448N00	48 15 N	124 25 W	488	CH ₄ , CO ₂ , H ₂
Pacific Ocean (15N)	N/A	POC915N00	15 00 N	145 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (20N)	N/A	POC920N00	20 00 N	141 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (25N)	N/A	POC925N00	25 00 N	139 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (30N)	N/A	POC930N00	30 00 N	135 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (35N)	N/A	POC935N00	35 00 N	137 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CO, H ₂
Pacific Ocean (40N)	N/A	POC940N00	40 00 N	136 00 W	10	¹³ CO ₂ , H ₂
Pacific Ocean (45N)	N/A	POC945N00	45 00 N	131 00 W	10	¹³ CO ₂ , H ₂
Park Falls	United States of America	LEF445N00	45 55 N	90 16 W	868	¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , H ₂ , HCFCs, HFCs, N ₂ O, SF ₆ , VOCs
Point Arena	United States of America	PTA438N00	38 57 N	123 43 W	17	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂
Ragged Point	Barbados	RPB413N00	13 10 N	59 25 W	45	C ₂ Cl ₄ , C ₂ HCl ₃ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CHCl ₃ , HCFCs, HFCs, N ₂ O, PFCs, SF ₆ , SO ₂ F ₂
Ragged Point	Barbados	RPB413N00	13 10 N	59 25 W	45	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , O ₃
Sable Island	Canada	WSA443N00	43 55 N	60 01 W	5	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , N ₂ O, SF ₆
Saturna	Canada	SAT448N00	48 46 N	123 07 W	178	O ₃
Shemya Island	United States of America	SHM452N00	52 43 N	174 04 E	40	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Southern Great Plains	United States of America	SGP436N00	36 46 N	97 30 W	314	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O, SF ₆ , VOCs
St. Croix	United States of America	AVI417N00	17 45 N	64 45 W	3	CH ₄ , CO ₂
St. David's Head	United Kingdom of Great Britain and Northern Ireland	BME432N00	32 22 N	64 39 W	30	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂

LIST OF OBSERVATIONAL STATIONS (continued)

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
Trinidad Head	United States of America	THD441N00	41 02 N	124 09 W	120	C ₂ Cl ₄ , C ₂ HCl ₃ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CHCl ₃ , HCFCs, HFCs, N ₂ O, PFCs, SF ₆ , SO ₂ F ₂
Trinidad Head	United States of America	THD441N00	41 02 N	124 09 W	120	¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , HCFCs, HFCs, N ₂ O, O ₃ , SF ₆ , VOCs
Tudor Hill	United Kingdom of Great Britain and Northern Ireland	BMW432N00	32 16 N	64 52 W	30	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , O ₃ , VOCs
Wendover	United States of America	UTA439N00	39 52 N	113 43 W	1320	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
West Branch	United States of America	WBI441N00	41 43 N	91 21 W	241.7	¹³ CO ₂ , C ¹⁸ O ₂
REGION V (South-West Pacific)						
Baring Head	New Zealand	BHD541S00	41 24 S	174 52 E	85	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂
Baring Head	New Zealand	BHD541S00	41 24 S	174 52 E	85	¹³ CH ₄ , ¹⁴ CO ₂ , CH ₄ , CO, CO ₂ , N ₂ O, O ₃ , VOCs
Bukit Koto Tabang	Indonesia	BKT500S00	0 12 S	100 19 E	864.5	NO ₂ , SO ₂
Bukit Koto Tabang	Indonesia	BKT500S00	0 12 S	100 19 E	864.5	CH ₄ , CO, CO ₂ , O ₃
Bukit Koto Tabang	Indonesia	BKT500S00	0 12 S	100 19 E	864.5	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O, SF ₆ , VOCs
Cape Ferguson	Australia	CFA519S00	19 16 S	147 03 E	2	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Cape Grim	Australia	CGO540S00	40 40 S	144 40 E	94	CO ₂ , O ₃
Cape Grim	Australia	CGO540S00	40 40 S	144 40 E	94	C ₂ Cl ₄ , C ₂ HCl ₃ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CHCl ₃ , CO, H ₂ , HCFCs, HFCs, N ₂ O, PFCs, SF ₆ , SO ₂ F ₂
Cape Grim	Australia	CGO540S00	40 40 S	144 40 E	94	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Cape Grim	Australia	CGO540S00	40 40 S	144 40 E	94	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , H ₂ , HCFCs, HFCs, N ₂ O, SF ₆ , VOCs
Cape Kumukahi	United States of America	KUM519N00	19 31 N	154 49 W	3	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , H ₂ , HCFCs, HFCs, N ₂ O, SF ₆ , VOCs
Christmas Island	Kiribati	CHR501N00	1 42 N	157 10 W	3	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Danum Valley GAW	Malaysia	DMV504N00	4 58 N	117 49 E	426	CO ₂ , O ₃
Baseline Station						
Guam	United States of America	GMI513N00	13 25 N	144 46 E	2	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Gunn Point	Australia	GPA512S00	12 14 S	131 02 E	25	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Jakarta	Indonesia	JKR506S00	6 10 S	106 49 E	7	NO ₂ , SO ₂
Kaitorete Spit	New Zealand	NZL543S00	43 49 S	172 37 E	3	CH ₄
Lauder	New Zealand	LAU545S00	45 01 S	169 40 E	370	O ₃

LIST OF OBSERVATIONAL STATIONS (continued)

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
Macquarie Island	Australia	MQA554S00	54 28 S	158 58 E	12	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Mauna Loa	United States of America	MLO519N00	19 32 N	155 34 W	3397	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Mauna Loa	United States of America	MLO519N00	19 32 N	155 34 W	3397	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , H ₂ , HCFCs, HFCs, N ₂ O, O ₃ , SF ₆ , VOCs
Pacific Ocean (00N)	N/A	POC900N00	0 00 N	155 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (05N)	N/A	POC905N00	5 00 N	151 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (05S)	N/A	POC905S00	5 00 S	159 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (10N)	N/A	POC910N00	10 00 N	149 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (10S)	N/A	POC910S00	10 00 S	161 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (15S)	N/A	POC915S00	15 00 S	171 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (20S)	N/A	POC920S00	20 00 S	174 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (25S)	N/A	POC925S00	25 00 S	171 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (30S)	N/A	POC930S00	30 00 S	176 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific Ocean (35S)	N/A	POC935S00	35 00 S	180 00 E	10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Sand Island	United States of America	MID528N00	28 11 N	177 22 W	7.7	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Tanah Rata	Malaysia	TAR504N00	4 28 N	101 22 E	1545	O ₃
Tutuila (Cape Matatula)	United States of America	SMO514S00	14 14 S	170 34 W	42	C ₂ Cl ₄ , C ₂ HCl ₃ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CHCl ₃ , HCFCs, HFCs, N ₂ O, PFCs, SF ₆ , SO ₂ F ₂
Tutuila (Cape Matatula)	United States of America	SMO514S00	14 14 S	170 34 W	42	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , H ₂ , HCFCs, HFCs, N ₂ O, O ₃ , SF ₆ , VOCs
REGION VI (Europe)						
Adrigole	Ireland	ADR651N00	51 40 N	9 43 W	50	CCl ₄ , CFCs, CH ₃ CCl ₃ , N ₂ O
Angra do Heroismo	Portugal	ANG638N00	38 40 N	27 13 W	74	O ₃
BEO Moussala	Bulgaria	BEO642N00	42 10 N	23 35 E	2925	CO, CO ₂ , NO, NO ₂ , NO _x , O ₃ , SO ₂
Baltic Sea	Poland	BAL655N00	55 21 N	17 13 E	28	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Begur	Spain	BGU641N00	41 58 N	3 13 E	13	CH ₄ , CO ₂
Beja	Portugal	BEJ638N00	38 01 N	7 52 W	246	O ₃
Black Sea	Romania	BSC644N00	44 10 N	28 40 E	3	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Bragança	Portugal	BRG641N00	41 47 N	6 43 W	690	SO ₂
Brotjackkriegel	Germany	BRT648N00	48 49 N	13 13 E	1016	CO ₂ , O ₃
Burgas	Bulgaria	BUR642N00	42 28 N	27 28 E	16	NO ₂ , SO ₂
Castelo Branco	Portugal	CAS639N00	39 49 N	7 28 W	386	O ₃
Danki	Russian Federation	DAK654N00	54 53 N	37 47 E	140	O ₃
Deuselbach	Germany	DEU649N00	49 46 N	7 02 E	480	CH ₄ , CO ₂ , O ₃
Dobele	Latvia	DBL656N00	56 22 N	23 11 E	42	O ₃
Doñana	Spain	DON637N00	37 02 N	6 32 W	5	NO ₂ , O ₃ , SO ₂

LIST OF OBSERVATIONAL STATIONS (continued)

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
Dwejra Point	Malta	GOZ636N00	36 02 N	14 10 E	30	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Eskdalemuir	United Kingdom of Great Britain and Northern Ireland	EDM655N00	55 19 N	3 12 W	242	O ₃
Finokalia	Greece	FIK635N00	35 20 N	25 40 E	150	CH ₄ , CO ₂
Fundata	Romania	FDT645N00	45 28 N	25 18 E	1383.5	NO ₂ , SO ₂
Fundata	Romania	FDT645N00	45 28 N	25 18 E	1383.5	CO ₂ , NO ₂ , O ₃
Giordan Lighthouse	Malta	GLH636N00	36 04 N	14 13 E	160	²²² Rn, CH ₄ , CO, CO ₂ , NO, NO ₂ , NO _x , O ₃ , SO ₂
Hegyhatsal	Hungary	HUN646N00	46 57 N	16 38 E	248	CO ₂
Hegyhatsal	Hungary	HUN646N00	46 57 N	16 38 E	248	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O, SF ₆
Heimaey	Iceland	ICE663N00	63 23 N	20 16 W	100	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , O ₃ , VOCs
Hohe Warte	Austria	HHE648N00	48 15 N	16 22 E	202	NO, NO ₂ , SO ₂
Hohe Warte	Austria	HHE648N00	48 15 N	16 22 E	202	NO, NO ₂ , SO ₂
Hohenpeissenberg	Germany	HPB647N00	47 47 N	11 01 E	985	²²² Rn, CO, H ₂ O ₂ , NO, NO ₂ , NO _x , NO _y , O ₃ , PAN, ROOH, SO ₂ , VOCs
Hohenpeissenberg	Germany	HPB647N00	47 47 N	11 01 E	985	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , VOCs
Ile Grande	France	LPO648N00	48 48 N	3 35 W	10	CH ₄ , CO ₂
Iskrba	Slovenia	IRB645N00	45 34 N	14 52 E	520	NO ₂ , O ₃ , SO ₂
Ivan Sedlo	Bosnia and Herzegovina	IVN643N00	43 46 N	18 01 E	970	NO ₂ , SO ₂
Jarczew	Poland	JCZ651N00	51 49 N	21 58 E	180	NO ₂ , SO ₂
Jungfrauoch	Switzerland	JFJ646N00	46 32 N	7 59 E	3580	CO ₂
Jungfrauoch	Switzerland	JFJ646N00	46 32 N	7 59 E	3580	CH ₄ , CO, CO ₂ , H ₂ , N ₂ O, NO, NO ₂ , NO _x , NO _y , O ₃ , PAN, SF ₆ , SO ₂ , VOCs
Jungfrauoch	Switzerland	JFJ646N00	46 32 N	7 59 E	3580	C ₂ Cl ₄ , C ₂ HCl ₃ , CBrClF ₂ , CBrF ₃ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CHCl ₃ , HCFCs, HFCs, PFCs, SF ₆ , SO ₂ F ₂
K-pusztá	Hungary	KPS646N00	46 58 N	19 33 E	125	CO ₂ , NO ₂ , O ₃ , SO ₂
Kamenicki Vis	Serbia	KAM643N00	43 23 N	21 56 E	813	NO ₂ , SO ₂
Kloosterburen	Netherlands (the)	KTB653N00	53 23 N	6 25 E	0	CO, NO, NO ₂ , NO _x , SO ₂
Kollumerwaard	Netherlands (the)	KMW653N00	53 19 N	6 16 E	0	CH ₄ , CO, CO ₂ , NO, NO ₂ , NO _x , O ₃ , SO ₂
Kosetice	Czech Republic	KOS649N00	49 34 N	15 04 E	534	CH ₄ , CO, NO, NO ₂ , O ₃ , SO ₂
Kovk	Slovenia	KVK646N00	46 07 N	15 05 E	600	O ₃
Krvavec	Slovenia	KVV646N00	46 17 N	14 31 E	1720	CO, O ₃
La Cartuja	Spain	CAR637N00	37 12 N	3 36 W	720	NO ₂ , SO ₂
Lampedusa	Italy	LMP635N00	35 31 N	12 37 E	45	CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Br ₂ , CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₃ I, CH ₄ , CHCl ₃ , CO ₂ , HCFCs, HFCs, N ₂ O, SF ₆
Lampedusa	Italy	LMP635N00	35 31 N	12 37 E	45	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂
Lazaropole	The former Yugoslav Republic of Macedonia	LZP641N00	41 31 N	20 41 E	1320	NO ₂ , SO ₂
Leba	Poland	LEB654N00	54 45 N	17 31 E	2	NO ₂ , SO ₂
Lisboa / Gago Coutinho	Portugal	LIS638N00	38 46 N	9 07 W	105	O ₃
Logroño	Spain	LOG642N00	42 27 N	2 30 W	370	NO ₂ , SO ₂
Mace Head	Ireland	MHD653N00	53 19 N	9 54 W	8	O ₃

LIST OF OBSERVATIONAL STATIONS (continued)

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
Mace Head	Ireland	MHD653N00	53 19 N	9 54 W	8	CH ₄ , CO ₂
Mace Head	Ireland	MHD653N00	53 19 N	9 54 W	8	C ₂ Cl ₄ , C ₂ HCl ₃ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CHCl ₃ , CO, H ₂ , HCFCs, HFCs, N ₂ O, PFCs, SF ₆ , SO ₂ F ₂
Mace Head	Ireland	MHD653N00	53 19 N	9 54 W	8	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , H ₂ , HCFCs, HFCs, N ₂ O, SF ₆ , VOCs
Mahón	Spain	MHN639N00	39 52 N	4 19 E	78	NO ₂ , O ₃ , SO ₂
Monte Cimone	Italy	CMN644N00	44 10 N	10 41 E	2165	CO ₂
Monte Cimone	Italy	CMN644N00	44 10 N	10 41 E	2165	CH ₄ , CO, H ₂ , N ₂ O, O ₃ , SF ₆
Monte Velho	Portugal	MVH638N00	38 04 N	8 48 W	43	O ₃
Neuglobsow	Germany	NGL653N00	53 10 N	13 01 E	65	CH ₄ , CO, CO ₂ , O ₃
Noia	Spain	NIA642N00	42 43 N	8 55 W	685	NO ₂ , O ₃ , SO ₂
Ocean Station "M"	Norway	STM666N00	66 00 N	2 00 E	5	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Ocean Station Charlie	Russian Federation	STC652N00	52 45 N	35 30 W	5	CO ₂
Ocean Station Charlie	United States of America	STC654N00	54 00 N	35 00 W	6	CO ₂
Ochsenkopf	Germany	OXK650N00	50 01 N	11 48 E	1185	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , VOCs
Oulanka	Finland	OUL666N00	66 19 N	29 23 E	310	NO ₂ , O ₃ , SO ₂
Pallas-Sammaltunturi	Finland	PAL667N00	67 58 N	24 07 E	560	CH ₄ , CO ₂ , O ₃
Pallas-Sammaltunturi	Finland	PAL667N00	67 58 N	24 07 E	560	¹³ CO ₂ , C ¹⁸ O ₂ , CBrF ₃ , CH ₄ , CO, CO ₂ , VOCs
Payerne	Switzerland	PAY646N00	46 49 N	6 57 E	490	CO, NO, NO ₂ , NO _x , O ₃ , SO ₂ , VOCs
Penhas Douradas	Portugal	PEN640N00	40 25 N	7 32 W	1380	O ₃
Pic du Midi	France	PDM642N00	42 56 N	0 08 E	2877	CO, O ₃
Pic du Midi	France	PDM642N00	42 56 N	0 08 E	2877	CH ₄ , CO ₂
Plateau Rosa	Italy	PRS645N00	45 55 N	7 42 E	3480	CH ₄ , CO ₂ , O ₃
Pleven	Bulgaria	PLV643N00	43 25 N	24 36 E	64	NO ₂ , SO ₂
Plovdiv	Bulgaria	PLD642N00	42 07 N	24 45 E	179	NO ₂ , SO ₂
Puszcza Borecka/Diabler Góra	Poland	DIG654N00	54 08 N	22 04 E	157	CO ₂ , NO ₂ , O ₃ , SO ₂
Puy de Dome	France	PUY645N00	45 46 N	2 57 E	1465	CO, O ₃
Puy de Dome	France	PUY645N00	45 46 N	2 57 E	1465	CH ₄ , CO ₂
Ridge Hill	United Kingdom of Great Britain and Northern Ireland	RGL651N00	51 59 N	2 32 W	204	CH ₄ , CO ₂ , N ₂ O, SF ₆
Rigi	Switzerland	RIG646N00	46 04 N	8 26 E	1031	CO, NO, NO ₂ , NO _x , O ₃ , SO ₂ , VOCs
Roquetes	Spain	ROQ640N00	40 49 N	0 28 E	50	NO ₂ , O ₃ , SO ₂
Rucava	Latvia	RCV656N00	56 09 N	21 10 E	18	NO ₂ , O ₃ , SO ₂
San Pablo de los Montes	Spain	SPM639N00	39 32 N	4 20 W	917	NO ₂ , O ₃ , SO ₂
Schauinsland	Germany	SSL647N00	47 55 N	7 55 E	1205	CH ₄ , CO, CO ₂ , N ₂ O, NO, NO ₂ , O ₃ , PAN, SF ₆
Sede Boker	Israel	WIS631N00	31 07 N	34 52 E	400	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Semenic	Romania	SEM645N00	45 07 N	21 58 E	1432	NO ₂ , SO ₂
Shepelevo	Russian Federation	SHP659N00	59 58 N	29 07 E	4	O ₃

LIST OF OBSERVATIONAL STATIONS (continued)

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
Shetland	United Kingdom of Great Britain and Northern Ireland	SIS660N00	60 04 N	1 15 W	30	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Site J	Denmark	GRL666N00	66 30 N	46 12 W	2030	CH ₄
Sniezka	Poland	SNZ650N00	50 43 N	15 43 E	1603	NO ₂ , SO ₂
Sofia	Bulgaria	SOF642N00	42 38 N	23 22 E	586	NO ₂ , SO ₂
Sonnblick	Austria	SNB647N00	47 02 N	12 56 E	3106	CH ₄ , CO, CO ₂ , NO, NO ₂ , NO _y , O ₃
Stephansplatz	Austria	STP648N00	48 13 N	16 22 E	171	NO, NO ₂ , SO ₂
Stephansplatz	Austria	STP648N00	48 13 N	16 22 E	171	NO, NO ₂ , SO ₂
Stîna de Vale	Romania	STN646N00	46 40 N	22 37 E	1116	NO ₂ , SO ₂
Summit	Denmark	SUM672N00	72 34 N	38 28 W	3238	¹³ CO ₂ , C ¹⁸ O ₂ , CBrClF ₂ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO ₂ , HCFCs, HFCs, N ₂ O, O ₃ , SF ₆ , VOCs
Suwalki	Poland	SWL654N00	54 07 N	22 56 E	184	NO ₂ , SO ₂
Tacolneston Tall Tower	United Kingdom of Great Britain and Northern Ireland	TAC652N00	52 31 N	1 08 E	56	C ₂ Cl ₄ , C ₂ HCl ₃ , CBrClF ₂ , CBrF ₃ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CHCl ₃ , CO, CO ₂ , H ₂ , HCFCs, HFCs, N ₂ O, PFCs, SF ₆ , SO ₂ F ₂
Terceira Island	Portugal	AZR638N00	38 46 N	27 22 W	40	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Teriberka	Russian Federation	TER669N00	69 12 N	35 06 E	40	CH ₄ , CO ₂
Utö	Finland	UTO659N00	59 46 N	21 22 E	7	NO ₂ , O ₃ , SO ₂
Varna	Bulgaria	VRN643N00	43 12 N	27 55 E	41	NO ₂ , SO ₂
Viana do Castelo	Portugal	VDC641N00	41 42 N	8 48 W	16	SO ₂
Vindeln	Sweden	VDL664N00	64 15 N	19 46 E	271	O ₃
Virolahti	Finland	VIR660N00	60 31 N	27 40 E	4	NO ₂ , O ₃ , SO ₂
Waldhof	Germany	LGB652N00	52 47 N	10 46 E	74	CO ₂ , O ₃
Wank Peak	Germany	WNK647N00	47 31 N	11 09 E	1780	CO ₂ , NO _x , SO ₂
Westerland	Germany	WES654N00	54 55 N	8 19 E	12	CO ₂ , O ₃
Zabljak	Montenegro	ZBL643N00	43 08 N	19 07 E	1450	NO ₂ , SO ₂
Zavodnje	Slovenia	ZRN646N00	46 25 N	15 00 E	770	O ₃
Zeppelinfjellet (Ny-Alesund)	Norway	ZEP678N00	78 54 N	11 52 E	475	CO ₂
Zeppelinfjellet (Ny-Alesund)	Norway	ZEP678N00	78 54 N	11 52 E	475	CCl ₄ , CFCs, CH ₃ CCl ₃ , N ₂ O, O ₃ , SO ₂
Zeppelinfjellet (Ny-Alesund)	Norway	ZEP678N00	78 54 N	11 52 E	475	C ₂ Cl ₄ , C ₂ HCl ₃ , CBrClF ₂ , CBrF ₃ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CHCl ₃ , HCFCs, HFCs, PFCs, SF ₆ , SO ₂ F ₂
Zeppelinfjellet (Ny-Alesund)	Norway	ZEP678N00	78 54 N	11 52 E	475	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Zingst	Germany	ZGT654N00	54 25 N	12 43 E	1	CH ₄ , CO ₂ , O ₃
Zoseni	Latvia	ZSN657N00	57 04 N	25 32 E	182	NO ₂ , O ₃ , SO ₂
Zugspitze	Germany	ZUG647N00	47 25 N	10 58 E	2960	CO ₂
Zugspitze	Germany	ZUG647N00	47 25 N	10 58 E	2960	CH ₄ , CO, CO ₂ , NO, NO _x , NO _y , O ₃
Zugspitze / Schneefernerhaus	Germany	ZSF647N00	47 25 N	10 58 E	2656	SO ₂
Zugspitze / Schneefernerhaus	Germany	ZSF647N00	47 25 N	10 58 E	2656	CH ₄ , CO, CO ₂ , N ₂ O, NO, NO ₂ , NO _y , O ₃ , PAN, SF ₆
Ähtäri	Finland	AHT662N00	62 34 N	24 11 E	180	NO ₂ , O ₃ , SO ₂

LIST OF OBSERVATIONAL STATIONS (continued)

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
ANTARCTICA						
Arrival Heights	New Zealand	ARH777S00	77 47 S	166 40 E	184	O ₃
Arrival Heights	New Zealand	ARH777S00	77 47 S	166 40 E	184	¹³ CH ₄ , CH ₄ , CO, N ₂ O
Casey Station	Australia	CYA766S00	66 16 S	110 31 E	60	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Concordia, Dôme C	France	DCC775S00	75 05 S	123 20 E	3233	O ₃
Halley Bay	United Kingdom of Great Britain and Northern Ireland	HBA775S00	75 34 S	26 30 W	33	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
Jubany	Argentina	JBN762S00	62 13 S	58 40 W	15	CO ₂
King Sejong	Republic of Korea	KSG762S00	62 12 S	58 46 W	0	CO ₂
Marambio	Argentina	MBI764S00	64 14 S	56 37 W	198	O ₃
Mawson	Australia	MAA767S00	67 37 S	62 52 E	32	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
McMurdo Station	United States of America	MCM777S00	77 49 S	166 34 E	11	CH ₄ , O ₃
Mizuho	Japan	MZH770S00	70 42 S	44 17 E	2230	CH ₄
Neumayer	Germany	NMY770S00	70 39 S	8 15 W	42	O ₃
Palmer Station	United States of America	PSA764S00	64 55 S	64 00 W	10	¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , H ₂ , HCFCs, HFCs, N ₂ O, SF ₆ , VOCs
South Pole	United States of America	SPO789S00	89 58 S	24 48 W	2810	¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
South Pole	United States of America	SPO789S00	89 58 S	24 48 W	2810	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , C ₂ Cl ₄ , CBrClF ₂ , CBrF ₃ , CCl ₄ , CFCs, CH ₂ Cl ₂ , CH ₃ Br, CH ₃ CCl ₃ , CH ₃ Cl, CH ₄ , CO, CO ₂ , H ₂ , HCFCs, HFCs, N ₂ O, O ₃ , SF ₆ , VOCs
Syowa Station	Japan	SYO769S00	69 00 S	39 34 E	16	O ₃
Syowa Station	Japan	SYO769S00	69 00 S	39 34 E	16	CO ₂
Syowa Station	Japan	SYO769S00	69 00 S	39 34 E	16	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂ , VOCs
MOBILE STATION						
Aircraft (over Bass Strait and Cape Grim)	Australia	AIA999900				¹³ CO ₂ , CH ₄ , CO, CO ₂ , H ₂ , N ₂ O
Aircraft Observation of Atmospheric trace gases by JMA	Japan	AOA999900				CH ₄ , CO, CO ₂ , N ₂ O
Aircraft: Orleans	France	ORL999900			150	CH ₄ , CO ₂
Akademik Korolev, R/V	United States of America	AKD999900				CH ₄
Alligator liberty, M/V	Japan	ALG999900				CO ₂
Atlantic Ocean	United States of America	AOC9XXX00			10	CH ₄ , CO ₂
Comprehensive Observation Network for TRace gases by AirLiner (CONTRAIL)	Japan	EOM999900				CH ₄ , CO ₂

LIST OF OBSERVATIONAL STATIONS (continued)

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
Comprehensive Observation Network for TRace gases by AIrLiner (CONTRAIL)	Japan	EOM999900				¹³ CH ₄ , CH ₃ D
Discoverer 1983 & 1984, R/V	United States of America	DIS999900				CH ₄
Discoverer 1985, R/V	United States of America	DSC999900				CH ₄
Drake Passage	United States of America	DRP999900				¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO ₂
HATS Ocean Projects	United States of America	HOP999900				HFCs
INSTAC-I (International Strato/Tropospheric Air Chemistry Project)	Japan	INS999900				¹³ CO ₂ , CH ₄ , CO ₂
John Biscoe, R/V	United States of America	JBS999900				CH ₄
Keifu Maru, R/V	Japan	KEF999900				CO ₂
Kofu Maru, R/V	Japan	KOF999900				CO ₂
Korolev, R/V	United States of America	KOR999900				CH ₄
Long Lines Expedition, R/V	United States of America	LLE999900				CH ₄
MRI Research, 1978-1986, R/V	Japan	MRI999900				CH ₄
MRI Research, Hakuho Maru, R/V	Japan	HKH999900				CO ₂
MRI Research, Kaiyo Maru, R/V	Japan	KIY999900				CO ₂
MRI Research, Mirai, R/V	Japan	MMR999900				CO ₂
MRI Research, Natushima, R/V	Japan	NTU999900				CO ₂
MRI Research, Ryofu Maru, R/V	Japan	RFM999900				CO ₂
MRI Research, Wellington Maru, R/V	Japan	WLT999900				CO ₂
Mexico Naval H-02, R/V	United States of America	MXN999900				CH ₄
NOPACCS - Hakurei Maru -	Japan	HAK999900				TIC
Observation of Atmospheric Chemistry Over Japan	Japan	OAJ999900				CFCs, N ₂ O
Oceanographer, R/V	United States of America	OCE999900				CH ₄
Pacific Ocean	New Zealand	BSL999900				¹³ CH ₄ , CH ₄ , VOCs
Pacific Ocean	United States of America	POC9XXX00			10	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Pacific-Atlantic Ocean	United States of America	PAO999900				CH ₄
Polar Star, R/V	United States of America	PLS999900				CH ₄
Ryofu Maru, R/V	Japan	RYF999900				CFCs, CH ₄ , CO ₂ , N ₂ O, TIC

LIST OF OBSERVATIONAL STATIONS (continued)

Station	Country/Territory	Index Number	Location		Altitude (m)	Parameter
			Latitude (° ')	Longitude (° ')		
Santarem	Brazil	SAN999900				CH ₄ , CO, CO ₂ , N ₂ O, SF ₆
South China Sea	United States of America	SCS9XXX00			15	¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO, CO ₂ , H ₂
Soyo Maru, R/V	Japan	SOY999900				CO ₂
Surveyor, R/V	United States of America	SUR999900				CH ₄
The Observation of Atmospheric Methane Over Japan	Japan	OAM999900				CH ₄
The Observation of Atmospheric Sulfur Hexafluoride Over Japan	Japan	OAS999900				SF ₆
WEST COSMIC - Hakurei Maru No.2 -	Japan	HAK999901				TIC
Wakataka-Marui	Japan	WAK999900				CO ₂
Western Pacific	United States of America	WPC9XXX00			10	¹³ CH ₄ , ¹³ CO ₂ , C ¹⁸ O ₂ , CH ₄ , CO ₂
northern and western Pacific	Japan	NWP999900				N ₂ O

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Drake Passage		
Discoverer 1985, R/V		
Guam		
Grifton		
John Biscoe, R/V		

LIST OF CONTRIBUTORS (continued)

Station Country/Territory	Name	Address
	Key Biscayne	
	Korolev, R/V	
	Kitt Peak	
	Cape Kumukahi	
	Park Falls	
	Long Lines Expedition, R/V	
	McMurdo Station	
	Sand Island	
	Mauna Loa	
	Mexico Naval H-02, R/V	
	Niwot Ridge (T-van)	
	Niwot Ridge (Saddle)	
	Oceanographer, R/V	
	Olympic Peninsula	
	Pacific-Atlantic Ocean	
	Polar Star, R/V	
	Pacific Ocean	
	Palmer Station	
	Point Arena	
	South China Sea	
	Southern Great Plains	
	Shemya Island	
	La Jolla	
	Tutuila (Cape Matatula)	
	South Pole	
	Ocean Station Charlie	

LIST OF CONTRIBUTORS (continued)

Station Country/Territory	Name	Address
Surveyor, R/V		
Trinidad Head		
Wendover		
West Branch		
Moody		
Western Pacific		
(United States of America)		
NOAA/ESRL/HATS Network		
Tierra del Fuego (Argentina)	Geoffrey S. Dutton James W. Elkins Stephen A. Montzka	Halocarbons and Other Atmosphere Trace Species Group (HATS)/NOAA/ESRL Global Monitoring Division 325 Broadway R/GMD1 Boulder, CO 80305-3328, U.S.A.
Cape Grim (Australia)		
Alert (Canada)		
Summit (Denmark)		
Mace Head (Ireland)		
BACPAC 99		
BLAST1		
BLAST2		
BLAST3		
Barrow		
CLIVAR 01		
Gas Change Experiment		
Harvard Forest		
HATS Ocean Projects		
Grifton		

LIST OF CONTRIBUTORS (continued)

Station Country/Territory	Name	Address
Cape Kumukahi		
Park Falls		
Mauna Loa		
Niwot Ridge (C-1)		
PHASE I-04		
Palmer Station		
Tutuila (Cape Matatula)		
South Pole		
Trinidad Head		
(United States of America)		

NOAA/ESRL Surface Ozone Network

Ragged Point (Barbados)	Audra McClure-Begley Irina Petropavlovskikh	NOAA/ESRL Global Monitoring Division 325 Broadway, R/GMD1, Boulder, CO 80305, U.S.A
Summit (Denmark)		
Heimaey (Iceland)		
Arrival Heights Lauder (New Zealand)		
Tudor Hill (United Kingdom of Great Britain and Northern Ireland)		
Barrow McMurdo Station Mauna Loa Niwot Ridge (C-1) Niwot Ridge (Saddle) Tutuila (Cape Matatula) South Pole Trinidad Head (United States of America)		

LIST OF CONTRIBUTORS (continued)

Station Country/Territory	Name	Address
CSIRO Flask Network		
Aircraft (over Bass Strait and Cape Grim)	Ray Langenfelds	Commonwealth Scientific and Industrial Research Organisation (CSIRO)
Cape Ferguson	Paul Krummel	CSIRO Marine and Atmospheric Research
Cape Grim	Marcel van der Schoot	Private Bag 1
Casey Station	Paul Steele	Aspendale, Vic, Australia 3195
Gunn Point	Colin Allison	
Mawson		
Macquarie Island (Australia)		
Alert		
Estevan Point (Canada)		
Cape Rama (India)		
Shetland (United Kingdom of Great Britain and Northern Ireland)		
Mauna Loa		
South Pole (United States of America)		
ALE/GAGE/AGAGE Network		
Cape Grim (Australia)	Ray Wang	School of Earth and Atmospheric Sciences, Georgia Institute of Technology
	Stefan Reimann	311 Ferst Drive
	Simon O'Doherty	School of Earth and Atmospheric Sciences
Ragged Point (Barbados)	Paul Krummel	Georgia Institute of Technology
	Paul Steele	Atlanta, GA 30332-0340
	Ray F. Weiss	
Adrigole	Chris Harth	
Mace Head (Ireland)	Chris Lunder	
	Dickon Young	
	Jens Muhle	
Zeppelinfjellet (Ny-Alesund) (Norway)	Martin Vollmer	
Jungfraujoch (Switzerland)		
Cape Meares		
Tutuila (Cape Matatula)		
Trinidad Head (United States of America)		

GLOSSARY

ATMOSPHERIC SPECIES:

CCl₄	tetrachloromethane (carbon tetrachloride)
CFC-11	chlorofluorocarbon-11 (trichlorofluoromethane, CCl ₃ F)
CFC-12	chlorofluorocarbon-12 (dichlorodifluoromethane, CCl ₂ F ₂)
CFC-113	chlorofluorocarbon-113 (1,1,2-trichlorotrifluoroethane, CCl ₂ FCFClF ₂)
CFCs	chlorofluorocarbons
CH₃Cl	chloromethane (methyl chloride)
Halon-1211	chlorodifluorobromomethane (CBrClF ₂)
Halon-1301	bromotrifluoromethane (CBrF ₃)
HCFC-141b	hydrochlorofluorocarbon-141b (1,1-dichloro-1-fluoroethane, CH ₃ CCl ₂ F)
HCFC-142b	hydrochlorofluorocarbon-142b (1,1-difluoro-1-chloroethane, CH ₃ CClF ₂)
HCFC-22	hydrochlorofluorocarbon-22 (chlorodifluoromethane, CHClF ₂)
HCFCs	hydrochlorofluorocarbons
HFC-134a	hydrofluorocarbon-134a (1,1,1,2-tetrafluoroethane, CH ₂ FCF ₃)
HFC-152a	hydrofluorocarbon-152a (1,1-difluoroethane, CHF ₂ CH ₃)
HFCs	hydrofluorocarbons
CH₄	methane
CH₃CCl₃	trichloroethane (methyl chloroform)
CO	carbon monoxide
CO₂	carbon dioxide
N₂O	nitrous oxide
NO	nitrogen monoxide
NO₂	nitrogen dioxide
NO_x	nitrogen oxides
O₃	ozone
SF₆	sulphur hexafluoride
SO₂	sulphur dioxide
VOCs	volatile organic compounds

UNITS:

ppm	parts per million
ppb	parts per billion
ppt	parts per trillion

Others:

ENSO	El Niño-Southern Oscillation
M/V	merchant vessel
R/V	research vessel

LIST OF WMO/WDCGG PUBLICATIONS

DATA REPORTING MANUAL:

WDCGG No. 1 January 1991

WMO WDCGG DATA REPORT:

(period of data accepted)

WDCGG No. 2 Part A	October	1992	October	1990	~	August	1992
WDCGG No. 2 Part B	October	1992	October	1990	~	August	1992
WDCGG No. 3	October	1993	September	1992	~	March	1993
WDCGG No. 5	March	1994	April	1993	~	September	1993
WDCGG No. 6	September	1994	September	1993	~	March	1994
WDCGG No. 7	March	1995	April	1994	~	December	1994
WDCGG No. 9	September	1995	January	1995	~	June	1995
WDCGG No.10	March	1996	July	1995	~	December	1995
WDCGG No.11	September	1996	January	1996	~	June	1996
WDCGG No.12	March	1997	July	1996	~	November	1996
WDCGG No.14	September	1997	December	1996	~	June	1997
WDCGG No.16	March	1998	July	1997	~	December	1997
WDCGG No.17	September	1998	January	1998	~	June	1998
WDCGG No.18	March	1999	July	1998	~	December	1998
WDCGG No.20	September	1999	January	1999	~	June	1999
WDCGG No.21	March	2000	July	1999	~	December	1999
WDCGG No.23	September	2000	January	2000	~	June	2000
WDCGG No.25	March	2001	July	2000	~	December	2000

WMO WDCGG DATA CATALOGUE:

WDCGG No. 4	December	1993
WDCGG No.13	March	1997
WDCGG No.19	March	1999
WDCGG No.24	March	2001

WMO WDCGG DATA SUMMARY:

WDCGG No. 8	October	1995
WDCGG No.15	March	1998
WDCGG No.22	March	2000
WDCGG No.26	March	2002
WDCGG No.27	March	2003
WDCGG No.28	March	2004
WDCGG No.29	March	2005
WDCGG No.30	March	2006
WDCGG No.31	March	2007
WDCGG No.32	March	2008
WDCGG No.33	March	2009
WDCGG No.34	March	2010
WDCGG No.35	March	2011
WDCGG No.36	March	2012
WDCGG No.37	March	2013

WMO WDCGG CD-ROM:

(period of data accepted)

CD-ROM No. 1	March	1995	October	1990	~	December	1994
CD-ROM No. 2	March	1996	October	1990	~	June	1995
CD-ROM No. 3	March	1997	October	1990	~	June	1996
CD-ROM No. 4	March	1998	October	1990	~	December	1997
CD-ROM No. 5	March	1999	October	1990	~	December	1998

CD-ROM No. 6	March	2000	October	1990	~	December	1999
CD-ROM No. 7	March	2001	October	1990	~	December	2000
CD-ROM No. 8	March	2002	October	1990	~	January	2002
CD-ROM No. 9	March	2003	October	1990	~	December	2002
CD-ROM No.10	March	2004	October	1990	~	December	2003
CD-ROM No.11	March	2005	October	1990	~	December	2004
CD-ROM No.12	March	2006	October	1990	~	December	2005
CD-ROM No.13	March	2007	October	1990	~	November	2006
CD-ROM No.14	March	2008	October	1990	~	November	2007

WMO WDCGG DVD:

(period of data accepted)

DVD No. 1	March	2009	October	1990	~	November	2008
DVD No. 2	March	2010	October	1990	~	November	2009
DVD No. 3	March	2011	October	1990	~	November	2010
DVD No. 4	March	2012	October	1990	~	November	2011
DVD No. 5	March	2013	October	1990	~	November	2012
DVD No. 6	March	2014	October	1990	~	November	2013